

**Integrated Ocean Drilling Program
Expedition 311 Preliminary Report**

Cascadia Margin Gas Hydrates

28 August–28 October 2005

Expedition 311 Scientists

PUBLISHER'S NOTES

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ABSTRACT

A transect of four sites (U1325, U1326, U1327, and U1329) across the Northern Cascadia margin was cored during Integrated Ocean Drilling Program Expedition 311 to study gas hydrate occurrences and formation models for accretionary complexes. In addition to the transect sites, a fifth site (U1328), representing a cold vent with active fluid and gas flow, was visited.

The four transect sites represent different stages in the evolution of gas hydrate across the margin from the earliest occurrence on the westernmost first accreted ridge (Site U1326) to its final stage at the eastward limit of gas hydrate occurrence on the margin in shallower water (Site U1329).

Logging while drilling/measurement while drilling carried out at the start of the expedition prior to coring provided a set of measurements that guided subsequent coring and special tool deployments at all five sites. Additional wireline logging at each site and two vertical seismic profiles at Sites U1327 and U1328 were completed. A total of 1217.76 m of sediment core was recovered using the advanced piston corer and extended core barrel systems, interspersed with 24 (16 successful) pressure core sampler runs for onboard degassing experiments and 19 Fugro piston corer/HYACE deployments; 4 of these pressure cores were stored under in situ pressure for subsequent shore-based studies.

Indirect evidence of the presence of gas hydrate included increased electrical resistivities and *P*-wave velocities on downhole logs and low-salinity interstitial water anomalies, numerous infrared cold spots, and decreases in void gas C_1/C_2 ratios, as well as gas hydrate-related sedimentological moussy/soupy textures in recovered cores. Gas hydrate was also observed directly in the recovered cores, and >30 gas hydrate samples were preserved in liquid nitrogen for shore-based studies. The combined observations show that gas hydrate mainly occurs within coarser-grained turbidite sands and silts.

The occurrence of gas hydrate appears to be controlled by several key factors, and the concentration of gas hydrate changes significantly as those factors vary in the sediments along the margin. The key controlling factors are (1) local methane solubility linked with pore water salinity, (2) fluid/gas advection rates, and (3) availability of suitable host material (coarse-grained sediments). In the previous model for gas hydrate formation in an accretionary margin, the highest concentrations of gas hydrate were expected to occur localized near the base of the gas hydrate stability zone above

the bottom-simulating reflector (BSR), with concentrations gradually decreasing upward as a result of pervasive fluid advection from overall tectonically driven fluid expulsion. However, the results of Expedition 311 show that this model is too simple and that there are additional overprinting factors. Although evidence for widespread gas hydrate-related BSRs was observed in the data, by far the largest concentrations of gas hydrate were observed at the top of the gas hydrate occurrence zone, at a point where the amount of methane in the pore fluid exceeds the local methane solubility threshold. This was especially observed at Sites U1326 and U1327, where gas hydrate deposits several tens of meters thick occur at a shallow depth of ~100 meters below seafloor (mbsf) with concentrations locally exceeding 80% of the pore volume. Another site of very large gas hydrate concentrations was found at the cold vent Site U1328, where beds of massive gas hydrate occur within the top ~40 mbsf with concentrations exceeding 80% of the pore space as a result of focused fluid/gas migration from underneath.

INTRODUCTION

The area of investigation for Expedition 311 is the accretionary prism of the Cascadia subduction zone (Fig. F1). The Juan de Fuca plate converges nearly orthogonally to the North American plate at a present rate of ~45 mm/y (e.g., Riddihough, 1984). Seaward of the deformation front, the Cascadia Basin consists of pre-Pleistocene hemipelagic sediments overlain by rapidly deposited Pleistocene turbidites for a total sediment thickness of ~2500 m. Most of the incoming sediment is scraped off the oceanic crust and folded and thrust upward to form elongated anticlinal ridges with elevations as high as 700 m above the adjacent basins. The thrust faults near the deformation front penetrate nearly the entire sediment section (Davis and Hyndman, 1989).

The primary objective of Expedition 311, on the northern Cascadia margin, is to constrain geologic models for the formation of gas hydrate in subduction zone accretionary prisms. This expedition follows the goals for gas hydrate drilling as proposed by the Ocean Drilling Program (ODP) Gas Hydrates Program Planning Group:

- Study the formation of natural gas hydrate in marine sediments.
- Determine the mechanism of development, nature, magnitude, and global distribution of gas hydrate reservoirs.
- Investigate the gas transport mechanism, and migration pathways through sedimentary structures, from site of origin to reservoir.
- Examine the effect of gas hydrate on the physical properties of the enclosing sediments, particularly as it relates to the potential relationship between gas hydrates and slope stability.
- Investigate the microbiology and geochemistry associated with gas hydrate formation and dissociation.

The specific objectives of this expedition are to test gas hydrate formation models and constrain model parameters, especially models that account for the formation of concentrated gas hydrate occurrences through upward fluid and methane transport. These objectives require (1) high-quality data on the vertical concentration distribution of gas hydrate and free gas and variation landward in the accretionary prism and (2) estimates of the vertical fluid and methane fluxes through the sediment section as a function of landward distance from the deformation front. The study concentrates on the contrast between dispersed pervasive upward flow and focused flow of fluid and methane in fault zones. The pervasive permeability may be on the sediment grain

scale, on a centimeter scale (the scaly fabric observed in previous ODP clastic accretionary prism cores), or in closely spaced faults.

Natural gas hydrate occurs in marine continental slope and onshore Arctic permafrost environments. The Arctic occurrences can exhibit very high gas hydrate concentrations but appear to contain less total gas than marine gas hydrate occurrences. Recent studies have indicated that the largest occurrences of gas hydrate may lie in nearly horizontal layers several hundred meters beneath the seafloor of continental slopes, especially in the large subduction zone accretionary sedimentary prisms. Gas hydrate and underlying free gas produce the ubiquitous bottom-simulating reflectors (BSRs) along numerous continental margins of the world. Gas hydrates do occur on passive margins, but they are less common and appear to be usually at lower concentrations.

The two marine gas hydrate areas that have received the most detailed scientific study, including previous drilling during the ODP, are the Blake Ridge region off the east coast of the United States (a passive margin setting) and the Cascadia margin off the coast of Oregon and Vancouver Island (subduction zone accretionary prism). Important new information on Arctic gas hydrate occurrences has been obtained from the Mallik-1998 and Mallik-2002 drilling projects in northern Canada (Dallimore et al., 1999; Dallimore and Collett, 2005). If our conclusions for the occurrence and distribution of gas hydrate in nature are correct, gas hydrate formed within accretionary prisms is the most important both for understanding the geologic controls on the occurrence of gas hydrate and for the role that natural gas hydrate plays in the global carbon cycle. Within accretionary prisms, the largest amount of gas hydrate is predicted to occur in a widespread layer located just above the BSR (Hyndman and Davis, 1992). However, also important is focusing a portion of the upward methane flux into localized migration conduits or channels to form concentrated near-seafloor gas hydrate accumulations. The amount of gas hydrate in the widespread layer above the BSR, compared to that in local concentrations near the seafloor, is especially important for understanding the response of marine gas hydrate to climate change. Near-seafloor gas hydrate accumulations will respond much more quickly to ocean temperature changes compared to gas hydrate several hundred meters below the seafloor. For the region of ODP Sites 889/890, Taylor et al. (2002) calculated that a 30 m thick hydrate deposit lying at the base of the stability field would dissociate due to seabed warming in ~8000 y. However, near-seafloor gas hydrate could dissociate much more quickly and be much more responsive to human-induced global warming (e.g., Wood et al., 2002). Off Vancouver Island a gas hydrate-related BSR occurs in a 30 km wide band parallel to the coast beneath much of the continental slope (Fig. F1). Gas hy-

hydrate is thought to be concentrated in a layer 50–100 m thick, just above the base of the gas hydrate stability field, which is located 200–300 m below seafloor (mbsf) based on earlier interpretations (see, e.g., Hyndman et al., 2001, for a summary of these findings). The gas hydrate concentrations, estimated from downhole logging data collected during ODP Leg 146, may reach ~30% of pore space. The surveys and studies that have been carried out and the evidence for the presence and content of gas hydrate are summarized in two comprehensive review articles (Hyndman et al., 2001; Spence et al., 2000).

Gas hydrate concentrations were previously estimated from downhole acoustic and resistivity logs, multichannel seismic (MCS) data, vertical seismic profile (VSP) velocities, and pore water freshening (e.g., Yuan et al., 1996, 1999; Hyndman et al., 1999). Gas hydrate concentrations determined from the different methods vary slightly but were estimated to be on average 20%–35% of the pore space over a 100 m thick interval above the BSR. These are high concentrations for a marine gas hydrate occurrence and are not observed on other margins. ODP drilling and logging at the Blake Ridge offshore South Carolina (passive margin environment) showed gas hydrate concentrations that were on average <10% (Paull, Matsumoto, Wallace, et al., 1996). Results from Leg 204 at southern Hydrate Ridge located on the southern Cascadia margin offshore Oregon, with a very similar tectonic environment to offshore Vancouver Island, showed very low gas hydrate concentrations of <5% except for the unusual summit of Hydrate Ridge (Trehu et al., 2004). An earlier approach by Ussler and Paull (2001) for interpreting chlorinity data from Leg 146 by using a smooth chlorinity baseline suggested lower gas hydrate concentrations at Sites 889/890 but apparently contradicted other geophysical results using electrical resistivity and seismic velocity. In preparation for Expedition 311, the gas hydrate concentrations along the northern Cascadia margin were recalculated using Leg 146 acoustic/electrical resistivity logs and pore water chlorinity/salinity data (Riedel et al., 2005). New estimates show that the concentrations could alternatively be between 5% and 10% on average of the pore volume from ~130 mbsf to the BSR (~230 mbsf).

No conclusive decision can be made using interstitial water chlorinity data from Leg 146 to assess the occurrence of gas hydrate without a more complete analysis of the baseline chlorinity data and without additional pore water chemistry data (lithium and strontium data) to confirm the model by Torres et al. (2004) derived from Leg 204 results. Therefore, gas hydrate concentrations along the northern Cascadia margin as estimated from chlorinity data may either be as low as a few percent, as suggested earlier by Ussler and Paull (2001), or as high as 40%, previously suggested by Hyndman

et al. (1999). Gas hydrate concentrations also were calculated from the resistivity data using Archie's equation. Archie's equation consists of several empirical parameters (referred to as a , m , and n), which were determined to be $a = 1.4$, $m = 1.76$, and $n = 1.0$ by Hyndman et al. (1999) using the Leg 146 core data at Sites 888 and 889/890. However, Collett (2000) redefined these parameters to $a = 1.0$, $m = 2.8$, and $n = 1.9$ by using the Site 889 resistivity and neutron porosity logs. The differences in these empirical parameters result in highly different estimated gas hydrate concentrations. The Leg 146 logging data are of relatively poor quality and neither the new parameters nor the previous analyses fit all data well, and a large uncertainty remains in the results. However, results from Leg 204, obtained from a tectonically similar environment, suggest that the Archie parameters by Collett (2000) may be preferable.

Finally, new reference velocity profiles were calculated from the Site 889 porosities using various published empirical relations between porosity and velocity (Jarrard et al., 1995; Hyndman et al., 1993) and the Lee et al. (1993) weighted equation. All the newly proposed baselines are significantly shifted toward higher seismic velocities relative to the former baseline defined by Yuan et al. (1996) and Hyndman et al. (2001). Thus, they result in gas hydrate concentrations that are only between 5% and 10% of the pore volume. However, significant uncertainty remains in the applicability of the empirical parameters for each of the individual equations. If the model by Torres et al. (2004) proposed for Hydrate Ridge (Leg 204) is applicable to Site 889 and if the Archie parameters by Collett (2000) are correct, these two methods (although individually very uncertain) confirm results from the acoustic velocity analyses of Lee et al. (1993). Thus, the total gas hydrate concentration at Sites 889/890 may be much lower than previously assumed.

A general model for deep-sea gas hydrate formation by removal of methane from upwardly expelled fluids was developed earlier for the Expedition 311 area (Fig. F2) (Hyndman and Davis, 1992). Mainly microbial methane, inferred to be produced over a thick sediment section, migrates vertically to form gas hydrate when it enters the stability field. The gas hydrate concentration is predicted to be greatest just above the BSR. A model has also been proposed for how free gas and the resulting BSR will be formed as the base of the gas hydrate stability moves upward due to post-Pleistocene seafloor warming, uplift, and sediment deposition (e.g., Paull and Ussler, 1997; von Huene and Pecher, 1998). In addition, physical and mathematical models have been developed for the formation of gas hydrate from upward methane advection and diffusion (e.g., Xu and Ruppel, 1999).

Recently, evidence for focused fluid/gas flow and gas hydrate formation has been identified on the Vancouver margin. The most studied site is an active cold vent field associated with near-surface gas hydrate occurrences close to Sites 889/890 (e.g., Riedel et al., 2002). Studies include high-resolution subbottom profiling, three-dimensional (3-D) seismic surveys, piston coring, and ocean-bottom video surveying and sampling with the remotely operated vehicle *ROPOS*. These vents represent fault-related conduits for focused fluid and/or gas migration associated with massive gas hydrate formation within the fault zone and represent, therefore, the opposite mechanism to the widespread fluid flow. It is thus far unknown how important these cold vents are in the total budget of fluid flow in an accretionary prism. Drilling at the vent field will help constrain the significance of fault-related fluid and/or gas flow.

During Expedition 311, a series of four holes was drilled along a margin-perpendicular transect (Figs. F3, F4) representing different stages in the evolution of the gas hydrate stability zone; an additional fifth site was established to examine an apparent active cold vent.

At the southwestern end of the transect, Site U1326 (prospectus Site CAS-03C) is located on top of the first uplifted ridge, followed by Site U1325 (prospectus Site CAS-02C) located in the first slope-sediment basin. The third Site U1327 (prospectus Site CAS0-01B) is located near Leg 146 Site 889. The fourth site of the transect, Site U1329 (prospectus Site CAS-05D), represents the eastward limit of the gas hydrate occurrence on the margin. The additional fifth site is located at an active cold vent (Site U1328; prospectus Site CAS-06A).

Recently acquired bathymetry data reveal a collapse structure near the originally proposed primary Site CAS-03B (Fig. F5) where we now expect a more complicated geologic history. We decided to switch the former alternate location Site CAS-03C to primary status to avoid coring directly into the slump feature.

The original proposal 553-Full2 “Cascadia Gas Hydrate” included long-term monitoring experiments using Circulation Obviation Retrofit Kit (CORK)-II, Advanced CORK (ACORK), modular formation dynamic testing (MDT), and distributed temperature sensing (DTS). An extra site for a dual-hole hydrogeologic ACORK experiment was proposed near Site 889 (prospectus Site CAS-01C). With this shortened expedition approach under the Integrated Ocean Drilling Program (IODP), the long-term monitoring experiments were postponed until a second (yet unscheduled) expedition. Other components of the original proposal that had to be deferred included coring, wireline

logging, and logging-while-drilling operations at prospectus Sites CAS-04B and CAS-05B and a walk-away VSP at prospectus Site CAS-01B.

At each site we tried to accomplish a three-hole conventional coring, specialized coring, and logging approach to maximize the scientific objectives. The first hole (Hole A) was always dedicated to logging-while-drilling/measurement-while-drilling (LWD/MWD) operations. All LWD/MWD holes were drilled within the first week of the expedition, followed by coring and conventional wireline operations. The second hole (Hole B) in most cases was for continuous coring, temperature measurements, and pressure core sampler (PCS) pressure coring to establish complete downhole profiles of gas hydrate proxies (such as chlorinity, infrared images, etc.). The third hole (Hole C), when included, was dedicated to special tool deployments, especially the two HY-ACINTH pressure coring systems, the Fugro Pressure Corer (FPC) and HYACE Rotary Corer (HRC), and additional PCS pressure coring.

During this expedition, we conducted LWD/MWD operations prior to coring each site. It was essential to know before deploying the available pressure core systems what the gas hydrate concentrations and vertical distribution were to choose the optimum depths for pressure coring operations. Because these holes were drilled prior to coring, the LWD/MWD data had to be monitored to detect gas entering the hole, since the standard method of hydrocarbon safety management during IODP is core gas ratio measurements. The primary measurement used during Expedition 311 for gas monitoring during LWD/MWD operations was the annular pressure while drilling (APWD) measured by the EcoScope tool in the borehole annulus (the space between the drill string and the borehole wall). On the basis of a calculation of the effect of free gas on the borehole fluid density, it was determined that a pressure decrease of >100 psi from the general trend of fluid pressure indicated that a significant amount of gas had been released into the drilling fluid, requiring some preventative action. For example, a pressure decrease of 100 psi corresponds to 25% gas saturation in a borehole drilled to 300 mbsf. It was also decided to monitor sudden pressure increases of >100 psi, which had been reported as precursors to gas flow into the annulus (Aldred et al., 1998). Figure F6 shows the results for LWD/MWD APWD measurements acquired during the expedition. For the entire LWD/MWD operations, no APWD response was >30 psi above or below the baseline, and hence, no corrective actions were required under the adopted LWD/MWD safety protocol. The observed positive APWD excursions were all related to drilling-related difficulties, such as an increase in pumping rates during drilling hard intervals or cuttings restricting flow in the annulus. Interestingly, small negative excursions were noted at some sites at depths below the

BSR (e.g., Site U1328) that were also associated with loss coherence of the sonic waveforms acquired by the SonicVision tool, which may be the result of the presence of free gas.

RESULTS

Site U1325

Site U1325 (prospectus Site CAS-02C) is located near the southwestern end of the margin-perpendicular transect established during Expedition 311 and is within a major slope basin that developed eastward of the deformation front behind a steep ridge of accreted sediments (Fig. F4). Bathymetry data show that the seafloor in the western part of this slope basin is relatively flat with water depths of ~2200 m. Around Site U1325 the seafloor becomes gradually shallower before it rises rapidly to the east to form the plateau of the second main accreted ridge at water depths of ~1200 m, on which Sites U1327 and 889 are located. A BSR is clearly visible in the eastern part of the slope basin (Fig. F7), but it fades to the west (common depth points [CDPs] 1180–1280 along MCS Line 89-08). The BSR also shows the typical frequency-dependent reflection strength pattern (Chapman et al., 2002) as observed at all other sites. At Site U1325 the BSR is less strong than at the core of a buried ridge of accreted sediments, which is located ~700 m west of this site. The BSR appears to be at $\sim 230 \pm 5$ mbsf based on seismic travelttime-depth conversion, using an average *P*-wave velocity of 1636 m/s. This average velocity was determined at Site 889 to match observed travelttime depths of the BSR and VSP data (MacKay et al., 1994). We acknowledge that this velocity may not be appropriate in this setting. However, lack of good control in seismic interval velocities from MCS semblance analyses of Lines 89-08 and PGC9902_ODP7 prohibited accurate depth migration.

The primary research objectives for this site are linked to the transect-concept of this expedition. The objectives include

1. Studying the distribution of gas hydrates,
2. Defining the nature of the BSR,
3. Developing baseline geochemical and microbiological profiles, and
4. Obtaining data needed to ground-truth remotely acquired imaging techniques such as seismic or controlled-source electromagnetic (CSEM) surveys.

The slope basin is expected to show a different geochemical regime and related geophysical properties than the uplifted ridges of accreted sediments. A summary of major logging and core data is shown in comparison with two key seismic data sets (coincident but of different seismic frequency) in Figures **F8** and **F9**.

Four holes were occupied at Site U1325. Hole U1325A was dedicated for the LWD/MWD program to a total depth (TD) of 300 mbsf (Table **T1**). Hole U1325B was spudded with the advanced piston corer (APC) system, but problems arose due to thick sand accumulation, causing a switch to the extended core barrel (XCB) for one core section and a switch back to APC. The hole was then advanced by combination of APC, XCB, and pressure coring to 206.5 mbsf. Interspersed with the XCB cores was a Davis-Villinger Temperature Probe (DVTP) run at 140.5 mbsf, which yielded high-quality temperature data. Five pressure cores were deployed in Hole U1325B, but all attempts failed to retrieve sediments under pressure. The last run with the PCS resulted in a stuck tool, which led to the abandonment of the hole. Hole U1325C was drilled to 188 mbsf, where core recovery in Hole U1325B began to deteriorate. Coring operations resumed with the XCB system, interspersed with two pressure coring runs (one failed FPC run and one successful run with the PCS) and two DVTP deployments, deepening the hole to a TD of 304.3 mbsf. Because of degrading weather conditions, it was decided that two separate logging runs would be attempted without calipers to reduce the potential risk of damage to the tool strings. The first run included the phasor Dual Induction Tool (DIT) and the Hostile Environment Gamma Ray Sonde (HNGS). The second run included the Dipole Sonic Imager (DSI), the Scintillation Gamma Ray Tool (SGT), and the Temperature/Acceleration/Pressure (TAP) tool. We were only able to lower the first tool string to 259.8 mbsf, 44.5 m shallower than total depth for coring, and the hole was successfully logged followed by some difficulty re-entering the drill pipe. The second tool string (sonic without the Formation Micro-Scanner [FMS]) was deployed and was only able to reach a TD of 185.8 mbsf; the available hole was logged successfully. The repeat pass of the sonic tool string reached 183.0 mbsf. To ensure that a mudline core was obtained at this site, Hole U1325D was established with one APC core only (4.69 m recovery), indicating a seafloor depth of 2193.2 meters below sea level (mbsl; 2204.8 meters below rig floor [mbrf]). Analyses of the pore water showed that the first core in Hole U1325C had missed the mudline by ~3 m.

The sediments cored at Site U1325 are mostly Quaternary in age, based on the examination of diatoms from Holes U1325B and U1325C. The 304.3 m thick sedimentary

section cored at Site U1325 was divided into four lithostratigraphic units based on visual inspection of the recovered cores and analysis of smear slides (Fig. F7).

Lithostratigraphic Subunit IA (0–52.50 mbsf in Hole U1325B and 0–4.69 mbsf in Hole U1325D) is characterized by very abundant, thick, coarse-grained sand layers within fine-grained (clay and silty clay) detrital interlayers. The thickness of the sand layers suggests the depositional environment of a distributary channel within the slope basin with mass transport parallel to the uplifted bounding ridges. Sedimentation rate for the upper 129 mbsf of sediments (lithostratigraphic Units I, II, and upper part of Unit III) was estimated at 43 cm/k.y., the highest sedimentation rate calculated for any site occupied during Expedition 311. Lithostratigraphic Subunit IB (24.00–52.23 mbsf in Hole U1325B) is composed of fine-grained (clay to silty clay) detrital sediments with few thin silty/sandy interlayers from turbidites. The occurrence of only marine diatoms suggests that the terrigenous sediment flux was somehow restricted. In addition, the occurrence of abundant bioturbation and sulfide mottles suggests more hemipelagic dominated sedimentation. Mousselike texture related to the presence of gas hydrate is present at depths as shallow as 26 mbsf. The boundary between lithostratigraphic Subunits IA and IB is marked by a seismic horizon (unconformity) that can be traced for ~1.5 km around Site U1325 (Fig. F7). It separates undeformed seafloor-parallel slope sediments from underlying deformed and slightly faulted sediments. It is likely that at Site U1325 we penetrated a fault within lithostratigraphic Unit II near the base of the unit (Fig. F7).

Lithostratigraphic Unit II (52.23–102.30 mbsf in Hole U1325B) is characterized by fine-grained (clay to silty clay) detrital sediments with intervals of silty/sandy interlayers. We interpret the coarser interlayers as turbiditic deposits. Their frequent occurrence might indicate times of active tectonism. The boundary between lithostratigraphic Units I and II is marked by the absence of diatoms, bioturbation, less intense sulfide mottles, and a sharp decrease in the occurrence of sponge spicules. Mousselike texture related to the presence of gas hydrate was also observed in lithostratigraphic Unit II.

Lithostratigraphic Unit III (102.30–189.44 mbsf in Hole U1325B and 188.80–192.08 mbsf in Hole U1325C) is characterized by fine-grained (clay to silty clay) detrital sediments with abundant sponge spicules and diatoms. We interpret this interval as mixed hemipelagic-turbiditic sediments. The presence of authigenic carbonate cement suggests that diagenetic processes are active in lithostratigraphic Unit III. Soupy and mousselike sediment textures related to the presence of gas hydrates are present

in lithostratigraphic Unit III within the depth interval from -120 to -150 mbsf. The boundary between lithostratigraphic Units III and IV is marked by the sudden decrease in diatoms.

Lithostratigraphic Unit IV (197.40–206.56 mbsf in Hole U1325B and 198.40–304.30 mbsf in Hole U1325C) is characterized by fine-grained (clay to silty clay) detrital sediments with few silty/sandy interlayers from turbiditic deposits. The poor recovery in this unit limits the observation of sedimentary structures; however, the appearance of several coarser-grained layers indicates low-energy turbidity currents.

The stratigraphic section logged at Site U1325 was divided into three “logging units” based on obvious changes in the LWD/MWD and wireline gamma ray, density, electrical resistivity, and acoustic measurements. There is no obvious correlation between the logging units and lithostratigraphic units.

Logging Unit 1 (0–122 mbsf) is characterized by a well-defined gradual increase in density with depth and a corresponding decrease in porosity. This increase in density is matched by a corresponding increase of resistivity with depth, from ~1 Ωm near the seafloor to ~1.5 Ωm at 122 mbsf. The boundary between logging Units 1 and 2 corresponds to the 0.3 Ma age boundary noted in biostratigraphy.

Logging Unit 2 (122–260 mbsf) is characterized by a uniform density that averages ~1.9 g/cm³. In contrast, the resistivity logs in Unit 2 show alternating thin intervals of high and low resistivity, spanning the range 1–15 Ωm . These high and low resistivities likely correspond to intercalated layers that have high and low gas hydrate concentrations, respectively.

Logging Unit 3 (260–350 mbsf) displays a uniform background resistivity (~1 Ωm) and density (~2 g/cm³). The most striking feature of this unit is the presence of several borehole enlargements.

LWD/MWD-derived resistivity-at-the-bit (RAB) images from Hole U1325A suggest that gas hydrates occur concentrated in thin sand layers within the depth interval between 173 and 240 mbsf (logging Unit 2). The LWD/MWD porosity and resistivity logs from Hole U1325A further show that it is a very heterogeneous gas hydrate-bearing section, composed of alternating layers of gas hydrate-saturated sands and clay-rich layers with little to no gas hydrate. This interpretation is in general agreement with the marked freshening of the interstitial waters observed in sampled sand layers.

There appears to be little correlation between the core-derived physical property data and defined lithostratigraphic or logging units. Sediment density and porosity determined from moisture and density (MAD) analyses show a typical sediment compaction trend with depth. Slight deviations in the density and porosity profiles with depth reflect small-scale lithologic variability and generally match the downhole measured LWD/MWD data. Infrared (IR) imaging of the recovered cores was routinely carried out on the catwalk to detect and characterize the nature of gas hydrates in the cores. A large number of IR-imaged cold spots were detected in the cores from Holes U1325B and U1325C and were partially subsampled for focused interstitial water analyses and microbiology studies. In many cases the IR-imaged cold temperature anomalies correlated with layers of high resistivity and low interstitial water salinities and chloride concentrations, which have been shown to be associated with the occurrence of gas hydrate.

Six deployments of temperature tools were attempted at Site U1325, and all six provided usable data. Three APC3 and one DVTP deployment in Hole U1325B as well as two additional DVTP deployments in Hole U1325C occurred during times of relatively low heave and yielded excellent quality data. A linear fit to the points indicates a thermal gradient of $0.060^{\circ} \pm 0.003^{\circ}\text{C}/\text{m}$ and a seafloor intercept of $3.03^{\circ} \pm 0.55^{\circ}\text{C}$. The temperature data indicate that the base of the methane hydrate stability zone for this site, assuming a hydrostatic pore pressure gradient and seawater salinity, is at 275 ± 25 mbsf.

Pressure coring tools were deployed seven times at Site U1325, including two PCS cores, two HRC cores, and one FPC core in Hole U1325B. Hole U1325C included one FPC core from above the projected depth of the BSR and a PCS core from well below the estimated depth of the BSR. Pressure coring proved to be extremely difficult at Site U1325; only the deepest PCS deployment (Core 311-U1325C-10P; 256.5 mbsf) successfully recovered core under pressure, which was investigated by a controlled ship-board degassing experiment. All other attempts to deploy pressure cores failed for various reasons, partially due to difficult lithologic conditions (the presence of unconsolidated fine sand) and the potential effect of adverse ship heave conditions. Degassing of Core 311-U1325C-10P yielded 2.07 L of methane gas and may have contained small amounts of gas hydrates (0.4%) or free gas (0.3%) depending where the base of gas hydrate stability (BGHS) is situated (see uncertainty in temperature-derived BGHS above).

A total of 92 interstitial water samples were processed from two holes cored at Site U1325. The salinity and chlorinity profiles at this site appear to indicate an advective transport system with the occurrence of higher than seawater salinity and chlorinity, which likely comes from a deeper source (~36 and 600 mM, respectively). The chlorinity of this fluid is 7.3% above the modern seawater value, more than twice the value of the interglacial maximum ocean value (3.5%). The elevated solute concentration may instead be caused by low-temperature diagenetic reactions in the deeper parts of the basin. The most plausible candidate for such a reaction is alteration of volcanic ash to clay minerals and/or zeolites. In the zone extending from ~70 to 240 mbsf (the BSR is at $\sim 230 \pm 5$ mbsf), salinity and chlorinity data show discrete excursions to fresher values indicating that gas hydrate was present in the cores and dissociated prior to processing the samples. Intense microbial activity at Site U1325 results in sulfate depletion, phosphate and alkalinity production, and significant Ca and Mg depletion in the interstitial waters of the upper 3 mbsf. At this site the sulfate/methane interface (SMI) was placed between 4 and 5 mbsf, based on samples recovered from Hole U1325D.

Organic geochemical studies at Site U1325 included analysis of the composition of volatile hydrocarbons (C_1 – C_5) and nonhydrocarbon gases (i.e., O_2 and N_2) from head-space gas samples, void gas samples, and gas samples recovered during the only PCS degassing experiment. With the exception of several samples that contained a large percentage of air, the void gas was almost entirely methane and a small percentage of carbon dioxide (~0.1%–0.5%). Trace quantities of C_{2+} hydrocarbons (<5 ppmv) were present above the seismically inferred BSR (230 mbsf). With greater depth, the concentrations of ethane, propane, and isobutane increased but did not exceed 15 ppmv. The molecular ratios of methane to ethane (C_1/C_2) were the highest observed during Expedition 311. Ratio values from where gas voids were first observed (9.1 mbsf) to the seismically inferred BSR ranged from ~61,000 to ~173,000 and showed no apparent trend with depth.

Microbiological subsampling was routinely conducted on cores recovered from Holes U1325B and U1325C. On each core run, perfluorocarbon tracer (PFT) was continuously metered into the drilling fluid and fluorescent microspheres were deployed to investigate potential drilling fluid contamination of the core. These analyses confirmed that the center of each whole-round sample remains undisturbed for microbiological subsampling. Additional IR images were taken on the cut ends of each microbiological core section to document the thermal warming process of the core before subsampling.

At Site U1325 the pore water chemistry is distinctly different than what was found at other sites during this expedition, in that chlorinity and salinity increase, rather than decrease, with depth. This trend, attributed to a shallow, more saline sourced fluid, is used as evidence for an advection-driven system; it also shows that this part of the basin is likely detached from the more broad-scale fluid-flow regime seen at the other sites.

Seismic data (Fig. F7) suggest that the sediments cored at Site U1325 are similar to those seen in a seafloor-parallel package of sediments on the southwest side of the buried accreted ridge, but they have been deformed and tilted with small offset faults associated with the uplift of the underlying ridge. Enhanced advective fluid flow from the buried accreted basin resulted in gas hydrate formation within the overlying deformed section, preferentially in the coarser-grained turbidite sequences. In contrast, the southwest part of the basin has relatively little deformation and postulated reduced fluid flow and therefore less gas hydrate, which also explains the lack of a strong BSR in that part of the basin.

Although the BSR is not very prominent near Site U1325, the stratigraphic section beneath the projected BSR generally shows high-amplitude reflectors and updip amplitude truncations at the BSR, typical of free gas trapped beneath the BSR. Data from downhole logging and coring also revealed lateral homogeneity between the holes at this site, which is a testament to the orientation of the holes along structural strike in a northwest-southeast direction, avoiding the steep southwest-northeast-dipping slope seen in the seismic data.

Site U1326

Site U1326 (prospectus Site CAS-03C) is located on an uplifted ridge of accreted sediments at the southwest end of the multisite transect established during Expedition 311 (Fig. F5). Recently acquired bathymetry data reveal a collapse structure near the originally proposed primary Site CAS-03B, which may have complicated the recent geologic history of this site. We decided to switch the former alternate location CAS-03C to the primary site to avoid coring directly into the slump feature.

The headwall of the slump feature near prospectus Site CAS-03B is ~250 m high and the slump has eroded a ~2.5 km long section into the ridge. Further examination of the bathymetry data from this area also reveals the occurrence of continuous linear features crossing the ridge in an almost east-west direction. These linear features are

clearly associated with faults as seen on Figure **F10**. The outcrop of these faults at the seafloor generate a surface displacement of up to 25 m and can be seismically traced from the surface down through the sedimentary section to depths below the BSR. The occurrence of these deep throughgoing faults is generally limited to the area of the slump scar. The seismic character of the ridge also changes from the southwest to the northeast across the ridge, with the southwest-facing part of the ridge characterized by strong semicontinuous reflectivity (Figs. **F11**, **F12**), while the seismic reflectivity disappears underneath the northeast-facing flank of the ridge. The differences in the seismic character across the ridge cannot be easily explained, and it is one of the research focuses of the site. The BSR is present underneath most of the ridge (Fig. **F10**), but it is almost absent underneath the slump feature in association with the most heavily faulted sediments.

The objectives of coring and logging at this site are tied to completing the transect of scientific research holes across the northern Cascadia margin. Site U1326 is the closest location to the deformation front, and it probably represents the tectonically youngest occurrence of gas hydrate on the northern Cascadia margin; as such, our primary research objectives include

1. Studying the distribution of gas hydrates,
2. Defining the nature of the BSR,
3. Developing baseline geochemical and microbiological profiles, and
4. Obtaining data needed to ground-truth remotely acquired imaging techniques such as seismic or CSEM surveys.

The operational plan to achieve these objectives was based on a general three-hole concept, which included an LWD/MWD logging hole, a continuous core hole to characterize geochemical and microbiological baselines and proxies for the occurrence gas hydrate, and an additional “tool” hole for specialized pressure coring and selected spot-coring using conventional systems. The special tools hole was also intended to be logged with the conventional IODP wireline logging tool strings.

After transiting from Site U1325, the LWD/MWD bottom-hole assembly (BHA) was tripped to the seafloor. Hole U1326A was spudded at 1445 h on 22 September 2005 at an estimated seafloor tag depth of 1828.1 mbsl (1839.0 mbrf). To avoid blowing out the top of the hole, we adopted a controlled-spud program that featured reduced drilling fluid circulation rates of only 100 gpm and a bit rotation rate of 10–15 rpm. As the hole was advanced, both the circulation rate and bit rotation rate were increased

to maintain a 50 m/h rate of penetration. The LWD/MWD safety protocol was followed without incident and required no corrective action. The hole was drilled to a TD of 300 mbsf; the drill string was pulled clear of the seafloor at 0600 h on 23 September. Hole U1326B was spudded 15 m northeast of Hole U1326A at 1205 h on 23 October, but the first core failed to establish the depth of the mudline and it was decided to start a new hole. Without offsetting from the location of Hole U1326B, Hole U1326C was spudded at 1245 h on 23 October. The first core established a seafloor depth of 1828.0 mbsl (1839.6 mbrf). On the fourth APC core (~30 mbsf), we unexpectedly hit APC refusal and switched to XCB coring. Hole U1326C was advanced by XCB coring to 82.7 mbsf (Table T1), which was followed by three consecutive pressure core deployments within a high-electrical resistivity zone identified on LWD/MWD downhole logs. In this case, the FPC and HRC pressure core runs were added to the traditional continuous core hole to increase the total number of pressure core runs at this site. The FPC was the first pressure core system deployed at this site; it recovered a partial core (15 cm) at less than full pressure. The second system to be deployed was the PCS, which recovered a partial core under pressure. The third pressure core system deployed was the HRC, which was “packed off” with sand, and the cutting shoe and the lower part of the autoclave was unscrewed and left behind in the hole, resulting in the termination of Hole U1326C at a TD of 86.7 mbsf. After tripping the BHA back to the seafloor, the ship was moved ~30 m southwest (15 m from Hole U1326A). Hole U1326D was spudded at 1130 h on 24 October and drilled to 78.8 mbsf in preparation for continued coring to a target depth of 300 mbsf. Because of problems associated with the heave state and schedule limitations, all pressure coring operations were suspended for the remainder of the hole. XCB coring deepened the hole to 271.4 mbsf. With the forecast of deteriorating weather conditions on the morning of 26 October, it was decided to stop coring and complete the hole by drilling to 300 mbsf. Despite the marginal conditions, we did attempt three deployments with the DVTP (252.2, 271.4, and 300 mbsf), which yielded degraded data. After completing coring and drilling operations, we then decided to conduct a single downhole logging run with a nonstandard IODP tool string, which included the SGT, DIT, and the DSI. At 2315 h on 26 October, the logging tool string was lowered to 298.4 mbsf; two successful logging passes were made, with the tool string back on deck at 0345 h. The drill string was pulled clear of the seafloor at 0530 h on 27 October, ending operations in Hole U1326D.

The 271.40 m thick Quaternary sedimentary section cored at Site U1326 was divided into three lithostratigraphic units (Fig. F13). Lithostratigraphic Unit I (0–1.50 mbsf in Hole U1326B and 0–24.13 mbsf in Hole U1326C) is characterized by fine-grained

(clay to silty clay) detrital sediments with thin silty/sandy interlayered turbidite sequences. Foraminifers, shell fragments, sponge spicule remains, mottling, and bioturbation, together with the high marine versus nonmarine ratio of diatoms, indicate hemipelagic sedimentation mixed with turbiditic inputs (coarse-grained facies). Authigenic carbonate cement is also present in lithostratigraphic Unit I.

Lithostratigraphic Unit II (24.13–82.70 mbsf in Hole U1326C and 88.40–146.30 mbsf in Hole U1326D) is also dominated by fine-grained (clay to silty clay) detrital sediments with intervals of turbidite deposited silty/sandy interlayers, but diatoms are absent and other biogenic components are rare. The frequent occurrence of turbidites might indicate times of active tectonism. Below 30.4 mbsf, abundant soft-sediment deformation and dipping strata show that tectonism is more active at the westernmost end of the transect drilled during this expedition. Authigenic carbonate cement and a carbonate concretion also are present in lithostratigraphic Unit II. Soupy and mousseliike textures indicative of gas hydrate occurrence were observed throughout this unit. The lithostratigraphic Unit II/III boundary is marked by the appearance of diatoms, although their abundance is low.

Lithostratigraphic Unit III (146.30–271.40 mbsf in Hole U1326C) is characterized by fine-grained (clay to silty clay) detrital sediments with few thin silty/sandy turbidite interlayers. We interpret this interval as mixed hemipelagic-turbiditic deposition. Soupy sediment textures related to the presence of gas hydrate are also present in lithostratigraphic Unit III.

Precoring LWD/MWD logging was conducted at Site U1326 to direct special tool deployments, such as the PCS, HRC, and FPC pressure cores. The stratigraphic section logged at Site U1326 was divided into three “logging units” based on obvious changes in LWD and wireline gamma ray, density, electrical resistivity, and acoustic measurements. The three identified logging units do not correspond to the defined lithostratigraphic units.

Logging Unit 1 (0–72 mbsf in Hole U1326A) is characterized by a resistivity trend that steadily increases from $\sim 1 \Omega\text{m}$ near the seafloor to $\sim 1.5 \Omega\text{m}$ at the bottom of the unit. This increase in resistivity with depth is matched by an increase in density. This unit shows only a few small resistivity “spikes” (e.g., at ~ 52 and 60 mbsf in the LWD logs) that may be attributed to gas hydrates.

Logging Unit 2 (72–240 mbsf in Hole U1326A) is characterized by uniform density with depth and includes a prominent high-resistivity section from 72 to 107 mbsf in

Hole U1326A, with peaks $>40 \Omega\text{m}$. This interval of high resistivity also includes a number of alternating thin intervals of very low resistivities. These high and low resistivities likely correspond to intercalated layers that have high and low gas hydrate concentrations, respectively. The *P*-wave velocities measured by wireline logging in Hole U1326D show highly variable values (between 1750 and >3000 m/s) in the same high-resistivity interval. The high values of *P*-wave velocity are consistent with high concentrations of gas hydrate.

The top of logging Unit 3 (240–300 mbsf in Hole U1326A) is marked by an increase in density, which remains constant to 300 mbsf. The background resistivity also increases slightly and displays a clear peak, reaching 5 m in the interval 255–261 mbsf. This resistivity peak could be due to gas hydrate or free gas. Free gas seems more likely, as this interval is located below the BSR depth at this site, estimated from seismic reflection data at 234 mbsf. On the other hand, *P*-wave velocities are at least 1700 m/s in this unit and there is no sign of the low velocities as expected in a free gas zone.

LWD/MWD-derived RAB images from Hole U1326A also indicate the presence of gas hydrate in the high-resistivity interval observed in logging Unit 2, with gas hydrate-rich layers alternating with low-resistivity layers that likely contain little or no gas hydrate (Fig. F13). When the RAB images are examined in greater detail, they also show apparent stratigraphic layers in the interval from 50 to 300 mbsf, uniformly dipping to the north-northeast with dips between 45° and 85° . This is in agreement with the west-northwest–east-southeast strike of the uplifted ridge penetrated by Hole U1326A.

The Archie interpretation of water saturations from electrical resistivity logging data yielded evidence for relatively shallow gas hydrate occurrences at this site within the depth interval from ~ 50 to 100 mbsf, with inferred water saturations as low as 20% (Fig. F13). This interpretation is in general agreement with the freshening of the pore waters observed in sand layers at Site U1326. The IR core images taken on the catwalk reveal cold anomalies due to gas hydrate dissociation, which give an independent map of gas hydrate occurrence and concentrations. In some cases IR-imaged cold temperature anomalies correlate with layers of high resistivity. In other cases, however, cold intervals in the IR images are harder to correlate to high-resistivity peaks in the wireline or LWD/MWD logs because of incomplete core recovery and intrasite geologic variability.

Cold temperature anomalies were observed at a wide range of depths from 70 to 250 mbsf, and catwalk sampling was conducted based on these scans. Many interstitial water samples were taken based on IR data to extend the chlorinity anomaly database available for calibrating IR data as a proxy for gas hydrate saturation. We note that the maximum depth of observed IR anomalies is deeper than the anticipated BSR depth of 230 mbsf. It was also suggested that the chlorinity anomaly associated with gas hydrate may actually extend to a depth as great as 270 mbsf at this site. This apparent conflict between the depth of the BSR and the observed occurrence of gas hydrate can be explored with closer examination of gas hydrate stability condition at this site. Because of difficult weather conditions throughout the duration of the coring operations at this site, we deployed only four temperature tools. For the most part, all of the data from the temperature tool deployments were degraded to some extent. At this time we have extrapolated the temperature data from Site U1325 to estimate temperature conditions at Site U1326. With a seafloor temperature of $\sim 3.03^{\circ}\text{C}$ and a geothermal gradient of $-60^{\circ}\text{C}/\text{km}$, the depth to the base of the methane hydrate stability zone was estimated at Site U1326 to ~ 250 mbsf. A predicted methane hydrate stability zone depth of 250 mbsf would help explain most of the deeper IR and chlorinity inferred gas hydrate occurrences, but the several isolated anomalies below this depth require further postcruise consideration.

At Site U1326, only one PCS core (Core 311-U1326C-12P) was recovered successfully under pressure and investigated by controlled shipboard degassing experiments. This core was taken at 84.2 mbsf. The degassing of this single PCS core yielded 19.1 L of gas, which was determined to be equivalent to a pore space gas hydrate saturation of 40%, which is in close agreement with the gas hydrate saturations estimated from the LWD/MWD Archie resistivity calculations in the same interval.

This site was selected to further characterize the occurrence of what could be the tectonically youngest gas hydrate occurrence on the Northern Cascadia margin, considering its location within the westernmost portion of the deformation front. Hydrocarbon headspace gas measurements from Holes U1326C and U1326D show that methane is the dominant hydrocarbon gas within the cored interval at this site. Although the C_{2+} hydrocarbon void gas concentrations were <125 ppmv for all samples, their relative abundances and distribution were valuable for describing the gas hydrate system at Site U1326. Low C_1/C_2 ratios within the interval from ~ 35 to ~ 72 mbsf were determined to be directly associated with two recovered gas hydrate samples. With greater depth, ethane concentrations returned to the near-surface concentrations close to the depth of the seismically inferred BSR. Isobutane concentrations

were also elevated within the same interval, which may indicate the presence of a Structure II gas hydrate. The occurrence of more complex hydrocarbon gases generally indicates contribution of a deeper gas source. The salinity and chlorinity profiles at this site also point to a deeper fluid source with a chloride concentration higher than seawater, indicative of low-temperature diagenetic reactions in the deeper parts of the site. Except for the difficulties experienced with pressure coring and special temperature tool deployments, all of the objectives set for Site U1326 were considered fulfilled.

Site U1327

Site U1327 (prospectus Site CAS-01B) is located near Sites 889/890, approximately at the midslope of the accretionary prism over a clearly defined BSR estimated to be at 223 mbsf (Fig. F14, F15). The primary research objectives for this site are linked to the transect concept of this expedition. Sites 889/890 provided critical baseline data for the development of the objectives of this research expedition. Historical work with data from Sites 889/890 has shown that the concentration of gas hydrate in sediments has been estimated by using (1) deviations in interstitial water chlorinities from measured baseline conditions, (2) electrical resistivity measurements by applying Archie's relation, and (3) both seismic- and downhole logging-derived *P*- and *S*-wave velocities. But results from Leg 146 left many questions unanswered (Riedel et al., 2005), such as:

- What is the geochemical reference profile for chlorinity and other geochemical gas hydrate proxies?
- What is the electrical resistivity baseline for the sediments drilled at this site and what are the appropriate empirical Archie coefficients?
- What are the reference profiles for acoustic *P*- and *S*-wave velocities?

Answers to these questions are needed to calibrate remote sensing techniques, such as reflection seismic and CSEM surveys, which can be obtained through research coring and downhole logging.

Five holes were occupied at Site U1327 (Table T1). Hole U1327A was dedicated to LWD/MWD measurements to a TD of 300 mbsf. Initially we planned to drill to 350 mbsf, but tight time constraints during LWD/MWD operations made it necessary to reduce the depth of the planned penetration at this site. The first APC core in Hole U1327B missed the mudline with a full core, the core was curated, and a new hole was spudded without offsetting the ship. Hole U1327C was then cored (10 APC cores, 22

XCB cores, and 3 PCS cores; recovery = 88.3%) to 300.0 mbsf. The PCS was deployed three times in Hole U1327C, and four advanced piston corer temperature (APCT) tool measurements were made (Cores 311-U1327C-3H, 5H, 7H, and 9H). Three additional deployments using the DVTP and Davis-Villinger Temperature-Pressure Probe (DVTPP) were attempted, but data quality is marginal due to strong heave conditions and instrument failures. In Hole U1327D, which was drilled and cored as a special tools hole, two APC cores were also taken from the surface to 16.4 mbsf for a high-resolution microbiological and geochemical study of the SMI. In this hole, the PCS was deployed three times, together with four deployments of the HRC and two deployments of the FPC system. The last pressure core was taken at 246.5 mbsf, and the hole was advanced to a TD of 300 mbsf for the logging program.

The first tool string deployed was the triple combination (triple combo), which reached 295.4 mbsf, and the hole was logged to near the top of the logging run without incident. A combination of a ship heave event (>3 m) and the oversized borehole apparently caught and tore off the density tool caliper arm. The damaged tool string was returned to the ship without further incident. The VSP logging program in Hole U1327D was started the following morning with the 1 h long preshooting marine-mammal observation period and a 30 min pressure ramp-up time for the generator-injector gun. The VSP program was successfully carried out with 16 clamping stations covering the depth range from 181 to 276 mbsf before the arm on the Well Seismic Tool (WST) apparently broke. The WST was returned with some difficulties to the ship. Because of the critical nature of the acoustic wireline logging program to the expedition, it was decided to drill a new hole and acquire additional pressure cores and a wireline acoustic logging survey.

Hole U1327E was advanced to 3 mbsf, and a single 9.5 m long APC core was taken for high-resolution microbiological and geochemical sampling of the SMI, which was partially missed during an earlier attempt. Two additional PCS cores and one HRC core were taken, out of which only the second PCS at ~80.5 mbsf recovered sediment under pressure. Excessive heave >3.5 m forced the termination of pressure coring operations, and the hole was advanced by drilling to a completion depth of 300 mbsf to prepare for the second wireline logging run. Because of the excessive heave conditions it was decided to run a nonstandard tool string composed of the HNGS-DIT-DSI, which allowed measurement of natural gamma, resistivity, and acoustic data (V_p and V_s). Two successful runs of this tool string were completed, recovering high-quality downhole logging data.

Site U1327 is located near two prominent topographic highs, composed almost entirely of accreted sediments that rise >200 m above the surrounding seafloor (Fig. F14). The ridges of accreted sediment lack any coherent seismic reflections and are associated with underlying thrust faults that contributed to the overall uplift of the area between the two ridges (Westbrook, Carson, Musgrave, et al., 1994; Riedel, 2001). Site U1327 is situated near the northwest flank of the easternmost ridge and is characterized by a ~90 m thick cover of slope-basin sediments, underlain by a thick section of accreted sediments (Fig. F15). The stratigraphy at Site U1327 was divided into three lithostratigraphic units (Fig. F16).

Lithostratigraphic Unit I (0–90.1 mbsf in Hole U1327C; age = Holocene–Pleistocene; <0.3 Ma) is composed of dark greenish gray and dark gray clay and silty clay, often interbedded with silt, clayey silt, sandy silt, sand, and gravel layers. Lithostratigraphic Unit I is characterized by fine-grained detrital sediments (clay and silty clay) with abundant coarse-grained layers up to 6 cm thick indicating turbiditic deposits and a sedimentation rate of 22 cm/k.y. Authigenic carbonate cement is present in Unit I, but no evidence for dolomite precipitation is found, which is inferred from interstitial water analyses. A unique characteristic of Unit I is the anomalous occurrence of lithified carbonate nodules and rocks of various lithologies, such as granite, which probably are dropstones from rafting icebergs. The lithostratigraphic boundary between Units I and II is marked by a sharp decrease in sand and silt layers and the onset of diatom-rich sediments. It is also the seismically inferred boundary between slope-basin type and accreted sediments (Fig. F15).

Lithostratigraphic Unit II (90.1–170.4 mbsf in Hole U1327C; age = Pleistocene; >0.3–1.0 Ma) is composed of dark greenish gray and dark gray clay, clay with diatoms, and silty clay, silty clay with diatoms, and diatom silty clay locally interbedded with sandy silt and sand layers and lenses. Very few carbonate cements are present in Unit II, but lithified carbonate nodules and exotic rocks (some up to 8 cm in diameter) as described in Unit I are very abundant. Within Unit II soupy and mussel-like sediment textures related to the presence of observed gas hydrate can be correlated to IR-imaged cold spot anomalies detected on the catwalk. The sedimentation rate of Unit II is estimated at ~16 cm/k.y., and the intervals with a high ratio of nonmarine versus marine diatoms indicate the increasing contribution of terrigenous detrital sediments from land sources via turbidites. The great abundance of marine diatoms along with resting spores within lithostratigraphic Unit II suggest blooming in shallow-water shelf environment and coastal upwelling then reworking by turbidity currents.

Lithostratigraphic Unit III (170.4–TD in Hole U1327C; age = Pleistocene; >1 Ma) is composed of dark greenish gray and dark gray silty clay in the upper part of the unit, whereas the lower part (below 248 mbsf) is dark greenish gray silty clay with diatoms. Lithostratigraphic Unit III is distinguished from lithostratigraphic Unit II by the sudden absence of diatoms, as well as the degree of induration of the sediments. Diatoms reappear at the bottom of Unit III. Few carbonate cements are present and lithified carbonate nodules and exotic rock clasts were found, especially within Core 311-U1327C-24P taken at 197.3 mbsf. The depositional environment for Unit III was interpreted as an abyssal plain with sediment transport and deposition dominated by low-energy turbidity currents. The sedimentation rate in this section, as determined from biostratigraphic analyses of diatoms, is low for this site at ~12 cm/k.y., which is consistent with a basin-plain setting.

The downhole logged section at Site U1327 was divided into three “logging units” based on obvious changes in the LWD/MWD and wireline gamma ray, density, electrical resistivity, and acoustic measurements. There is no apparent correlation between the logging units and lithostratigraphic units at this site.

Logging Unit 1 (0–120 mbsf) is characterized by a resistivity trend that steadily increases from ~1 Ωm near the seafloor to ~2 Ωm at the bottom of the unit. *P*-wave velocities in this unit average ~1550 m/s.

Logging Unit 2 (120–240 mbsf) is characterized by a constant background resistivity value of ~2 Ωm and relatively high *P*-wave velocities that average ~1750 m/s over most of the interval. This unit shows a number of thin (0.5–1.0 m thick) high-resistivity sections that may be attributed to gas hydrates, and the occurrence of relatively high downhole measured *P*-wave velocities in these same intervals also indicates the presence of gas hydrate. A thick section of continuous high electrical resistivities and *P*-wave velocities combined with low density measurements was observed in the LWD/MWD data at 120–138 mbsf in Hole U1327A, which was the target of several pressure coring deployments (see below for details). However, it was later determined that this high-resistivity zone was laterally discontinuous and was penetrated in the adjacent Holes U1327C and U1327D at a much greater depth (ranging from 20 to 35 m deeper) and was not intersected in Hole U1327E.

The top of logging Unit 3 (240–300 mbsf) is defined by a sharp decrease in *P*-wave velocity. Below 240 mbsf, *P*-wave velocities drop to very low values near fluid velocities (~1500 m/s), suggesting the presence of small amounts of free gas. Logging Unit 3 also

displays a small drop in resistivity compared to logging Unit 2. Although resistivity tends to be just above $2 \Omega\text{m}$ in Unit 2, it is just below $2 \Omega\text{m}$ in Unit 3.

Variations in the core-derived physical properties apparently do not correlate well with the identified lithostratigraphic and logging units, with the exception of the boundary between lithostratigraphic Units I and II, marked by a sharp increase in magnetic susceptibility. Sediment density and porosity determined from MAD analyses show a typical trend with depth, suggesting normal compaction. Slight deviations from this trend reflect small-scale lithologic variability and generally match the porosity and density trend determined from the LWD/MWD data.

Infrared imaging of the recovered cores was routinely carried out on the catwalk to detect and characterize the nature of gas hydrates in the cores. A large number of IR-identified cold spots were detected in the cores from Holes U1327C and U1327D; however, an apparent depth mismatch in the occurrence of major cold core sections between the holes indicated significant intrasite geologic variability. Apparent mismatches between the IR-inferred gas hydrate occurrences in Hole U1327C and the LWD/MWD resistivity-inferred gas hydrate occurrences in Hole U1327A further document the lack of lateral continuity at this site. In situ temperature measurements with the APCT (four measurements), DVTP (one deployment, instrument failure), and DVTTP3 (two deployments, poor quality due to heave and instrument calibration problems) were conducted. The data suggest a geothermal gradient of $\sim 61^\circ\text{C}/\text{km}$, which is considerably higher than what was estimated for Site 889 ($\sim 54^\circ\text{C}/\text{km}$).

The PCS was deployed eight times at Site U1327 (three times in Hole U1327C, three times in Hole U1327D, and twice in Hole U1327E), five of which recovered sediment under pressure. In addition to the PCS deployment, the HRC was used four times (with three cores recovered under pressure) and the FPC was used twice, but only one of the FPC cores was recovered under pressure. Degassing of the five PCS cores from this site that were recovered under pressure showed variable gas concentrations with depth. The deepest PCS core was taken at 246.5 mbsf, which is ~ 25 m deeper than the seismically inferred BSR depth. This pressure core (Core 311-U1327D-17P) had a pore space methane concentration of 208 mM, which is equivalent to a free gas concentration in the pore space of the sediment of 0.9%. The other four PCS cores were taken from within the predicted depth interval of the methane hydrate stability zone. Core 311-U1327E-3P is the shallowest PCS core taken at this site, from ~ 80.5 mbsf. It yielded 1.31 L of methane (81 mM) or, equivalently, 0.2% gas hydrate concentration. The remaining three PCS cores were taken at 122.3, 155.6, and 197.8 mbsf (Core 311-

U1327C-15P yielded 1.1 L [94 mM] of methane; Core 311-U1327D-10P yielded 9.4 L [673 mM] of methane; and Core 311-U1327C-24P yielded 4.0 L [251 mM] of methane). Of these three cores, only Core 311-U1327D-10P yielded enough gas to infer the occurrence of a significant amount of gas hydrate, with an estimated gas hydrate pore space concentration of ~8%. The other two PCS cores yielded gas hydrate pore space concentrations ranging from 0.2% to 1.8%.

Three HRC cores (Cores 311-U1327D-4E at 125.3 mbsf, 311-U1327D-12E at 170.5 mbsf, and 311-U1327D-14E at 217.7 mbsf) and one FPC core (Core 311-U1327D-13Y at 203.6 mbsf) were successfully recovered and transferred under pressure to storage chambers for shore-based analyses. All of these cores were imaged with X-ray and logged for *P*-wave velocity and density within their storage vessels. Some of the recovered pressure cores exhibited evidence of gas hydrate, including high *P*-wave velocities and anomalous low density readings.

High-resolution interstitial water sampling was carried out to characterize the SMI, which was determined to be located between 9 and 10 mbsf. An overall smooth decreasing trend in pore water chlorinity and salinity is observed, similar to results from Sites 889/890 (Westbrook, Carson, Musgrave, et al., 1994). At ~120 mbsf the chlorinity and salinity profiles exhibit distinct anomalies that likely indicate freshening as a result of the dissociation of gas hydrate during the core recovery process (Fig. F16). Chloride values as low as 70 mM (equivalent salinity = 3.7) were measured in two samples, correlated to sand layers (up to 3 cm thick) that also showed clear IR cold spot anomalies. Such a low pore water chlorinity is equivalent to a pore space gas hydrate concentration of ~80%. No chlorinity anomalies occur below ~225 mbsf, which is close to the seismically inferred depth of the BSR. However, the strongly decreasing chlorinity profile from the surface to below the BSR at Site U1327 suggests mixing of standard marine pore water with a deep-sourced, relatively fresh fluid. Interestingly, interstitial water chlorinity and salinity remain almost constant beneath the BSR.

Organic geochemical studies at Site U1327 included analysis of the composition of volatile hydrocarbons (C_1 – C_5) and nonhydrocarbon (i.e., O_2 and N_2) gases from headspace gas samples, void gas samples, and gas samples recovered during PCS degassing experiments. The predominant hydrocarbon gas found in the cores from Site U1327 was methane; however, we did see an increase in ethane concentrations in the void and headspace gases collected from the stratigraphic section overlying the projected depth of the BSR. This increase in ethane concentrations near the BSR can also be seen in the plot of the C_1/C_2 void gas ratios (Fig. F16), which decrease with depth toward

the BSR. Other studies have shown that void gas ratios can decrease in a gas hydrate with the preferential selection of ethane into the gas hydrate structure. But the observed increase in the C_1/C_2 ratio below the projected depth of the BSR may indicate a methane-enriched free gas accumulation. In general, the C_1/C_2 ratios were generally high (>1000), indicating a predominant microbial origin for the observed methane.

Microbiological subsampling was routinely conducted on cores recovered from Hole U1327C. On each core run PFT was continuously metered into the drilling fluid, and fluorescent microspheres were deployed on selected cores to investigate potential drilling fluid contamination of the core. These analyses confirmed that the center of each whole-round sample remains undisturbed for microbiological subsampling. Additional IR images were taken on the cut ends of each microbiological core section to document the thermal warming process of the core before subsampling.

The primary research objectives of this site are linked to the results from Sites 889/890 of Leg 146 to delineate critical geochemical, geophysical, and microbiological data for gas hydrate proxy analyses. Site U1327, along with Sites 889/890, is the center of a research study area that since 1992 has been the focus of many interdisciplinary gas hydrate studies.

Geochemical analyses of recovered gas samples and interstitial fluids documented a relatively complex fluid regime for this site, similar to what was observed at Sites 889/890 and U1329. The rapid chlorinity concentration decrease with depth strongly suggests fluid communication with a deep-seated, fresher fluid source. Earlier interpretations of the chlorinity profile measured at Sites 889/890 attributed the pore water freshening entirely to gas hydrate dissociation upon core recovery (Hyndman et al., 1999). However, the new interstitial water data collected at Site U1327, in combination with other gas hydrate proxy data (such as IR images and downhole logging measurements), strongly support the development of a new undisturbed baseline for interstitial water chlorinity along the Cascadia margin that progressively decreases with depth, similar to what was suggested by Ussler and Paull (2001). In other words, gas hydrate-related pore water freshening at Site U1327 is interpreted to be limited to only the interval from ~120 to ~220 mbsf, which is near the depth of the seismically identified BSR. Geochemical analyses of hydrocarbon gases from this site reveal a similar story as the analysis of the interstitial fluids with the gas chemistry in the inferred gas hydrate-bearing section dominated by gas hydrate formation processes with an apparent increase in the more stable gas hydrate natural gas components.

Identifying gas hydrate and constraining its relative abundance in combination with providing physical models for the BSR are critical geophysical objectives at this site. A suite of geophysical experiments was conducted including precoring LWD/MWD logging, conventional wireline logging, and VSP to characterize the physical properties of gas hydrate-bearing sediments. The presence of gas hydrate is generally characterized by increases in measured electrical resistivity and acoustic velocity. However, the LWD/MWD data from Hole U1327A showed an 18 m thick zone of increased resistivity and acoustic velocity values combined with a decrease in density, which was interpreted to contain large amount of gas hydrate, with concentrations >50% of the pore volume. However, the same interval was penetrated in the adjacent core holes, Hole U1327C and Hole U1327D, at much greater depths and with lower estimated gas hydrate concentrations inferred from the wireline electrical and acoustic logs. In the repeat wireline logging in Hole U1327E, no evidence of this interval was found. This shows large intrasite variability in gas hydrate content that is probably controlled by lithostratigraphic changes or structural complexities. This observation has dramatic implications on the calibration of geophysical surveying techniques such as seismic methods or CSEM surveys that are commonly used to detect and quantify gas hydrates.

The combined use of IR imaging, interstitial pore water analyses, and void/headspace gas sampling revealed that many IR cold spots identified on the catwalk occur in relatively thin but coarser-grained intervals of sand and/or silty sand (interpreted to represent turbidite deposits) with gas hydrate concentrations in several cases exceeding 80% of the pore space. The gas hydrate forms preferentially in these coarser-grained sections. Therefore, it is likely that the 18 m thick LWD/MWD-measured high-resistivity and acoustic velocity interval (120–138 mbsf) in Hole U1327A is a localized feature made up of discontinuous thinly bedded gas hydrate-saturated turbidite sand lenses, that when surveyed with a logging tool with relatively lower vertical resolution the gas hydrate-bearing section appears more massive in nature as a single unit on the downhole log.

The results of the VSP seismic experiment taken in combination with the downhole acoustic logging from Hole U1327E also shows a relatively thick zone below the depth of the BSR that probably contains a considerable amount of free gas. The acoustic velocity decrease at the BSR is not abrupt; it is more transitional across a ~25 m thick interval from >1800 m/s in the gas hydrate stability zone to as low as 1280 m/s in the underlying section. This confirms earlier interpretations of the frequency-dependent reflection behavior of the BSR from vertical incidence seismic data. The BSR was

thought to be a gradient zone in which velocity decreases over several meters (Chapman et al., 2002).

This site was intended to characterize the gas hydrate occurrence on the northern Cascadia margin to establish an understanding of the geologic, geophysical, and geochemical controls on the occurrence and abundance of gas hydrates along the five-site transect established during this expedition. All of the objectives set for Site U1327 are considered fulfilled.

Site U1328

Site U1328 (prospectus Site CAS-06A) is located within a seafloor cold vent field, 2 km × 4 km, consisting of at least four vents associated with near-surface faults. The cold vents are characterized by near-vertical seismic blank (or wipeout) zones that are between 80 and several 100 m wide and show a clear east-west trend as identified from 3-D seismic imaging. The most prominent vent in the field, referred to as Bulls-eye vent, is the target of this site and has been the subject of intensive geophysical and geochemical studies since 1999 (e.g., Riedel et al., in press). Site U1328 is different from all of the other sites visited during this expedition in that it represents an area of active, focused fluid flow. A number of seismic surveys cross this site, including two single-channel seismic (SCS) 3-D surveys and a regional MCS data set. In Figure F17 we show MCS Line XL07 and the main lithostratigraphic units, hole locations, and penetration depths at this site.

Several different mechanisms have been proposed for the origin of the seismic blanking observed in the Bulls-eye vent area and for the implications concerning the nature of the fluid venting (Riedel et al., 2002, in press; Zühlendorff and Spiess, 2004; Wood et al., 2002). The objectives of coring and logging at this location are to test the different models for the cold vent structure and associated causes of seismic blanking, the rate of methane advection, and potential loss of methane into the water column. It was important to obtain a high-resolution temperature profile for this area to assess any evidence of active fluid flow. Pressure coring using the PCS system will help assess the occurrence of free gas within the gas hydrate stability zone below Bulls-eye vent and possibly ascertain the cause of the observed seismic blanking. Additional pressure coring using the HRC and FPC systems will be used to recover high-quality gas hydrate-bearing core for shore-based analyses.

Five holes were occupied at Site U1328. Hole U1328A was dedicated to LWD/MWD measurements to a TD of 300 mbsf. Hole U1328B was continuously cored (six APC, two XCB, and two PCS cores; recovery = 72.8%) to a depth of only 56.6 mbsf and was terminated after strong winds and severe ship heave conditions required us to pull out of the hole. Two APCT measurements were made (APCT and APC3). As a calibration experiment, the DVTPP was also deployed at the bottom of Hole U1328B. After waiting on weather for 16 h, conditions improved to the point to allow drilling and coring operations to continue. Hole U1328C was drilled from the seafloor to the maximum depth of Hole U1328B (56.5 mbsf). Hole U1328C was then continuously cored (4 APC cores, 22 XCB cores, and 1 PCS core; recovery = 80.7%) to a TD of 300 mbsf. In addition, three APCT and two DVTPP temperature measurements were made. Hole U1328C was wireline logged with the triple combo and FMS-sonic tool strings. The triple combo logged from a total depth of 294 mbsf, and two uphole passes of the FMS-sonic tool string logged the same depth interval. The wireline logging program in Hole U1328C included a VSP survey, with the deepest clamping position at 286 mbsf. The VSP included 35 clamping positions, with the uppermost station at 106 mbsf. Hole U1328D was cored as a special high-resolution combined microbiology and geochemistry research hole with two XCB cores and a single FPC core taken at the bottom of the hole. Hole U1328E was a special tools hole, which also included the deployment of six XCB cores within the upper 46.0 mbsf to recover additional samples of gas hydrate. Seven pressure cores were taken (three PCS, two HRC, and two FPC), separated by XCB cores and drilled intervals. Hole U1328E also included two DVTPP deployments for temperature data. In Figures F18 and F19, data from LWD/MWD logging (RAB images, porosity, and Archie-derived pore water saturations), core-derived physical property data (porosity), interstitial water chlorinity, and C_1/C_2 ratios from void gas samples are shown in comparison to the two key crossing seismic lines from this site.

At Site U1328 a 300 m thick sequence of Quaternary (0–1.6 Ma) slope and slope basin sediments was recovered (Fig. F17). The biostratigraphy determined for Site U1328 was based on the examination of diatoms from Holes U1328B and U1328C. The stratigraphic section cored and logged at Site U1328 was divided into three lithostratigraphic units.

Lithostratigraphic Unit I (0–56.50 mbsf in Hole U1328B, 56.50–132.60 mbsf in Hole U1328C, 0–15.00 mbsf in Hole U1328D, and 0–92.26 mbsf in Hole U1328E) is characterized by fine-grained detrital sediments (dark greenish gray clay and silty clay) with abundant coarse-grained layers up to 6 cm thick indicative of turbidite deposits.

The boundary between lithostratigraphic Units I and II is marked by a sharp decrease of sand and silt layers and by the onset of diatom-bearing sediments.

Lithostratigraphic Unit II (132.60–197.10 mbsf in Hole U1328C and 197.00–198.00 mbsf in Hole U1328E) is characterized by fine-grained (clay to silty clay) detrital sediments with some silty interlayers from turbiditic deposits and siliceous fossils. Abundant marine diatoms along with resting spores within lithostratigraphic Unit II suggests the diatoms bloomed in a shallow-water shelf environment associated with coastal upwelling, and then the diatoms were reworked. The boundary between lithostratigraphic Units II and III is distinguished by the sudden absence of diatoms.

Lithostratigraphic Unit III (197.10–300.00 mbsf in Hole U1328C) is characterized by fine-grained (clay to silty clay) detrital sediments with very few silty interlayers from turbidite deposits, likely deposited in an abyssal plain environment. The presence of authigenic carbonate cements show that diagenetic processes are active in lithostratigraphic Unit III. Because of the large input of turbidite deposits, in lithostratigraphic Unit I we see a significant increase in sedimentation rate from 24.6 cm/k.y. in lithostratigraphic Unit II to 37.5 cm/k.y. in lithostratigraphic Unit I.

At Site U1328, gas hydrate was sampled and evidence of gas hydrate was found in the recovered cores in the form of soupy and mousselike textures. At this site, soupy texture is only present within lithostratigraphic Unit I. At Site U1328 mousselike textures are present occasionally within lithostratigraphic Units I to III.

The LWD/MWD and wireline logged section in Holes U1328A and U1328C was divided into three “logging units” based on obvious changes in the downhole measured gamma ray, density, electrical resistivity, and acoustic measurements; for the most part these logging units correspond closely to lithostratigraphic Units I, II, and III.

Logging Unit 1 (0–128 mbsf in Hole U1328A) is characterized by a well-defined increase in density with depth. The most striking feature in this unit is seen in the LWD/MWD logs between the seafloor and 46 mbsf, where high resistivities ($>25 \Omega\text{m}$) alternate with intervals of much lower resistivity ($1\text{--}2 \Omega\text{m}$). The high resistivities likely indicate the presence of gas hydrate, which were sampled during coring.

Logging Unit 2 (128–200 mbsf in Hole U1328A) is marked at the top by a clear decrease in density log values. Unit 2 is also characterized by a constant relatively low background resistivity value just above $1 \Omega\text{m}$ and only one significant high-resistivity log-inferred gas hydrate occurrence at a depth of 160 mbsf.

Logging Unit 3 (200–300 mbsf in Hole U1328A) is characterized by small increases in background resistivity and notably more variability resistivity and *P*-wave velocity log values.

Physical properties were measured in cores recovered from four of the holes at this site, with the core-derived MAD data comparing favorably to both the LWD/MWD and downhole wireline data. Numerous low-porosity outliers, however, were interpreted to represent sand-rich samples.

It has been shown that RAB images as derived from the LWD/MWD GeoVision tool can be used to image gas hydrate in sediment. High concentrations of gas hydrate are evident in the interval 0–46 mbsf of Hole U1328A, where resistivities are high. The RAB images from Hole U1328A also reveal numerous high-resistivity sinusoidal features in the upper 46 mbsf in Hole U1328A and again at 90–100 mbsf, which can be interpreted as dipping fractures containing gas hydrate or possibly free gas. These near-vertical fractures may act as gas migration conduits that feed the gas hydrate accumulation observed near the seafloor.

To estimate the amount of gas hydrate at Site U1328, the GeoVision-derived resistivity logs were used as input to the Archie relation to estimate pore volume water saturations. Gas hydrate saturation is the percentage of pore space in sediment occupied by gas hydrate, which is the complement of the water saturation. These calculations have shown highly variable gas hydrate saturations in the upper 46 mbsf, with values as high as 95% in several intervals up to 10 m thick. These layers of high gas hydrate concentrations alternate with layers of much higher water saturations.

The combined analysis of acoustic velocities and waveform amplitudes can also help identify the occurrence of gas hydrate and/or free gas. In the interval between 210 and 220 mbsf, near the seismically inferred depth of the BSR at -219 mbsf, the *P*-wave velocity drops slightly while the *S*-wave velocity increases. This could be due to the coexistence of free gas and gas hydrate in the vicinity of the BSR, which has been observed on Blake Ridge and Hydrate Ridge. Additional evidence for the occurrence of free gas below the BSR was noted in the analysis of the LWD/MWD acoustic coherence data, borehole fluid pressure response during drilling, and the wireline *P*-wave, resistivity, density, and neutron logs. The PCS degassing experiments of a single core in Hole U1328E also suggest the presence of a free gas phase below the BSR at this site (see below). However, it is important to note that the VSP survey obtained in Hole

U1328C yielded a very uniform velocity profile with depth and no clear velocity contrast around the expected depth of the BSR.

Although we deployed the APCT, APC3, DVTP, and DVTPP temperature tools a total of nine times, we recorded only three reliable data points, which was mostly the result of severe ship heave conditions. The in situ temperatures determined for Site U1328 are generally similar to those at Site U1327 and slightly higher than those determined at similar depths during Leg 146. For Site U1328 we observed a seafloor temperature of $\sim 3.5^{\circ}\text{C}$ and a geothermal gradient of $\sim 5.4^{\circ}\text{C}/100\text{ m}$, which corresponds to a predicted depth to the base of the methane hydrate stability zone ranging from 222 to 247 mbsf given the uncertainty in the data.

All cores from this site were systematically scanned upon arrival on the catwalk to detect IR anomalies indicative of gas hydrate dissociation during core recovery. Strong cold anomalies were detected in the shallowest cores from this site. A large amount of gas hydrate had been anticipated in the upper 30 mbsf based on LWD/MWD resistivity measurements and previous piston coring at this site (Riedel et al., 2002, in press; Solem et al., 2002), and a considerable amount was found. However, the LWD/MWD high resistivity–inferred gas hydrate occurrence at 90–100 mbsf in Hole U1328A was not seen on either the IR images or the wireline logs, both taken in Hole U1328C. The IR data may have missed this interval due to poor core recovery, but the lack of a high-resistivity layer in the wireline log points to horizontal inhomogeneity. This is not unexpected, in that the high-resistivity layer in the LWD/MWD image contains a steeply dipping fracture that is unlikely to be intersected at the same depth even in two closely spaced holes. The strongest anomalies were detected just above the seismically inferred BSR (210–222 mbsf), and a few small but distinct cold anomalies were observed below it, one of which exhibited pore water freshening indicative of the presence of gas hydrate (Core 311-U1328C-22X; ~ 249 mbsf). The possibility of gas hydrate beneath the BSR may indicate the presence of Structure II gas hydrate, which is stable below the base of the pure methane Structure I gas hydrate stability zone. However, it should be noted that there remains substantial uncertainty in the exact depth of the base of the gas hydrate stability field at this site and the depth of the BSR. All available data suggest that the gas hydrate stability zone could be deeper than the initially inferred depth of the BSR, which may result in shifting the observed IR anomalies to depths well within the gas hydrate stability field.

At Site U1328 we attempted 11 deployments of the 3 different pressure coring tools. The PCS was deployed six times at Site U1328 (twice in Hole U1328B, once in Hole

U1328C, and three times in Hole U1328E), five of which recovered sediment under pressure. In addition to the PCS deployment, the HRC was used two times and the FPC was used three times, but none of these cores were recovered under pressure. Degassing of the five successful PCS cores from this site showed variable gas concentrations with depth. Two PCS cores were taken within the near-surface gas hydrate-bearing section from 0 to 46 mbsf (Core 311-U1328B-4P at 15 mbsf and Core 7P at 26.5 mbsf), which yielded 21.5 and 2.5 L of methane, respectively, equivalent to pore volume gas hydrate saturations of ~15% and ~2%. X-ray images of Core 311-U1328B-4P under pressure show 2–6 mm thick low-density structures that disappeared after degassing and are interpreted as gas hydrate veins. This core showed large amounts of gas expansion and sediment extrusion during depressurization, as did Cores 311-U1328B-7P and 311-U1328E-10P to a lesser extent. The two PCS cores taken in the Holes U1328C and U1328E at 92.0 mbsf (Cores 311-U1328C-5P and 311-U1328E-10P) each yielded quite different amounts of methane gas and would equate to the occurrence of 22% (Core 311-U1328C-10P) and 0.7% of gas hydrate (Core 311-U1328E-5P). One PCS core (Core 311-U1328E-13P) was taken from a depth of 233.0 mbsf, very close to the seismically inferred BSR depth. This core yielded 60 L of methane, the largest recovery of gas during this expedition. This amount of methane would fill 58% of the pore space if it was present as free gas or 50% if it was present as gas hydrate. However, gamma ray density scans during degassing and X-ray scans after depressurization showed almost no evidence of gas expansion cracks or large voids in the core. It is speculated that most of the gas released during degassing of this core came from the space between the inner and outer core barrel, where a gas bubble was observed during X-raying. Volumetric analysis suggests the presence of a 540 mL headspace in the outer core barrel which corresponds will to the 410 mL volume that would have been occupied by the released amount of gas at in situ conditions. The apparent extra gas in the PCS might have swabbed into the core barrel during the coring process.

A total of 91 interstitial water geochemistry samples were processed from four of the holes cored at Site U1328. The composite chlorinity profile for this site shows four distinct zones. In the upper ~30 mbsf, we see a striking increase in chlorinity, with maximum values exceeding 850 mM. The observed excess solutes are interpreted to result from salt exclusion from the water lattice structure during in situ gas hydrate formation that has not been removed by advective or diffusive processes. The second zone from ~30 to ~150 mbsf is characterized by relatively constant chlorinity values ranging from 538 to 570 mM. A third zone, extending from ~150 to 250 mbsf, across the BSR, shows discrete excursions to fresher chlorinity values (as low as 348 mM),

suggesting that gas hydrate was present in the cores and dissociated prior to processing the samples. The chlorinity anomalies are consistent with observations of distinct negative thermal excursions in IR scans. Our observations further indicate that most of the gas hydrates occupy relatively thin sand layers. In the deepest zone, below 250 mbsf, the chlorinity remains nearly constant at 493 ± 3 mM, suggesting communication with a fluid at greater depth that is notably different in composition from the deep-seated fluid sampled at Sites U1327 and 889/890.

To form and maintain shallow gas hydrate deposits near the seafloor, methane has to be supplied continuously. Furthermore, the shallow elevated chlorinity values would dissipate rapidly if they were not continuously maintained by ongoing gas hydrate formation. However, the chlorinity profile with depth at Site U1328 indicates a diffusion-controlled system. These observations indicate that the methane needed to sustain the shallow gas hydrate formation is likely supplied along faults or fracture zones, probably the ones imaged by the RAB tool within the gas hydrate-bearing section immediately below the seafloor.

The shipboard organic geochemistry program for Site U1328 included analysis of hydrocarbons and nonhydrocarbon gases from headspace gas samples, void gas samples, gas samples recovered during PCS degassing experiments, and gases from gas hydrate samples. Methane was the most prominent hydrocarbon gas in all of the samples analyzed; however, ethane was also present in almost all of the headspace samples. Most of the gas hydrate samples and gas hydrate-bearing sediments collected within the upper 46 mbsf under the vent site exhibited elevated ethane concentrations. These near-surface samples also contained slightly more air contamination and elevated H₂S concentrations. It is also notable that the concentration of ethane and other gas hydrate-forming gases, including propane and isobutane, increase within the gas samples collected from the cores crossing the depth of the seismically inferred BSR where the chlorinity and IR temperature anomalies both indicate evidence for gas hydrate. The occurrence of propane and elevated isobutane to *n*-butane (*i*-C₄/*n*-C₄) ratio support the hypothesis that the gas hydrates near the depth of the BSR may contain a mixture of Structure I and Structure II gas hydrates. It is also interesting that the ethane concentrations remain elevated below the depth of the BSR. Structure II gas hydrate is stable to greater depths with increasing temperature than Structure I gas hydrate; however, the observed IR and chlorinity anomalies are not necessarily below the base of the Structure I methane hydrate stability field and therefore not automatically related to Structure II gas hydrate given the uncertainty in the temperature measurements. However, the gas geochemistry does strongly suggest the presence of some

mixture of Structure I and Structure II gas hydrates near the depth of the seismically inferred BSR.

Microbiological subsampling was routinely conducted on cores recovered from Holes U1328B and U1328C. On each core run PFT was continuously metered into the drilling fluid, and fluorescent microspheres were deployed on all cores in the continuously cored holes to investigate potential drilling fluid contamination of the core. These analyses confirmed that the center of each whole-round sample remains undisturbed for microbiological subsampling. Additional IR images were taken on the cut ends of each microbiological core section to document the thermal warming process of the core before subsampling.

The main objectives at this site were to

1. Document the depth distribution of gas hydrates and their relationship to fluid and gas chemistry, lithology, and faulting and
2. Determine the fluid advection rate feeding the surface vents.

Apparently, however, only the very shallowest (<40 mbsf) and deepest (250–300 mbsf) portion of the sedimentary section cored at Site U1328 is dominated by fluid advection. It appears that diffusion transport controls the chemical depth profiles from ~40 to 250 mbsf at this site. This apparent contradiction can be explained with the vent model brought forward by Riedel et al. (in press). The model, based on various geophysical and geochemical observations, predicted that the shallow gas hydrate accumulations to be the result of isolated feeder channels or fractures. Along these fractures the bulk of the methane gas is transported upward and results in the near-seafloor concentrated gas hydrate accumulation. Evidence of the existence of these fractures comes from the LWD/MWD downhole logging data showing at several depths steeply dipping resistivity anomalies typical of fractures filled with gas hydrate or possibly free gas. These potential gas migration conduits may be connected to the seafloor chemosynthetic cold vent communities observed by bottom video surveys (Riedel et al., 2002, in press).

The growth of gas hydrate near the seafloor has also resulted in solute exclusion and development of high pore water salinities. This brine cannot be maintained in situ over a long period of time, unless gas hydrate is constantly formed at relatively high rates, thus showing the cold vent must be supplied with fluids and gas from greater depths. This is also indicated by the detection of a methane plume in the water column just above an active vent outlet. However, venting is most likely episodic since

the methane plume is not always detectable. There is further evidence for the presence of Structure II gas hydrate at this site as seen by the presence of unusual amounts of propane and isobutane, which is unique when compared to the other sites cored during this expedition.

Various geophysical and geochemical data support the presence of a BSR at this site, as proposed by Riedel et al. (in press); however, the BSR at this site is difficult to identify with the available seismic data. The lack of good 3-D seismic velocity control could potentially result in misinterpretation of the existing two-dimensional (2-D) seismic data, in that the observed BSR may actually be a side diffraction and may not be exactly located beneath the active part of the cold vent and seismic blanking. In the LWD/MWD data a sharp decrease in electrical resistivity was observed at ~220 mbsf, although it is coincident with a high-density interval potentially related to the presence of carbonates. Acoustic wireline logging, however, shows a sharp discontinuity in *P*-wave and *S*-wave velocity at a depth of ~215 mbsf, close to the seismically inferred BSR depth of 219 ± 5 mbsf. The downhole temperature measurements from this site yielded a temperature gradient of ~54°C/km, which predicts the base of gas hydrate stability at $\sim 230 \pm 10$ mbsf for the observed pore water and gas geochemistries.

There is evidence for the occurrence of free gas at depths below the seismically inferred BSR, as demonstrated from pressure coring, LWD/MWD pressure monitoring, and elevated LWD/MWD electrical resistivity values. There are also seismic indicators for the occurrence of free gas below the BSR in the form of strong-amplitude anomalies. However, the VSP does not show any evidence of a velocity change over the entire logged interval from 105 to 285 mbsf, which will be further examined after the cruise.

Site U1329

Site U1329 (prospectus Site CAS-05D) is at the eastern end of the southwest-northeast-trending margin-perpendicular transect of sites occupied during this expedition and is located closest to shore (65 km) at a water depth of ~946 mbsl. The location of this site is interpreted to be at the eastern limit of gas hydrate occurrence on the Northern Cascadia margin.

The objectives of coring and downhole logging at this site are tied to completing the transect of scientific drill sites across the northern Cascadia margin to further con-

strain models for the formation of marine gas hydrate in a subduction zone accretionary prism. The depth to the BSR rapidly becomes shallower around this site and is only half the depth (~125 mbsf) than at Site U1327 (~225 mbsf) (Figs. F20, F21). At this eastern end-member site of gas hydrate evolution in the accretionary prism, the objectives include

1. Characterizing the distribution of gas hydrate,
2. Defining the nature of the BSR,
3. Developing baseline geochemical and microbiological profiles, and
4. Obtaining data needed to ground-truth remotely acquired imaging techniques such as seismic and CSEM surveys.

Five holes were occupied at Site U1329 (prospectus Site CAS-05D) (Table T1). Hole U1329A was dedicated to LWD/MWD measurements to a TD of 220 mbsf. Hole U1329B consisted of only one missed mudline APC core to 9.5 mbsf. Hole U1329C was continuously cored (17 APC, 5 XCB, and 3 PCS cores; recovery = 99.3%) to 189.5 mbsf, and was terminated before the TD of 220 mbsf when the PCS cutting shoe broke off and was left in the bottom of the hole. The PCS was deployed three times in Hole U1329C, and four APC temperature measurements were made (APCT and APC3). Swell from a passing low-pressure system resulted in ship heave that was too high for wireline logging or coring operations, and we waited 4 h for the weather to improve. With a forecast for improving weather the following day, we decided to abandon Hole U1329C to drill a dedicated logging hole. Hole U1329D was drilled from the seafloor to 201.0 mbsf, which included another 1.5 h of suspended operations due to excessive ship heave, and a single XCB core was taken to 210.5 mbsf. Hole U1329D was wireline logged with the triple combo and FMS-sonic tool strings. The triple combo and the first pass of the FMS-sonic were logged to 209 mbsf. The second pass of the FMS-sonic only reached 171 mbsf. Hole U1329E was a special tool hole to 127.1 mbsf, where five APC cores were taken for high-resolution microbiological and geochemical studies, with two additional APC temperature measurements. Three DVTP deployments were also conducted. Five pressure cores were taken (three PCS, one HRC, and one FPC) separated by drilled intervals. In Figures F22 and F23, data from LWD/MWD logging (RAB images, porosity and Archie-derived pore water saturations), core-derived physical property data (porosity), and interstitial water chlorinity, as well as C_1/C_2 ratios from void gas samples are shown in comparison to the two key crossing seismic lines from this site.

Site U1329 is located near the foot of a relatively steep slope, and sedimentation at this site is dominated by slope processes. The stratigraphy at Site U1329 was divided into three lithostratigraphic units.

Lithostratigraphic Unit I (0–37.18 mbsf in Hole U1329C and 0–33.31 mbsf in Hole U1329E) is characterized by fine-grained detrital sediments (clay and silty clay), locally interbedded with coarse-grained sediments. Authigenic carbonates are abundant in some cores from Unit I, including the location of dolomite in Holes U1329C and U1329E, which coincides with changes in pore water Ca and Mg profiles, suggesting dolomite formation is an ongoing process. The sedimentation rate in Unit I appears to be relatively slow at ~10–12 cm/k.y. The boundary between lithostratigraphic Units I and II is defined by the first occurrence of diatom ooze.

Lithostratigraphic Unit II (37.18–135.60 mbsf in Hole U1329C and 33.31–125.95 mbsf in Hole U1329E) is characterized by a high abundance of biogenic silica (mainly diatoms). The sedimentation rate within Unit II appears to increase relative to Unit III from ~4 to 10 cm/k.y. The boundary between lithostratigraphic Units II and III is marked by a conglomerate composed of partly lithified and lithified rounded clasts supported by a silty clay matrix. The conglomerate at the base of Unit II corresponds to an unconformity between upper Miocene and Pleistocene sediments (no sediments preserved from 2–6.7 Ma) that also can be traced seismically for several kilometers along Line 89-08 (Fig. F20). The unconformity is not clearly seen in the higher-frequency seismic data crossing Line SCS CAS05C-01-04 (Fig. F21).

Lithostratigraphic Unit III (135 mbsf–TD) is characterized by fine-grained (clay to silty clay) detrital sediments with only a few coarser interlayers from turbiditic deposits. The input of detrital sediments from turbidites, however, is more obvious deeper in the section with the presence of nonmarine diatoms. In the lowermost part of Unit III, a breccia deposit marks a major change in deposition, possibly representing a debris flow. The sediments in lithostratigraphic Unit III were deposited at a low sedimentation rate of ~0.8–2.8 cm/k.y.

The biostratigraphy at this site was established by the analysis of diatoms. These analyses showed an interval of Quaternary sediments from the surface to ~35 mbsf that corresponds to lithostratigraphic Unit I. The age of lithostratigraphic Unit II was shown to range from 0.3 to 2.0 Ma. This interval showed an abundance of diatoms and resting spores suggesting a depositional environment typical of a coastal shelf area with upwelling. The sediments of this interval may also have been transported

to this site from the shelf by turbidites. Sediments deeper than 137 mbsf, near the contact between lithostratigraphic Units II and III, contain diatoms of only late Miocene age (>6.7 Ma); thus, the Pliocene at this site is completely absent, representing ~4 m.y. of missing sediment, which was likely eroded.

The downhole logged section at Site U1329 was divided into three “logging units” based on obvious changes in the LWD/MWD and wireline gamma ray, density, electrical resistivity, and acoustic transit-time measurements.

Logging Unit 1 (0–130 mbsf), which corresponds to lithostratigraphic Units I and II, is characterized by low electrical resistivities and densities.

Logging Unit 2 (130–183 mbsf) is characterized by a small increase in resistivity and density compared to logging Unit 1. The top of logging Unit 2 corresponds to the bottom of lithostratigraphic Unit II, which was identified as an unconformity separating Pleistocene from upper Miocene sediments.

Logging Unit 3 (183 mbsf–TD) is characterized by an abrupt increase in electrical resistivity, density, and *P*-wave velocity, indicating an increasingly consolidated formation with lower log-derived porosities. Logging Units 2 and 3 correspond to lithostratigraphic Unit III.

Variations in physical properties do not correlate well with the lithostratigraphic and logging units. MAD analyses of porosity and bulk density as well as gamma ray attenuation density data from the multisensor track show anomalously uniform trends throughout most of the section cored from 25 to 180 mbsf. A slightly different trend is observed in the upper 25 mbsf where porosity rapidly decreases to values as low as 55%, coincident with the depth where rapid changes in the core-derived shear strength and electrical resistivity occur. Except for the top 15 mbsf, the magnetic susceptibility record shows extremely low and uniform values of $<50 \times 10^{-7}$ SI. Grain densities generally decrease downhole crossing all lithostratigraphic boundaries starting from values of ~ 2.8 g/cm³ near the seafloor to ~ 2.6 g/cm³ at 175 mbsf, but this trend reverses below this depth and again reaches values as high as ~ 2.8 g/cm³. This deep increase in grain density and associated increase in bulk density occurs near the top of logging Unit 3, which is characterized by a sudden increase in electrical resistivity and drop in porosity. Analysis of the LWD/MWD density logs from this site yielded sediment porosities ranging from $\sim 65\%$ near the seafloor to $\sim 23\%$ at 220 mbsf.

IR imaging of the recovered core assisted in immediate gas hydrate detection on the catwalk. At this site, core IR temperatures did not show any significant cold spot anomalies from gas hydrate dissociation that could be related to the presence of gas hydrate in the recovered core. In situ temperature measurements with the APC3, APCT, and DVTP tools indicated a temperature gradient of $72^{\circ} \pm 4^{\circ}\text{C}/\text{km}$ and a sea-floor intercept of $3.34^{\circ} \pm 0.34^{\circ}\text{C}$. This implies a heat flow of 68–84 mW/m² for measured thermal conductivities of 1.0–1.1 W/m·K. This is consistent with the regional heat flow across the northern Cascadia margin (predicted to be ~ 72 mW/m² here) but higher than the heat flow measured at Site 889 during Leg 146, indicating that heat flow at this site is not affected by sedimentation and fluid expulsion.

In total six PCS deployments were made, three in Hole U1329C and three in Hole U1329E. Of the six deployments, two runs did not recover sediment under pressure and all other runs retrieved sediments under pressure. All PCS cores that were successfully retrieved under pressure were investigated by shipboard degassing experiments. In addition to the PCS, one HRC and one FPC were deployed in Hole U1329E. A full core at near in situ pressure was recovered by the HRC that revealed 10 cm high-velocity zones indicative of gas hydrate. The FPC deployment recovered a core without pressure. Degassing of the four PCS cores from this site that were recovered under pressure showed variable gas concentrations with depth. PCS Core 311-U1329C-7P (56.1 mbsf) and Core 311-U1329E-7P (74.0 mbsf) from within the predicted depth of the methane hydrate stability zone had total pore space methane concentrations of 79 and 85 mM, respectively, which would equate to a gas hydrate pore volume concentrations of <1% for both core samples. PCS Core 311-U1329E-10P (125.5 mbsf), from near the base of the predicted methane hydrate stability zone for this site, had a methane concentration of 89 mM, which falls below methane saturation at this depth, with no gas hydrate or free gas. The deepest PCS core (Core 311-U1329C-23P; 188.5 mbsf) was taken at a considerable depth below the base of the gas hydrate stability zone. This deep PCS had a total pore space methane concentration of 152 mM, which suggests a potential pore volume free gas saturation of 2.4%. A potential gas release during the coring operation may cause this to be an overestimate, yet it still supports the presence of free gas below the BSR at this site.

Interstitial water analyses were carried out on samples from Hole U1329C with additional high-resolution sampling in the top 15 mbsf of Hole U1329E to capture the SMI. Interstitial water salinities decrease from ~ 34 to ~ 30 at a depth of ~ 50 mbsf but remain almost constant at values of ~ 30.5 until near the bottom of the cored section in Hole U1329D. The deepest water samples from this site (>180 mbsf) show a sharp

drop in salinities, ranging from 22 to 28. A distinct chlorinity minimum (~15 mM decrease relative to the rest of the trend) is observed at the projected depth of the seismically inferred BSR. Assuming that this chlorinity minimum is the result of gas hydrate dissociation during core recovery and handling, the observed chlorinity minimum would equate to ~2% in situ pore space filling of gas hydrate. The strong decrease in chlorinity and salinity below 180 mbsf are coincident with the pronounced increase in electrical resistivity in the LWD/MWD downhole logging data in logging Unit 3 and suggests communication with a deeper-sourced fresher fluid. The Ca and Mg depth profiles reflect carbonate and dolomite formation in the shallow section (SMI-60 mbsf) but also suggest commingling with a deeper-sourced fluid over a depth range from 60 mbsf to the BSR.

Organic geochemical studies at Site U1329 included analysis of the composition of volatile hydrocarbons (C_1 - C_5) and nonhydrocarbon (i.e., O_2 and N_2) gases from head-space gas samples, void gas samples, and gas samples recovered during PCS degassing experiments. The predominant hydrocarbon gas found in the cores from Site U1329 was methane; however, we did see an increase in ethane concentrations in the void gases collected from the stratigraphic section overlying the projected depth of the BSR. This increase in ethane concentrations near the BSR can also be seen in the plot of the C_1/C_2 void gas ratios. Other studies have shown void gas ratios (such as C_1/C_2) can be elevated in a gas hydrate with the preferential selection of ethane into the gas hydrate structure. But the observed increase in the C_1/C_2 ratio below the projected depth of the BSR may be more indicative of a methane-enriched free gas accumulation. The C_1/C_2 ratios were generally high, suggesting a microbial origin for the observed methane. The samples collected from deeper than 180 mbsf at this site also exhibited significant increases in C_2 - C_5 concentrations, which suggests the influence of a potential thermogenic hydrocarbon source.

Microbiological subsampling was routinely carried out on cores recovered from Hole U1329C. On each core run PFT was continuously metered into the drilling fluid, and additional fluorescent microspheres were deployed on selected cores to investigate potential drilling fluid contamination of the core. These analyses confirmed that the center of each whole-round sample remains undisturbed for microbiological subsampling. Additional IR images were taken on the cut ends of each microbiological core section to document the thermal warming process of the core before subsampling.

When considering the primary research objectives of this site, one of the most important features is its location on the relatively steep slope and the observation that sed-

imentation at this site is dominated by both slope turbidite deposition and erosional processes. The presence of thick turbidite sequences, including the presence of non-marine diatoms in the stratigraphic section, is a testament to the history of slope-dominated processes. The occurrence of a significant depositional hiatus or erosional event that juxtaposes Miocene with Pleistocene sediments at this site further documents the complex depositional history of this margin. The deepest core and down-hole logging penetrations at this site also revealed even a more complex geologic history with unique high-resistivity, low-porosity conglomerate deposits that may represent thick debris flows.

Geochemical analyses of recovered gas samples and interstitial fluids have also documented a relatively complex fluid regime for this site. The relatively low interstitial fluid chlorinities observed near the seafloor may reflect paleoceanographic changes in seawater salinity. However, the more rapid Cl concentration decrease with depth strongly suggests communication with a deep-seated, fresher fluid, especially below 140 mbsf. For the most part, geochemical analyses of gases from this site reveal a similar story as the analysis of interstitial fluids, with most of the cored section dominated by a simple methane gas-rich system. But near the bottom of the cored section there are indications for the influence of a deeper hydrocarbon source.

Identifying gas hydrate and supportive physical evidence of the BSR at this site remains an elusive goal. As seen in Figures **F20** and **F21**, the BSR is subtle at this site and is located only ~10 m above a prominent unconformity, which results in a complicated reflectivity pattern. The precoring LWD/MWD and conventional wireline logging data were used to identify and further characterize potential gas hydrate and related free gas occurrences at each site occupied during this expedition. The presence of gas hydrate is generally characterized by increases in measured electrical resistivities and acoustic velocities. The relative lack of notable resistivity or acoustic down-hole log anomalies at this site supports a general inference of very low to no gas hydrate occurrence at this site. The lack of either visual observations of gas hydrate or the occurrence of IR-detected thermal anomalies along the recovered cores also supports the assumption that this site contains little to no gas hydrates. However, geochemical analyses of interstitial fluids from near the depth of the predicted BSR may suggest the presence of a small amount of gas hydrate, with minimum dissolved chloride concentration at ~125 mbsf, coincident with the depth of the BSR. The HRC core taken from 114.6 mbsf, which contained high-velocity layers, also suggests the presence of gas hydrate at these depths. In addition, the relative increase in ethane concentrations immediately above the depth of the BSR may indicate the presence of

gas hydrate. Degassing experiments of three PCS cores that were collected from within the zone of predicted methane hydrate stability appear to confirm that this site contains very little gas hydrate, if any. However, a PCS core from 188.5 mbsf in Hole U1329C does suggest the occurrence of a free gas phase. The presence of free gas at greater depth below ~160 mbsf may also be supported by a distinct change in seismic reflection character in the low-frequency seismic data of Line 89-08 (Figs. F21, F23). In contrast to the shallower section of the seismic line above 1.5 s two-way traveltime (TWT), no clear reflection can be identified below and seismic frequency also is strongly reduced.

As noted, this site was intended to characterize the eastern limit of gas hydrate occurrence on the northern Cascadia margin and establish an understanding of the geologic controls on the occurrence of gas hydrates along the transect of sites established during this expedition. All of the objectives set for Site U1329 are considered fulfilled.

SUMMARY AND IMPLICATIONS

The specific objectives of this expedition were to test gas hydrate formation models and constrain model parameters, especially models that account for the formation of concentrated gas hydrate occurrences through upward fluid and methane transport. During Expedition 311 a series of research sites across a transect through the Northern Cascadia margin was established by drilling and coring four sites extending from the westernmost accreted ridge (Site U1326, representing the youngest gas hydrate formation) and the first slope basin (Site U1325) to the second accreted ridge midslope (Site U1327) and a site at the eastern boundary of the inferred gas hydrate occurrence along this margin (Site U1329). We also cored and logged a cold vent (Site U1328), which represents a site of focused fluid and gas flow.

We collected the necessary data for characterizing gas hydrate proxies such as interstitial water chlorinities, core-derived gas chemistry, core physical properties, and downhole measured logging data. Geophysical characterization of each site was accomplished by the LWD/MWD program, conventional wireline deployments, and two VSP experiments. Extensive evidence for the presence of gas hydrate at all sites was collected from electrical resistivity and *P*-wave velocity logs, IR imaging, interstitial water and gas sampling, as well as from direct sampling of visible gas hydrate (Figs. F24, F25).

The main objectives of this cruise are considered fulfilled with only few elements having only partial success rates, such as temperature tool deployments and pressure coring, which are mainly a result of poor weather conditions.

Occurrence of Gas Hydrate

Bottom-Simulating Reflectors

The occurrence of gas hydrate has historically been inferred from the presence of a BSR in seismic images (Shipley et al., 1979). BSRs as ascribed to gas hydrates were first mentioned in the literature in 1977 from the Blake Ridge (e.g., Tucholke et al., 1977) and have since been a tool for gas hydrate detection. The BSR is the result of an impedance contrast generated by the transition from gas hydrate-bearing sediments above to gas hydrate-free and potentially free gas-bearing sediments underneath the interface. The presence of gas hydrate in sediments increases seismic *P*-wave velocity, whereas the presence of free gas (even small amounts) drastically reduces seismic *P*-wave velocity. The density is not expected to change much across this interface because the density changes introduced by the presence of gas hydrate and free gas are relatively small. The BSR reflection shows a phase-reversed polarity relative to that of the seafloor due to the negative impedance contrast at the interface.

The BSR has also been described as a frequency-dependent reflection (Chapman et al., 2002), which results from a finite-thickness gradient zone, in which velocity decreases from the elevated values above to low values below the interface. In conventional seismic surveys with low frequencies, the BSR appears as a strong reflection (see, e.g., Fig. F3), but with increasing frequency the magnitude of this reflection decreases.

BSRs have been widely observed in seismic data collected across the northern Cascadia margin (Hyndman and Spence, 1992; Hyndman, 1995) and were used to map the general distribution of gas hydrates along this margin (Fig. F1). At all sites visited during Expedition 311, BSRs were previously imaged with various seismic surveying techniques ranging over frequencies from 20 to 650 Hz (for a summary, see Hyndman et al., 2001). The depth to the BSR is a critical element in gas hydrate research because it is usually a measure of the base of the gas hydrate stability zone. However, it has also been shown that the BSR can occur away from the actual base of the stability zone at shallower or even deeper depths (Xu and Ruppel, 1999), where it depicts the first occurrence of free gas in the subsurface. Gas hydrate can also occur below the BSR, or even at a second BSR, in the form of Structure II gas hydrate.

The depth to the BSR in TWT has been determined from several crossing sections at each site visited during the expedition, and the average time was then converted to depth using velocities defined at Sites 889/890 from a VSP (Mackay et al., 1994). The BSR at Site 889 occurs at a depth of ~275 ms TWT, equivalent to a depth of 225 mbsf with the average velocity of 1636 m/s for the sediment column from seafloor to the BSR. This velocity was used across the entire margin for all sites visited and thus introduces a significant uncertainty in the depth assignment of the BSR. We therefore defined two extreme velocity profiles that describe sediments without gas hydrates (average velocity = 1619 m/s) and with concentrations well above what was observed (and inferred) at Site 889 (average velocity = 1653 m/s). The velocity increase appears very small; however, this is the result of only a thin sediment column over which the presence of gas hydrate was inferred (i.e., only 100 m above the BSR), leaving 125 m of sediment above this non-gas hydrate-bearing column. The range in velocity results then in a shift in the BSR depth by about ± 2.5 m.

Throughout Expedition 311, special effort was made to better define the depth of the base of gas hydrate stability. Among the proxies used to define this boundary and to compare it to the seismically inferred BSR depth are

- Downhole temperature measurements,
- Well-log measurements of *P*-wave velocity and electrical resistivity,
- Pore water chlorinity,
- IR imaging,
- C_1/C_2 ratios of the void gas combined with the occurrence of propane and butane, and
- Visual core descriptions and notes on gas hydrate-related sediment textures.

There are notable differences between the individual techniques, as they are highly dependent on how the measurements are conducted and the resolution or sensitivity of a particular measurement and as they are biased by core recovery (IR, pore water chemistry, and gas chemistry) and sampling density (e.g., frequency of temperature tool deployments and linear regression analysis).

Gas Hydrate Stability Calculations

Expedition 311 featured 31 temperature tool deployments in an attempt to characterize the thermal regime of sites drilled along the Expedition 311 transect (Table T2). Three standard IODP temperature tools were deployed during the expedition, includ-

ing the APCT (nine times), DVTP (nine times), and DVTPP (five times). We also deployed the new APC3 tool a total of eight times. Figure F26 shows a compilation of all in situ temperature estimates from Expedition 311 compared to results from Site 889. The implied heat flow, assuming a constant thermal conductivity of 1.1 W/m-K, is also shown and compared to the regional heat flow determined by Hyndman and Wang (1993). Heat flow across the lower slope (Sites U1325, U1327, and U1328) appears to be depressed compared to the regional heat flow pattern, consistent with perturbation by a high sedimentation rate and upward fluid advection (Hyndman and Davis, 1992). Postcruise analysis will focus on detailed examination of data uncertainties and on processes to explain inter- and intrasite variation.

The primary goal of the temperature tool deployment program during Expedition 311 was to obtain the data needed to calculate the depth to the base of the gas hydrate stability zone at each of the sites visited during the expedition. Gas hydrate exists under a limited range of temperature and pressure conditions such that the depth and thickness of the zone of potential gas hydrate stability can be calculated. Most gas hydrate stability studies assume that the pore pressure gradient is hydrostatic (9.795 kPa/m). However, the seafloor temperature and geothermal gradient for any given site can be highly variable. The temperature data acquired during this expedition have been used to make a preliminary estimate of the depth to the base of the gas hydrate stability zone at each site. In these calculations a pure methane gas chemistry was assumed for the gas hydrate, whereas the interstitial water salinities from the onboard analysis of core samples were used to estimate the depth to the base of the methane hydrate stability zone at each site. The results of these preliminary calculations have been reported as range of depths for each site:

- Site U1325 = 250–300 mbsf.
- Site U1326 = 250–270 mbsf.
- Site U1327 = 225–250 mbsf.
- Site U1328 = 220–245 mbsf.
- Site U1329 = 127–129 mbsf.

For the most part, the calculated depth to the base of the methane hydrate stability zone for each site falls near the estimated depth of the BSR as inferred from seismic data. In the case of Sites U1325 and U1326, however, it appears that the estimated depth of the BSR is less than what would be expected from the gas hydrate stability calculations. It should also be noted that the linear regression used to estimate the subsurface temperatures in the above calculations may not be appropriate for envi-

ronments that are affected by fluid advection (e.g., the cold vent Site U1328). A more detailed temperature measuring program is required to better resolve the base of the gas hydrate stability zone and will be addressed in the Phase II drilling under this project by deployments of distributed fiber-optic temperature cables that allow continuous recording of the temperature field on a vertical sampling density of <1 m.

Core Analysis

Visual Descriptions

The occurrence of gas hydrate can be documented by a variety of methods. The most obvious and fundamental method is by direct visual observation of cores. During Expedition 311, visual observation of gas hydrates was primarily associated with the cold vent location, Site U1328, where a thick section of highly concentrated gas hydrate-bearing strata was inferred from LWD/MWD and previous precruise studies. From cores in this interval, numerous gas hydrate pieces were recovered from the upper ~35 mbsf, some measuring up to 5–8 cm in diameter (Fig. F25). Gas hydrate was also observed at other sites, more typically seen as small nodules or filling pores within sandy layers (Fig. F24).

Because gas hydrates, especially disseminated gas hydrates, begin to dissociate during core recovery, visual observation may not be possible because they will have decomposed before the cores can be processed on deck. Thus, indirect or proxy methods related to the physical and chemical consequences of dissociation must be used to document gas hydrate occurrences.

Gas hydrate dissociation releases gas and water during core retrieval, which should physically disturb the original sedimentary fabric. Previous observations have indicated two textures that are a product of dissociation (Westbrook, Carson, Musgrave, et al., 1994; Kastner et al., 1995; Lorensen et al., 2000). Soupy texture describes intervals that are water saturated and have lost all primary sedimentary structures, which are thought to represent areas that once contained larger pieces or veins of gas hydrate. Intervals with many small, round voids are termed mousseliike and are interpreted to represent dissociation of disseminated gas hydrates. During Expedition 311, soupy and mousseliike textures were common and were associated with intervals where direct and other indirect evidence of gas hydrate were noted. Site U1329 was the only site where these textures were not observed. At the cold vent Site U1328 in the upper shallow gas hydrate zone (lithostratigraphic Unit I), soupy texture was common, but not below this unit, whereas mousseliike textures were observed in the lower

units. Sites U1327, U1326, and U1325 also noted the presence of soupy and mousse-like textures. At Site U1326 the lower part of a section with a notable IR anomaly (Section 311-U1326-18X-4) was sectioned and opened immediately. The sediment in this core had a very unusual texture, remaining cohesive and “foamy” rather than “soupy” or “moussy.” Sediment in this section lost its foamy texture in less than an hour, collapsing to a stiff, dry sediment that occupied an estimated 25% of the original volume in the core liner.

Interstitial Water Chemistry and Infrared Imaging

The presence of gas hydrate in situ has often been inferred by freshening of pore water chlorinity/salinity relative to a defined baseline (e.g., Kastner et al., 1995). When gas hydrate forms in the subsurface the process initially increases pore water chlorinity/salinity because gas hydrate formation excludes the salt from the crystal structure. Over geologic time, the excess salt is removed by advective and diffusive processes and the background pore water chemistry is reestablished. In case of rapid and relatively recent gas hydrate formation, this process is incomplete and high-pore water chlorinity/salinity brines are observed (e.g., Site 1249 of Southern Hydrate Ridge during ODP Leg 204; Trehu, Bohrmann, Torres, Rack, et al., 2003). Such a system was also cored at the cold vent Site U1328 during this expedition.

If a core is retrieved on the catwalk that had gas hydrate present in situ, then gas hydrate has most likely completely dissociated, leaving freshwater behind, which can be detected by interstitial water analyses. A summary of all interstitial water chlorinity profiles linked to LWD/MWD RAB-derived borehole images is shown in Figure F27. The definition of an interstitial water chlorinity baseline has proven difficult because the effect of pore water freshening from gas hydrate dissociation and advection of fresher pore fluids from greater depth cannot easily be distinguished with chlorinity and salinity measurements alone. Additional shore-based pore water analyses, such as strontium, boron, or iodine isotopes, are required to distinguish these processes (Teichert et al., 2005; Fehn et al., *in press*).

The catwalk procedure for collecting interstitial water samples during this expedition had been adapted from previous “gas hydrate” expeditions and was strongly linked to IR imaging. The presence of gas hydrate in the core can be detected (if not visually) by a decrease in core temperature due to the cooling upon gas hydrate dissociation. One of the early systematic approaches to core temperature measurements was during Leg 164 using an array of digitally recorded thermocouple probes (Paull, Matsumoto, Wallace, et al., 1996). Initial development of the IR imaging technique was accom-

plished during Leg 201, where thermal anomalies in IR images were associated with gas hydrate and voids (Ford et al., 2003). Systematic IR thermal imaging of the surface of the core liner was first fully implemented during Leg 204 (Trehu, Bohrmann, Torres, Rack, et al., 2003). During Expedition 311 a slightly different version of the IR system was used to systematically scan the core for IR cold spot anomalies. A summary of all IR profiles from continuous coring are compared to Archie-derived pore water saturations (S_w) and LWD/MWD RAB-derived borehole images, shown in Figure F28.

Generally, several interstitial water samples were collected from most cores that did not show IR cold spot anomalies to define the in situ background pore water chemistry. In addition, we directly sampled IR-imaged cold spots in cores to measure the amount of pore water freshening associated with gas hydrate-related IR temperature anomalies. This will help to further calibrate the IR measurements and delineate anomalous trends from the background pore water chemistry. Experiments during Leg 204 showed that cold spot-related interstitial water freshening is focused spatially to a range of a few centimeters around the IR anomaly (Trehu et al., 2004). Each IR cold spot sampled was therefore carefully analyzed during this expedition. Once in the chemistry laboratory, the handheld IR camera was used to isolate the cold anomaly so a subsample of the sediment that had contained gas hydrate could be separated from background sediment (Fig. F25). The cold material was squeezed separately from the background material. In many cases, it was found that the sediment hosting the cold spot was coarser-grained sands and silts and the surrounding clay did not exhibit freshening.

Gas Chemistry

Organic geochemical studies focused on volatile hydrocarbon gases from headspace samples, void gases, and selected gas hydrate samples. At all sites, analyses showed that methane is the dominant hydrocarbon gas within the cored intervals. However, the presence and distribution of higher hydrocarbons (C_{2+}) provided useful information about the occurrence and possibly the type of gas hydrate. In particular, studies have shown that gas hydrate forms with the preferential selection of ethane into the gas hydrate structure, which is released during dissociation. At most sites, where other visual or proxy data indicated the presence of gas hydrates, void gases and selected gas hydrate and PCS samples contained elevated ethane concentrations and decreasing C_1/C_2 ratios (Fig. F29). In some intervals, elevated propane and isobutane were

also observed, which suggest there may also be Structure II gas hydrate present in the cored interval as well.

Downhole Logging Analysis

Because gas hydrate is characterized by unique chemical compositions and distinct electrical resistivity and acoustic physical properties, it is possible to identify and further characterize gas hydrate-bearing sediment properties, including porosities and gas hydrate saturations, with downhole logging tools commonly deployed from the ship during IODP. In the case of Expedition 311, the standard downhole wireline logging program was augmented with a precoring LWD/MWD dedicated drilling-logging program. The downhole LWD/MWD logging data acquired during the start of this expedition was used to direct the special pressure coring program and make decisions on other critical special tool deployments. The LWD/MWD data, along with downhole wireline logging data, were also used to identify and quantify the occurrence of gas hydrate at each site drilled along the Expedition 311 transect.

As shown in Figure F30, the LWD/MWD-derived RAB images of borehole resistivities reveal numerous apparent high-resistivity intervals (shown as light color bands) within the stratigraphic section above the projected depth of the BSR. Also shown in Figure F30 are the Archie-derived water saturations (which is the mathematical complement of gas hydrate saturation) for each of the LWD/MWD-logged holes. Except for the cold vent Site U1328, the downhole logging-inferred gas hydrate occurrences appear to be first observed at a minimum depth of ~70–100 mbsf; and in most cases we see evidence for gas hydrates to about the base of the predicted methane hydrate stability zone. The apparent low water saturations below the BSR likely depict the occurrence of free gas. The downhole logging data from Site U1329 revealed relatively little evidence of gas hydrate. The downhole logging data from Sites U1327 and U1326, however, are dominated by the appearance of thick high-resistivity intervals near a depth of 100 mbsf at each site. In the case of Site U1327, we see an 18 m thick (120–138 mbsf) high-resistivity interval with peak LWD/MWD resistivity values exceeding 12 Ωm , whereas in Hole U1326A we see a 35 m thick interbedded section (72–138 mbsf) of gas hydrate-bearing sands and non-gas hydrate-bearing clay sections. The logging-measured resistivities in gas hydrate-bearing sands at Site U1326 exceed 40 Ωm for part of the section. Archie analyses of these high-resistivity intervals at both sites indicate gas hydrate saturations possibly exceeding 50%–75% of pore volume. The gas hydrate occurrence at Site U1325, however, appears to be distributed throughout the cored section of the site with the gas hydrate concentrated in thinly

bedded sand units measuring between 5 and 20 cm thick. The downhole LWD/MWD logging data from Site U1328 reveal a very distinct zone of high resistivity extending from the seafloor to a depth of ~50 mbsf. Archie analysis of the resistivity logging data from Site U1328 indicates that this near-surface high-resistivity interval may contain a significant amount of gas hydrate. The closer examination of the RAB images from the cold vent Site U1328 also reveals numerous high-angle fractures cutting through the shallow gas hydrate-bearing section, which may have served as migration conduits for gas from potential deeper sources.

Pressure Core Analysis

Expedition 311 had the most ambitious pressure coring and onboard pressure core analysis program ever attempted in the history of ocean drilling. Pressure cores retrieved at in situ pressures were used to determine methane hydrate quantity using degassing techniques and mass balance calculations and methane hydrate distribution using nondestructive measurement of the physical properties of the cores at in situ pressures. Large improvements in temperature control over previous expeditions (e.g., Leg 204; Tréhu, Bohrmann, Torres, Rack, et al., 2003) made the recovery and analysis of pressure cores more practical. Pressure coring is crucial for understanding the concentration of gas hydrate and free methane gas in marine sediments, their nature and distribution, and their effect on the intrinsic properties of the sediment.

Pressure cores were collected using the IODP PCS and HYACINTH FPC and HRC pressure corers. The PCS is a downhole tool designed to recover a 1 m long sediment core with a diameter of 4.32 cm at in situ pressure up to a maximum of 69 MPa (Pettigrew, 1992; Graber et al., 2002). Two types of wireline pressure coring tools were developed in the European Union-funded HYACE/HYACINTH programs: a percussion corer (FPC) and a rotary corer (HRC), which were designed to cut and recover core in a wide range of lithologies where gas hydrate-bearing formations might exist.

During Expedition 311, the PCS retrieved pressurized sediments for onboard degassing experiments. Controlled release of pressure from the PCS through a manifold permits

- Collecting all gas discharged from the sediment's free gas and gas hydrate phase for quantitative and qualitative analyses,
- Estimating the in situ abundance of gas hydrate and free gas based on mass balance, methane solubility, and gas hydrate stability considerations (Dickens et al., 1997),

- Identifying the presence of gas hydrate from volume-pressure-time relations (Hunt, 1997; Dickens et al., 2000; Milkov et al., 2004), and
- Monitoring the controlled decomposition of gas hydrate with non-destructive methods in the course of the degassing experiment.

A total of 24 PCS runs were carried out during Expedition 311, of which 16 were successful and returned sediment under pressure. Each PCS was X-rayed before and after the degassing experiment for core characterization and to guide subsampling for interstitial water and gas headspace. Gas hydrate and free gas concentration estimates were determined from the mass balance calculations and are superimposed on the Archie-derived water saturation logs in Figure F30. The PCS-derived pore water gas hydrate and free gas concentrations agree well with estimates from LWD/MWD data. The results of all 16 successful PCS measurements are displayed in Figure F31, showing the calculated methane concentrations as function of depth and their position relative to the methane hydrate phase and methane solubility boundaries defined from the Xu model (Xu, 2002, 2004). The largest amount of gas hydrate within a PCS core was 40% of the pore space as inferred from degassing of Core 311-U1326C-12P, recovered from a depth of 83.7 mbsf within a high-electrical resistivity zone (Fig. F30). Other significantly large amounts of gas hydrates were inferred from Cores 311-U1328B-4P (15%) and 311-U1328E-10P (22%) at the cold vent (Table T3). A very large amount of free gas was inferred from the degassing of Core 311-U1328E-13P (58%) recovered from ~233 mbsf near the base of the gas hydrate stability zone. However, the core did not show any typical degassing features after the experiment, such as gas expansion cracks or large voids combined with mousseliike or soupy texture; it was concluded that the free gas was accidentally trapped within the outer chamber of the PCS and as it flowed into the borehole from a free gas-bearing formation.

In total, three HRC cores and one FPC core were stored under in situ pressures for future shore-based analyses. Onboard Expedition 311 these cores were X-rayed and analyzed with the multisensor core logger to measure density and *P*-wave velocity.

Remote Sensing Calibration

The presence of gas hydrate in marine sediments is mainly inferred from the presence of BSRs on seismic sections. Quantification of gas hydrate concentrations from remote sensing techniques such as seismic or CSEM methods was attempted earlier at this margin (e.g., Yuan et al., 1996, 1999; Yuan and Edwards, 2001; Hyndman and Spence, 1992), but these efforts need further calibration from borehole data, espe-

cially *P*-wave velocity, density, porosity, and electrical resistivity. Although borehole data may provide the needed calibration data, strong lateral heterogeneity in the derived gas hydrate concentrations and physical properties between adjacent holes only several tens of meters apart was observed at multiple sites visited during this expedition. Structural control of the gas hydrate occurrence, local dip of the formation, and other tectonic complications make estimates of gas hydrate concentration from remote-sensing techniques a challenging task and will be addressed postcruise in various studies.

Gas Hydrate Geologic System

In recent years significant progress has been made to address some of the key issues on the formation, occurrence, and stability of gas hydrates in nature. The concept of a “gas hydrate geologic system” approach is gaining some acceptance. In this approach the individual factors that contribute to the formation of a gas hydrate occurrences, such as gas source, gas and water migration, emplacement, and growth of the gas hydrate in a suitable host sediment, can be identified and quantified.

It has been shown that the availability of large quantities of hydrocarbon gas from either microbial or thermogenic sources or both is an important factor controlling the formation and distribution of natural gas hydrates (Kvenvolden, 1988). Carbon isotope analyses indicate that the methane in most oceanic gas hydrates is derived from microbial sources. However, molecular and isotopic analyses indicate a thermal origin for portions of the gas in several of the gas hydrate occurrences on the Cascadia margin (Pohlman et al., 2005). Additional factors controlling the availability of gas are the geologic controls on fluid migration. If effective migration pathways are not available, it is unlikely that a significant volume of gas hydrate would accumulate. Therefore, geologic parameters such as water and gas chemistry, as well as sediment permeability and the nature of faulting, must be evaluated to determine if the required gas and water can be delivered to the sedimentary section potentially hosting the gas hydrate.

Analysis of interstitial waters during Expedition 311 gave us important insight to fluid and gas sources and their movement along the Cascadia margin. At Sites U1327 and U1329 we see general decreases in salinity and chlorinity with depth, suggesting dilution. It was concluded that the dilution profile observed is primarily dominated by diffusive communication with an advective low-chlorinity fluid system at greater depth and not only to gas hydrate dissociation. In addition to chlorinity, the concen-

trations of Na, K, Ca, and Mg also decrease with depth, supporting the suggestion of communication with a deeper-sourced fresher fluid. At the vent Site U1328, salinities and chloride concentrations remain almost constant below the highly concentrated surface section of gas hydrate, again suggesting communication with a fluid at greater depth. But in this case the waters have undergone less freshening than at Site U1327. The salinity and chlorinity profiles at Sites U1326 and U1325, however, trend toward concentrations higher than seawater, which was interpreted as indicative of low-temperature diagenetic reaction. A plausible candidate for such a reaction is the alteration of volcanic ash to clay minerals and/or zeolites. These reactions consume water and, hence, increase in situ the chlorinity values. The analysis of core-derived hydrocarbon gases along the Expedition 311 transect, when linked to interstitial water chemistries, also show communication with a deeper source.

Hydrocarbon headspace gas measurements show that methane is the dominant hydrocarbon gas within the cored interval at every site along the transect. Although ethane and more complex hydrocarbon gas concentrations were low, the occurrence of more complex hydrocarbon gases generally indicates the contribution of a deeper gas source, much like the interstitial water chemistry results. It was also shown that the methane needed to sustain the shallow gas hydrate formation at the Site U1328 vent site was likely supplied along faults or fracture zones from a deeper source. It is also notable that the concentration of ethane and other gas hydrate-forming gases, including propane and isobutane, increase within the gas samples collected from the cores crossing the depth of the seismically inferred BSR at several sites. This gas geochemistry also strongly suggests the presence of a mixture of Structure I and Structure II gas hydrate.

It can be shown that by integrating the results of the interstitial water and gas geochemistry studies that the occurrence of gas hydrate along the Expedition 311 transect is controlled in part by the upward advection and diffusion of methane and other gas hydrate-forming gases from a deeper source. In the case of vent Site U1328 we see the influence of a focused flow system that has contributed to the development of a localized concentrated gas hydrate accumulation.

For the most part, the interpretation of downhole logging data and linked IR imaging and interstitial water analyses from the sites drilled along the Expedition 311 transect indicate that the occurrence of concentrated gas hydrate is mostly controlled by the presence of sand-rich sediments, which provide the conduits for gas migration and the containers in which for gas hydrate to grow. The occurrence of the relatively

thick, highly concentrated gas hydrate accumulations within 100 mbsf at Sites U1326 and U1327 are also partially controlled by the presence of suitable host sands; but in these two cases we may also be seeing the effect of regional gas solubility in water on the occurrence of gas hydrates. We see relatively little evidence of gas hydrate or gas above these intervals at any other site visited during Expedition 311, other than the Site U1328 vent site at which free gas has been observed venting from the seafloor (Fig. F32).

Thus, preexpedition models that had predicted the occurrence of a highly concentrated gas hydrate section (50–100 m thick) just above the base of the gas hydrate stability zone do not take into account the geologic complexity revealed during Expedition 311, in that the occurrence of gas hydrate appears to require the contribution from mixed gas sources, including a potential deep thermogenic gas source. The occurrence of migration conduits and host sands also controls the spatial distribution of gas hydrates on the Cascadia margin.

PRELIMINARY SCIENTIFIC ASSESSMENT

The revised IODP U.S. Implementing Organization (USIO) riserless vessel schedule included an additional 15 operational days for the Expedition 311. The revised schedule included a port call in Astoria, Oregon, at the start and in Victoria, British Columbia, at the end of Expedition 311, with 37 total operational days on site. The original operations plan for Expedition 311 was based on 22 days on site and included a visit of five sites in total.

The main change to accommodate the additional 15 days was in expanding the operations plan to a three hole approach at each site. Hole A was dedicated to LWD and MWD operations. Hole B was a continuous coring hole for APC/XCB coring, with additional temperature measurements using the APCT/DVTPP tools as well as monitoring of pressure/temperature conditions of each core by deploying the advanced piston corer methane (APCM) tool. We also planned to deploy the PCS at three intervals in Hole B. These three PCS cores were to be used for degassing experiments only. Because core recovery in gas hydrate-bearing sediments is challenging, we took advantage of the third hole to spot-core in missed intervals from Hole B. Hole C was mainly dedicated toward pressure coring, with a total of six deployments alternating the PCS, HPC, and FPC systems. Hole C was then wireline logged after coring operations were completed. The additional time and associated flexibility in the program allowed us to work through weather problems and operational complexities.

For this expedition, we carried out LWD/MWD operations prior to coring each site. The LWD/MWD logging program surpassed all expectations. The newly developed LWD/MWD safety protocol provided an effective means to deal with concerns associated with shallow gas hazards. The logging data guided special tool deployments (PCS, FPC, and HRC) in addition to providing high-quality downhole measurements, which were used to identify and characterize gas hydrate occurrences.

Uncertainty remains in the detailed velocity profile at each site along the transect to resolve critical issues related to the depth of the BSR. Additional VSPs and walkaway VSPs will be needed to address these issues.

Weather conditions certainly impacted operations throughout the duration of the expedition. A combination of adverse weather and severe sea state resulted in schedule delays adding up to >4 days. In addition, the deteriorating conditions resulted in reduced core recovery and quality. Both temperature tool and pressure core deployments were also adversely affected by general deterioration of weather conditions throughout the expedition.

An unforeseen problem arose with the recognition of the fact that many of the gas hydrate accumulations encountered during the expedition occurred in sand-rich sediment sections. Recovery with standard XCB and specialized pressure core systems were both negatively affected by the presence of sands.

The IODP and Transocean staff members need to be complimented in their ability to execute the complex operations and to demonstrate outstanding flexibility to adapt to changing conditions in requirements.

The original proposal 553-Full2 Cascadia margin gas hydrates included deployment of long-term monitoring devices such as ACORK instruments and fiber-optic DTS cables as well as coring and logging of a reference site in the deep Cascadia Basin west of the deformation front to characterize the incoming, undeformed, and gas hydrate-free sediment column. The long-term monitoring aspect of the proposal is closely linked to the NEPTUNE cable observatory and is planned to connect the ACORK and DTS instruments by a node near Site 889. The NEPTUNE program is laying out a cable route within the next 2 years that exactly follows the transect established during this expedition.

Although Expedition 311 addressed many key issues of gas hydrate occurrence on this margin, several important questions are still unanswered and need to be addressed

during Phase II drilling. One of the most important aspects in gas hydrate research is establishing the extent of the gas hydrate stability zone. Capturing the temperature field with coarsely spaced single-point measurements and regression analyses to establish linear temperature gradients was shown to be not only challenging due to the weather conditions but also inappropriate in a region dominated by advective fluid/gas flow. It is therefore crucial to deploy a system like DTS that allows measurements on a higher vertical resolution to map the gas hydrate stability zone in much more detail, especially around the base of the stability zone. This system will provide the answers to critical questions (e.g., of the potential occurrence of Structure II gas hydrate and temporal changes in the system).

Expedition 311 further showed evidence for strong lateral heterogeneity between adjacent bore holes, especially within sites where the gas hydrate occurs mainly within the accreted sediment complex (e.g., Site U1327). The occurrence of gas hydrate appeared to be driven by local variation of methane solubility as well as proximity of suitable host sediments (coarser-grained turbidite sands). A key question here is permeability in the sediments and how fluid and gas can migrate through the system. The permeability may be on a sediment grain scale, on a centimeter scale (the scaly fabric observed in previous ODP clastic accretionary prism cores), or in closely spaced faults. A key experiment to measure such permeability on these various scales is given by the proposed dual ACORK deployment. Two boreholes spaced several tens of meters apart equipped with the ACORK instruments are required to carry out a cross-hole hydrogeologic fluid-flow experiment. The two boreholes additionally offer the opportunity to deploy sensors for geophysical measurements (e.g., multicomponent geophones for active and passive seismic experiments, resistivity receivers for CSEM experiments, and strainmeters to measure long-term deformation [to be linked with the newly installed strainmeters at the Pacific Geoscience Centre in Sydney and other sites along the margin and as part of the proposed NEPTUNE cable observatory]).

REFERENCES

- Aldred, W., Cook, J., Carpenter, B., Hutchinson, M., Lovell, J., Rezmer-Cooper, I., and Leder, P.C., 1998. Using downhole annular pressure measurements to improve drilling performance. *Oilfield Rev.*, Winter:40–55.
- Chapman, N.R., Gettrust, J., Walia, R., Hannay, D., Spence, G.D., Wood, W., and Hyndman, R.D., 2002. High-resolution, deep-towed, multichannel seismic survey of deep-sea gas hydrates off western Canada. *Geophysics*, 67(4):1038–1047. doi:10.1190/1.1500364
- Collett, T.S., 2000. Quantitative well-log analysis of in-situ natural gas hydrates [Ph.D. dissert.]. Colorado School of Mines, Golden.
- Dallimore, S.R., and Collett, T.S. (Eds.), 2005. *Scientific Results from the Mallik 2002 Gas Hydrate Production Research Well Program, Mackenzie Delta, Northwest Territories, Canada*. Bull.—Geol. Surv. Canada, 585.
- Dallimore, S.R., Collett, T.S., and Uchida, T., 1999. Overview of science program, JAPEX/JNOC/GSC Mallik 2L-38 gas hydrate research well. In Dallimore, S.R., Uchida, T., and Collett, T.S. (Eds.), *Scientific Results from JAPEX/JNOC/GSC Mallik 2L-38 Gas Hydrate Research Well, MacKenzie Delta, Northwest Territories Canada*. Bull.—Geol. Surv. Canada.
- Davis, E.E., and Hyndman, R.D., 1989. Accretion and recent deformation of sediments along the northern Cascadia subduction zone. *Geol. Soc. Am. Bull.*, 101(11):1465–1480. doi:10.1130/0016-7606(1989)101<1465:AARDOS>2.3.CO;2
- Dickens, G.R., Paull, C.K., Wallace, P., and the ODP Leg 164 Scientific Party, 1997. Direct measurement of in situ methane quantities in a large gas-hydrate reservoir. *Nature (London, U. K.)*, 385:427–428. doi:10.1038/385426a0
- Dickens, G.R., Wallace, P.J., Paull, C.K., and Borowski, W.S., 2000. Detection of methane gas hydrate in the pressure core sampler (PCS): volume-pressure-time relations during controlled degassing experiments. In Paull, C.K., Matsumoto, R., Wallace, P.J., and Dillon, W.P. (Eds.), *Proc. ODP, Sci. Results*, 164: College Station, TX (Ocean Drilling Program), 113–126. [HTML]
- Fehn, U., Lu, Z., and Tomaru, H., in press. Data report: ¹²⁹I/I ratios and halogen concentrations in pore water of the Hydrate Ridge and their relevance for the origin of gas hydrates: a progress report. In Tréhu, A.M., Bohrmann, G., Torres, M.E., and Colwell, F.S. (Eds.), *Proc. ODP, Sci. Results*, 204 [Online]. Available from World Wide Web: http://www-odp.tamu.edu/publications/204_SR/107/107.htm.
- Ford, K.H., Naehr, T.H., Skilbeck, C.G., and the Leg 201 Scientific Party, 2003. The use of infrared thermal imaging to identify gas hydrate in sediment cores. In D'Hondt, S.L., Jørgensen, B.B., Miller, D.J., et al., *Proc. ODP, Init. Repts.*, 201, 1–20 [CD-ROM]. Available from: Ocean Drilling Program, Texas A&M University, College Station TX 77845-9547, USA. [HTML]
- Hunt, J.M., 1979. *Petroleum Geochemistry and Geology*: San Francisco (W.H. Freeman).
- Hyndman, R.D., 1995. The lithoprobe corridor across the Vancouver Island continental margin: the structural and tectonic consequences of subduction. *Can. J. Earth Sci.*, 32:1777–1802.
- Hyndman, R.D., and Davis, E.E., 1992. A mechanism for the formation of methane hydrate and seafloor bottom-simulating reflectors by vertical fluid expulsion. *J. Geophys. Res.*, 97:7025–7041.
- Hyndman, R.D., Moore, G.F., and Moran, K., 1993. Velocity, porosity, and pore-fluid loss from the Nankai subduction zone accretionary prism. In Hill, I.A., Taira, A., Firth, J.V., et al., *Proc. ODP, Sci. Results*, 131: College Station, TX (Ocean Drilling Program), 211–220.

- Hyndman, R.D., Moran, K., and Yuan, T., 1999. The concentration of deep sea gas hydrates from downhole resistivity logs and laboratory data. *Earth Planet. Sci. Lett.*, 172(1–2):167–177. [doi:10.1016/S0012-821X\(99\)00192-2](https://doi.org/10.1016/S0012-821X(99)00192-2)
- Hyndman, R.D., and Spence, G.D., 1992. A seismic study of methane hydrate marine bottom simulating reflectors. *J. Geophys. Res.*, 97:6683–6698.
- Hyndman, R.D., Spence, G.D., Chapman, N.R., Riedel, M., and Edwards, R.N., 2001. Geophysical studies of marine gas hydrate in northern Cascadia. In Paull, C.K., and Dillon, W.P. (Eds.), *Natural Gas Hydrates, Occurrence, Distribution and Detection*. Geophys. Monogr., 124:273–295.
- Hyndman, R.D., and Wang, K., 1993. Thermal constraints on the zone of major thrust earthquake failure: the Cascadia subduction zone. *J. Geophys. Res.*, 98:2039–2060.
- Jarrard, R.D., MacKay, M.E., Westbrook, G.K., and Sreaton, E.J., 1995. Log-based porosity of ODP sites on the Cascadia accretionary prism. In Carson, B., Westbrook, G.K., Musgrave, R.J., and Suess, E. (Eds.), *Proc. ODP, Sci. Results*, 146 (Pt 1): College Station, TX (Ocean Drilling Program), 313–335.
- Kastner, M., Kvenvolden, K.A., Whiticar, M.J., Camerlenghi, A., and Lorenson, T.D., 1995. Relation between pore fluid chemistry and gas hydrate associated with bottom-simulating reflectors at the Cascadia margin, Sites 889 and 892, *Proc. ODP, Sci. Results*, 146 (Pt 1): College Station, TX (Ocean Drilling Program), 175–187.
- Kvenvolden, K.A., 1988. Methane hydrate—a major reservoir of carbon in the shallow geosphere? *Chem. Geol.*, 71:41–51. [doi:10.1016/0009-2541\(88\)90104-0](https://doi.org/10.1016/0009-2541(88)90104-0)
- Lee, M.W., Hutchinson, D.R., Dillon, W.P., Miller, J.J., Agena, W.F., and Swift, B.A., 1993. Method of estimating the amount of in situ gas hydrates in deep marine sediments. *Mar. Pet. Geol.*, 10(5):493–506. [doi:10.1016/0264-8172\(93\)90050-3](https://doi.org/10.1016/0264-8172(93)90050-3)
- Lorenson, T.D., and the Leg 164 Shipboard Scientific Party, 2000. Graphic summary of gas hydrate occurrence by proxy measurements across the Blake Ridge, Sites 994, 995, and 997. In Paull, C.K., Matsumoto, R., Wallace, P.J., and Dillon, W.P. (Eds.), *Proc. ODP, Sci. Results*, 164: College Station, TX (Ocean Drilling Program), 247–249. [\[HTML\]](#)
- MacKay, M.E., Jarrard, R.D., Westbrook, G.K., and Hyndman, R.D., 1994. Origin of bottom-simulating reflectors: geophysical evidence from the Cascadia accretionary prism. *Geology*, 22(5):459–462. [doi:10.1130/0091-7613\(1994\)022<0459:OBSRG>2.3.CO;2](https://doi.org/10.1130/0091-7613(1994)022<0459:OBSRG>2.3.CO;2)
- Milkov, A.V., Dickens, G.R., Claypool, G.E., Lee, Y.-J., Borowski, W.S., Torres, M.E., Xu, W., Tomaru, H., Tréhu, A.M., and Schultheiss, P., 2004. Co-existence of gas hydrate, free gas, and brine within the regional gas hydrate stability zone at Hydrate Ridge (Oregon margin): evidence from prolonged degassing of a pressurized core. *Earth Planet. Sci. Lett.*, 222:829–843. [doi:10.1016/j.epsl.2004.03.028](https://doi.org/10.1016/j.epsl.2004.03.028)
- Paull, C.K., and Matsumoto, R., 2000. Leg 164 overview. In Paull, C.K., Matsumoto, R., Wallace, P.J., and Dillon, W.P. (Eds.), *Proc. ODP, Sci. Results*, 164: College Station, TX (Ocean Drilling Program), 3–10. [\[HTML\]](#)
- Paull, C.K., and Ussler, W., III, 1997. Are low salinity anomalies below BSRs a consequence of interstitial gas bubble barriers? *Eos, Trans. Am. Geophys. Union*, 78:F339.
- Pohlmann, J.W., Canuel, E.A., Chapman, N.R., Spence, G.D., Whiticar, M.J., and Coffin, R.B., 2005. The origin of thermogenic gas hydrates on the northern Cascadia margin as inferred from isotopic ($^{13}\text{C}/^{12}\text{C}$ and D/H) and molecular composition of hydrate and vent gas. *Org. Geochem.*, 36:703–716. [doi:10.1016/j.orggeochem.2005.01.011](https://doi.org/10.1016/j.orggeochem.2005.01.011)
- Shipley, T.H., Houston, M.H., Buffler, R.T., Shaub, F.J., McMillen, K.J., Ladd, J.W., and Worzel, J.L., 1979. Seismic evidence for widespread possible gas hydrate horizons on continental slopes and rises. *AAPG Bull.*, 63:2204–2213.
- Riddihough, R.P., 1984. Recent movements of the Juan de Fuca plate system. *J. Geophys. Res.*, 89:6980–6994.

- Riedel, M., 2001. 3D seismic investigations of northern Cascadia marine gas hydrates [Ph.D. dissert.]. Univ. Victoria, Canada.
- Riedel, M., Collett, T.S., and Hyndman, R.D. 2005. Gas hydrate concentration estimates from chlorinity, electrical resistivity, and seismic velocity. *Geol. Surv. Can. Open-File Rep.*, 4934.
- Riedel, M., Novosel, I., Spence, G.D., Hyndman, R.D., Chapman, N.R., and Lewis, T., in press. Geophysical and geochemical signatures associated with gas hydrate-related venting in the northern Cascadia margin. *Geol. Soc. Am. Bull.*, 118(1–2).
- Riedel, M., Spence, G.D., Chapman, N.R., and Hyndman, R.D., 2002. Seismic investigations of a vent field associated with gas hydrates, offshore Vancouver Island. *J. Geophys. Res.*, 107(B9). doi:10.1029/2001JB000269
- Solem, R.C., Spence, G.D., Vukajlovich, D., Hyndman, R.D., Riedel, M., Novosel, I., and Kastner, M., 2002. Methane advection and gas hydrate formation within an active vent field offshore Vancouver Island. *Fourth Int. Conf. Gas Hydrates: Yokohama, Japan, 19023*.
- Spence, G.D., Hyndman, R.D., Chapman, N.R., Riedel, M., Edwards, N., and Yuan, J., 2000. Cascadia margin, northeast Pacific Ocean: hydrate distribution from geophysical investigations. In Max, M.D. (Ed.), *Natural Gas Hydrate in Ocean and Permafrost Environments*: New York (Kluwer Academic Publisher).
- Taylor, A.E., Dallimore, S.D., Hyndman, R.D., and Wright, F., 2002. Comparing the sensitivity of terrestrial and marine gas hydrates to climate warming at the end of the last ice age. *Fourth Int. Conf. Gas Hydrates: Yokohama, Japan, 19023*.
- Teichert, B.M.A., Torres, M.E., Bohrmann, G., and Eisenhauer, A., 2005. Fluid sources, fluid pathways and diagenetic reactions across an accretionary prism revealed by Sr and B geochemistry. *Earth Planet. Sci. Lett.*, 239:106–121. doi:10.1016/j.epsl.2005.08.002
- Torres, M.E., Teichert, B.M.A., Tréhu, A.M., Borowski, W.S., and Tomaru, H., 2004. Relationship of pore water freshening to accretionary processes in the Cascadia margin: fluid sources and gas hydrate abundance. *Geophys. Res. Lett.*, 31:L22305. doi:10.1029/2004GL021219
- Torres, M.E., Wallmann, K., Tréhu, A.M., Bohrmann, G., Borowski, W.S., and Tomaru, H., 2004. Gas hydrate growth, methane transport, and chloride enrichment at the southern summit of Hydrate Ridge, Cascadia margin off Oregon. *Earth Planet. Sci. Lett.*, 226:225–241. doi:10.1016/j.epsl.2004.07.029
- Tréhu, A.M., Bohrmann, G., Rack, F.R., Collett, T.S., Goldberg, D.S., Long, P.E., Milkov, A.V., Riedel, M., Schultheiss, P., Torres, M.E., Bangs, N.L., Barr, S.R., Borowski, W.S., Claypool, G.E., Delwiche, M.E., Dickens, G.R., Gracia, E., Guerin, G., Holland, M., Johnson, J.E., Lee, Y.-J., Liu, C.-S., Su, X., Teichert, B., Tomaru, H., Vanneste, M., Watanabe, M., and Weinberger, J.L., 2004. Three-dimensional distribution of gas hydrate beneath southern Hydrate Ridge: constraints from ODP Leg 204. *Earth Planet. Sci. Lett.*, 222(3–4):845–862. doi:10.1016/j.epsl.2004.03.035
- Tucholke, B.E., Bryan, G.M., and Ewing, J.I., 1977. Gas-hydrate horizons detected in seismic-profiler data from the western North Atlantic. *AAPG Bull.*, 61:698–707.
- Ussler, W., III, and Paull, C.K., 2001. Ion exclusion associated with marine gas hydrate deposits. In Paull, C.K., and Dillon, W.P. (Eds.), *Natural Gas Hydrates: Occurrence, Distribution, and Detection*. Geophys. Monogr., 124:41–51.
- von Huene, R., and Pecher, I.A., 1998. Vertical tectonics and the origins of BSRs along the Peru margin. *Earth Planet. Sci. Lett.*, 166(1–2):47–55. doi:10.1016/S0012-821X(98)00274-X
- Westbrook, G.K., Carson, B., Musgrave, R.J., et al., 1994. *Proc. ODP, Init. Repts.*, 146 (Pt. 1): College Station, TX (Ocean Drilling Program).

- Wood, W.T., Gettrust, J.F., Chapman, N.R., Spence, G.D., and Hyndman, R.D., 2002. Decreased stability of methane hydrates in marine sediments owing to phase-boundary roughness. *Nature (London, U. K.)*, 420:656–660. [doi:10.1038/nature01263](https://doi.org/10.1038/nature01263)
- Xu, W., 2002. Phase balance and dynamic equilibrium during formation and dissociation of methane gas hydrate. *Fourth Int. Conf. Gas Hydrates: Yokohama, Japan, 19023*:199–200.
- Xu, W., 2004. Modeling dynamic marine gas hydrate systems. *Am. Mineral.*, 89:1271–1279.
- Xu, W., and Ruppel, C., 1999. Predicting the occurrence, distribution, and evolution of methane gas hydrate in porous marine sediments. *J. Geophys. Res.*, 104(B3):5081–5096. [doi:10.1029/1998JB900092](https://doi.org/10.1029/1998JB900092)
- Yuan, J., and Edwards, R.N., 2001. The assessment of marine gas hydrates through electrical remove sounding: hydrate without a BSR? *Geophys. Res. Lett.*, 27(16):2397–2400. [doi:10.1029/2000GL011585](https://doi.org/10.1029/2000GL011585)
- Yuan, T., Hyndman, R.D., Spence, G.D., and Desmons, B., 1996. Seismic velocity increase and deep-sea gas hydrate concentration above a bottom-simulating reflector on the northern Cascadia continental slope. *J. Geophys. Res.*, 101(B6):13655–13671. [doi:10.1029/96JB00102](https://doi.org/10.1029/96JB00102)
- Yuan, T., Spence, G.D., and Hyndman, R.D., 1994. Seismic velocities and inferred porosities in the accretionary wedge sediments at the Cascadia margin. *J. Geophys. Res.*, 99(B3):4413–4427. [doi:10.1029/93JB03203](https://doi.org/10.1029/93JB03203)
- Yuan, T., Spence, G.D., Hyndman, R.D., Minshull, T.A., and Singh, S.C., 1999. Seismic velocity studies of a gas hydrate bottom-simulating reflector on the northern Cascadia continental margin: amplitude modeling and full waveform inversion. *J. Geophys. Res.*, 104(B1):1179–1192. [doi:10.1029/1998JB900020](https://doi.org/10.1029/1998JB900020)
- Zühlsdorff, L., and Spiess, V., 2004. Three-dimensional seismic characterization of a venting site reveals compelling indications of natural hydraulic fracturing. *Geology*, 32(2):101–104. [doi:10.1130/G19993.1](https://doi.org/10.1130/G19993.1)

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Table T1. Expedition 311 operations summary.

Hole	Latitude	Longitude	Water depth (mbsl)	Cores (N)	Cored (m)	Recovered (m)	Recovery (%)	Drilled (m)	Penetration (m)
U1325A	48°38.691' N	126°58.991' W	2201.1	0	0	0.00		350.0	350.0
U1325B	48°38.694' N	126°58.999' W	2194.8	28	205.5	141.17	68.7	1.0	206.5
U1325C	48°38.701' N	126°59.007' W	2194.8	15	115.5	62.49	54.1	188.8	304.3
U1325D	48°38.701' N	126°59.007' W	2194.6	1	4.7	4.69	99.8	0.0	4.7
Site U1325 totals:				44	325.7	208.35	64.0	539.8	865.5
U1326A	48°37.635' N	127°3.029' W	1828.1	0	0	0.00		300.0	300.0
U1326B	48°37.638' N	127°3.023' W	1828.4	1	1.5	1.55	103.3	0.0	1.5
U1326C	48°37.638' N	127°3.023' W	1828.0	13	85.7	54.39	63.0	1.0	86.7
U1326D	48°37.628' N	127°3.043' W	1827.9	20	192.6	121.96	63.3	107.4	300.0
Site U1326 totals:				34	279.8	177.9	229.6	408.4	688.2
U1327A	48°41.887' N	126°51.921' W	1322.1	0	0	0.00		300.0	300.0
U1327B	48°41.889' N	126°51.914' W	1306.1	1	9.5	9.79	103.1	0.0	9.5
U1327C	48°41.889' N	126°51.914' W	1304.5	35	297	262.17	88.3	3.0	300.0
U1327D	48°41.895' N	126°51.906' W	1303.7	17	75.9	53.99	71.1	224.1	300.0
U1327E	48°41.900' N	126°51.896' W	1303.4	4	12.5	12.37	99.0	287.5	300.0
Site U1327 totals:				57	394.9	338.32	85.7	814.6	1209.5
U1328A	48°40.072' N	126°51.022' W	1268.1	0	0	0.00		300.0	300.0
U1328B	48°40.048' N	126°51.056' W	1267.8	10	54.5	39.65	72.8	2.0	56.5
U1328C	48°40.057' N	126°51.043' W	1267.7	27	242.5	195.64	80.7	57.5	300.0
U1328D	48°40.067' N	126°51.020' W	1266.7	3	15	10.03	66.9	0.0	15.0
U1328E	48°40.080' N	126°50.987' W	1264.7	13	49	13.33	27.2	186.0	235.0
Site U1328 totals:				53	361	258.65	71.6	545.5	906.5
U1329A	48°47.369' N	126°40.713' W	959.1	0	0	0.00		220.0	220.0
U1329B	48°47.375' N	126°40.720' W	953.0	1	9.5	10.02	105.5	0.0	9.5
U1329C	48°47.375' N	126°40.720' W	946.4	23	187.5	186.16	99.3	2.0	189.5
U1329D	48°47.362' N	126°40.716' W	946.4	1	9.5	1.03	10.8	201.0	210.5
U1329E	48°47.385' N	126°40.716' W	945.5	10	48	37.74	78.6	79.0	127.0
Site U1329 totals:				35	254.5	234.95	92.3	502.0	756.5
Expedition 311 totals:				223	1615.9	1218.17	75.4	2810.3	4426.2

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Table T2. Temperature tool deployments.

Core	Depth (mbsf)	Corrected temperature (°C)	Uncorrected temperature (°C)	Thermal conductivity (W/m·K)	Calibration correction (°C)	Temperature tool	Solution reliability
311-U1325B-							
4H	33.5	4.44	4.44	1.1	0	APC3	Excellent
6H	52.5	6.47	6.47	1.1	0	APC3	Excellent
8H	71.5	7.11	7.11	1.1	0	APC3	Excellent
19H–20H	150.2	12.66	12.66	1.1	0	DVTP	Excellent
311-U1325C-							
5H–6H	218.6	16.89	16.89	1.1	0	DVTP	Good
12H–13H	275.4	18.64	18.64	1.1	0	DVTP	Very good
311-U1326C-							
4H	30.4	4.3	4.3	1.1	0	APC3	Very good
311-U1326D-							
19X	252	NA†	NA	1.1	0	DVTP	Poor (heave)
21X	271.4	NA	NA	1.1	0	DVTP	Fair (heave)
BOH	300	NA	NA	1.1	0	DVTP	Fair (heave)
311-U1327C-							
3H	25.1	NA	NA	NA	-0.97	APCT16	Poor
5H	44.1	6.6	7.6	1.1	-0.97	APCT16	Good
7H	55.6	9.04	9.005	1.1	-0.97	APCT16	Good
7H	74.6	10.544	11.514	1.1	-0.97	APCT16	Excellent
19X–20X	161	NA	NA	NA	NA	DVTP	Instrument
25X–26X	209	15.1	15.1	1.1	0	DVTPP3	Excellent
34X–35X	295	NA	NA	NA	NA	DVTPP3	Poor
311-U1328B-							
8H	37.5	5.45	6.42	1.1	-0.97	APCT16	Good
10H	56.5	10*	10	1.1	0	APC3	Poor (heave)
10H–11H	56.5	ND	ND	ND	ND	DVTPP3	No data
311-U1328C-							
2H	75.5	7.57	8.54	1.1	-0.97	APCT16	Good
3H	85.0	9	9	1.1	0	APC3	Poor
6H	103.5	11	11	1.1	0	APC3	Poor (heave)
11X–12X	148.9	ND	ND	ND	ND	DVTPP3	No data
16X–17X	197.1	ND	ND	ND	ND	DVTPP3	No data
311-U1328E-							
10P–11Y	93.0	8.47	8.47	1.1	0	DVTP	Poor (heave)
11Y–12E	198.0	14.13	14.13	1.1	0	DVTP	Very good
311-U1329C-							
3H	27.1	NA	NA	NA	-0.97	APCT16	Poor
5H	46.1	6.6	7.6	1	-0.97	APCT16	Excellent
9H	76.6	9.04	9.005	1	0.03	APC3	Excellent
11H	95.6	10.544	11.514	1	-0.97	APCT16	Excellent
13H	114.6	11.79	11.759	1	0.03	APC3	Excellent
311-U1329E-							
4H	33.5	5.1	5.86/6.13	1	-0.97	APCT16	Fair
6H	64	7.76	8.69	1	-0.97	APCT16	Good
6H–7P	64	7.84	7.84	1	0	DVTP	Good
BOH	127.1	12.1	12.1	NA	0	DVTP	Excellent

Notes: NA = not applicable, ND = no data. BOH = bottom of hole. DVTP(P) = Davis-Villinger Temperature (Pressure) Probe, APCT = APC temperature tool, * = measurements severely affected by heave are estimated to nearest degree. † = tool did not stay in the seafloor long enough to obtain in situ temperature using CONEFIT. These data will be modeled postcruise in an attempt to place an upper bound on in situ temperature.

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Table T3. Pressure coring results.

Core	Date	Local time (h)	Top depth (mbsf)	Cored (m)	Recovered (m)	Recovery (%)	Pressurized core	Inferred S _H (%)
311-U1329C-								
7P	9/28/05	0255	55.6	1.0	1.00	100.0	Y	0.0
14P	9/28/05	1110	114.6	1.0	1.00	100.0	N	
23P	9/29/05	0210	188.5	1.0	1.00	100.0	Y	2.4
311-U1329E-								
5P	10/1/05	0630	33.5	1.0	0.00	0.0	N	
7P	10/1/05	1220	73.5	1.0	0.86	86.0	Y	0.7
8Y	10/1/05	1430	104.0	1.0	1.00	100.0	N	
9E	10/1/05	1625	114.5	1.0	1.00	100.0	Y	
10P	10/1/05	1940	125.0	1.0	0.95	95.0	Y	0.0
311-U1327C-								
6P	10/2/05	1435	44.1	1.0	1.00	100.0	N	
15P	10/3/05	0020	121.8	1.0	1.00	100.0	Y	0.2
24P	10/3/05	1715	197.3	1.0	1.00	100.0	Y	1.8
311-U1327D-								
3P	10/4/05	2320	83.0	1.0	0.00	0.0	N	
4E	10/6/05	0050	125.3	1.0	1.10	110.0	Y	
6Y	10/6/05	0340	132.0	1.0	0.84	84.0	N	
10P	10/6/05	1040	155.1	1.0	0.92	92.0	Y	8.0
12E	10/6/05	1540	170.5	1.0	1.14	114.0	Y	
13Y	10/6/05	2045	203.6	1.0	0.48	48.0	Y	
14E	10/7/05	0110	217.7	1.0	1.15	115.0	Y	
17P	10/7/05	1030	246.0	1.0	1.00	100.0	Y	0.9
311-U1327E-								
2P	10/9/05	1300	40.0	1.0	1.00	100.0	N	
3P	10/9/05	1550	80.0	1.0	0.95	95.0	Y	0.2
4E	10/9/05	2040	128.0	1.0	0.48	48.0	N	
311-U1328B-								
4P	10/11/05	1720	14.5	1.0	0.77	77.0	Y	15.0
7P	10/11/05	2105	26.0	1.0	0.78	78.0	Y	2.0
311-U1328C-								
5P	10/13/05	0510	92.0	1.0	0.96	96.0	Y	0.7
311-U1328D-								
3Y	10/15/05	2325	14.0	1.0	0.00	0.0	N	
311-U1328E-								
3E	10/16/05	0350	8.5	1.0	0.00	0.0	N	
5P	10/16/05	0600	15.1	1.0	0.00	0.0	N	
7Y	10/16/05	0830	25.7	1.0	0.87	87.0	N	
10P	10/16/05	1455	92.0	1.0	0.26	26.0	Y	22.0
11Y	10/16/05	2055	197.0	1.0	0.90	90.0	N	
12E	10/17/05	0130	215.5	1.0	0.00	0.0	N	
13P	10/17/05	0450	233.0	1.0	0.94	94.0	Y	50.0
311-U1325B-								
11P	10/18/05	1315	82.4	1.0	0.88	88.0	N	
17E	10/18/05	2145	129.9	1.0	0.00	0.0	N	
22Y	10/19/05	0725	169.4	1.0	0.17	17.0	N	
26E	10/19/05	1345	197.4	1.0	0.28	28.0	N	
28P	10/19/05	1925	206.1	0.4	0.00	0.0	N	

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Core	Date	Local time (h)	Top depth (mbsf)	Cored (m)	Recovered (m)	Recovery (%)	Pressurized core	Inferred S _H (%)
311-U1325C-								
5Y	10/21/05	0600	217.6	1	0.00	0.0	N	
10P	10/21/05	1750	256.1	1	0.96	96.0	Y	0.4
311-U1326C-								
11Y	10/24/05	0445	82.7	1	0.15	15.0	Y	
12P	10/24/05	0730	83.7	1	0.26	26.0	Y	0.4
13E	10/24/05	0905	85.7	1	0.00	0.0	N	

Figure F1. General location of drilling transect near previous ODP Sites 889/890. A bottom-simulating reflector is present on ~50% of the midcontinental slope (shaded area).

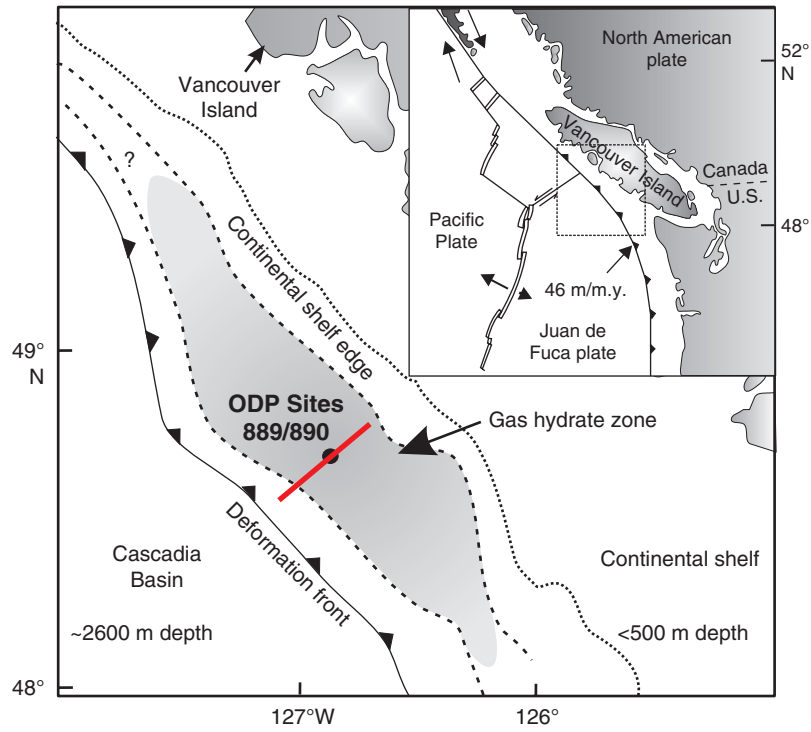


Figure F2. Marine gas hydrate cycle (after Hyndman and Davis, 1992). BSR = bottom-simulating reflector.

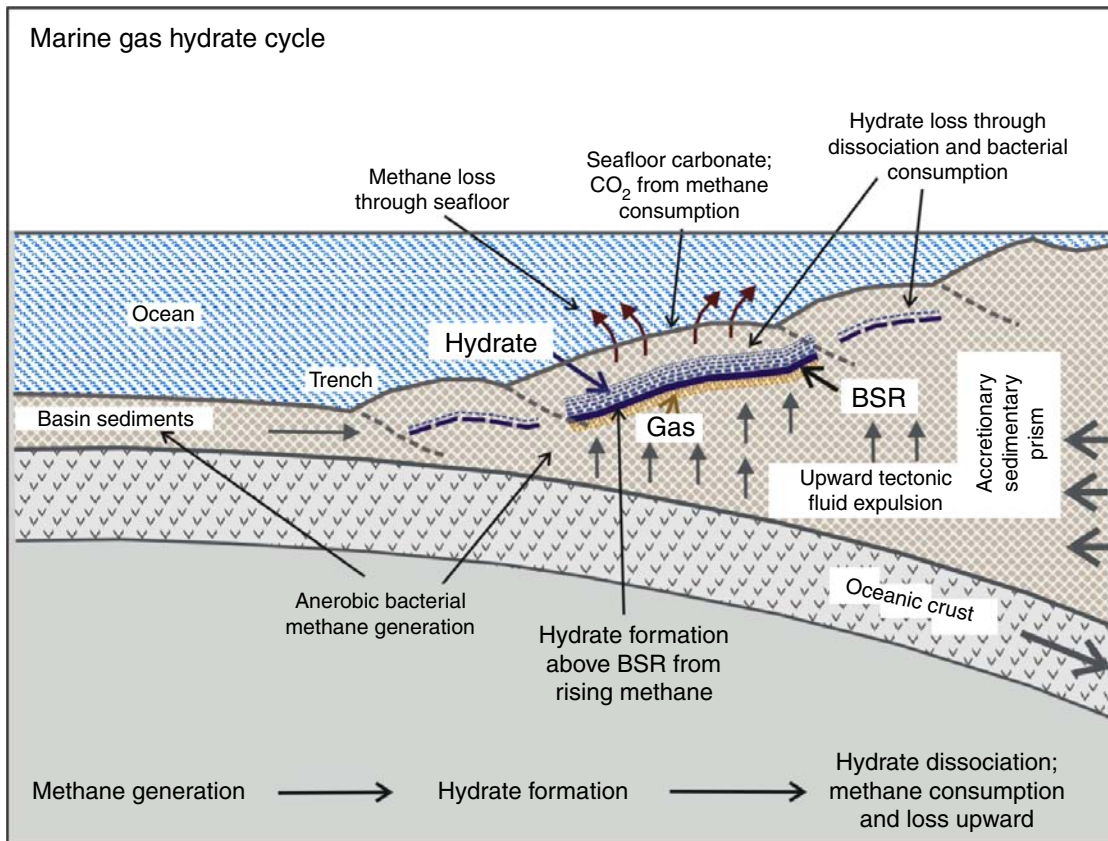


Figure F3. Overview of MCS Line 89-08 across the Northern Cascadia margin showing the transect with the four transect sites accomplished during Expedition 311 (Sites U1326, U1325, U1327, and U1328) as well as prospectus basin reference Site CAS-04B. The cold vent Site U1328 (prospectus Site CAS-06A) is ~2 nmi off this line toward the southeast.

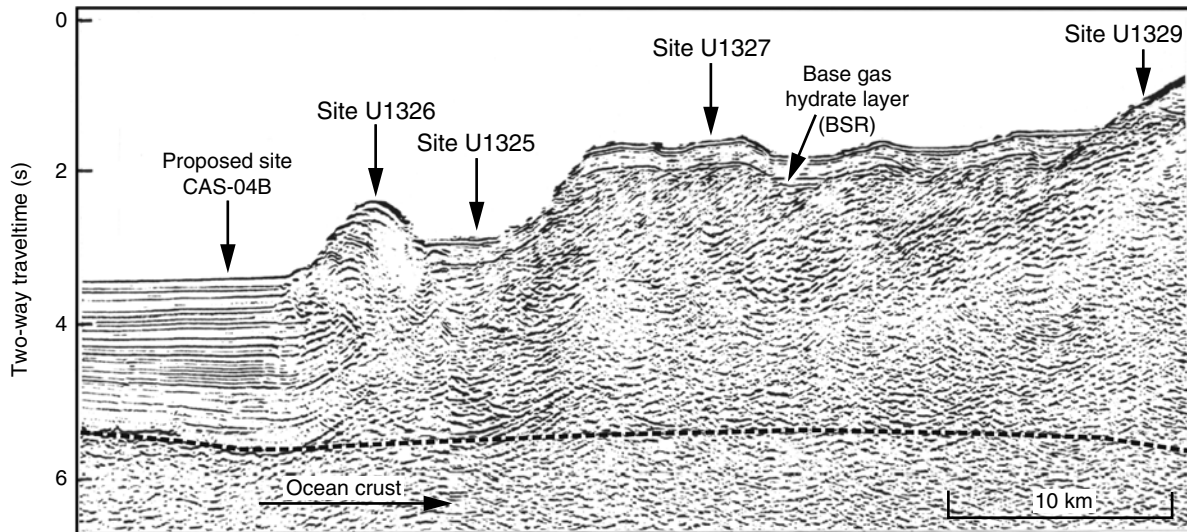


Figure F4. Multibeam bathymetry map along the transect across the accretionary prism offshore Vancouver Island (courtesy of D. Kelley, J. Delaney, and D. Glickson, University of Washington, and C. Barnes, C. Katnick, NEPTUNE Canada, University of Victoria; funded by the University of Washington and the W.M. Keck Foundation). MCS = multichannel seismic.

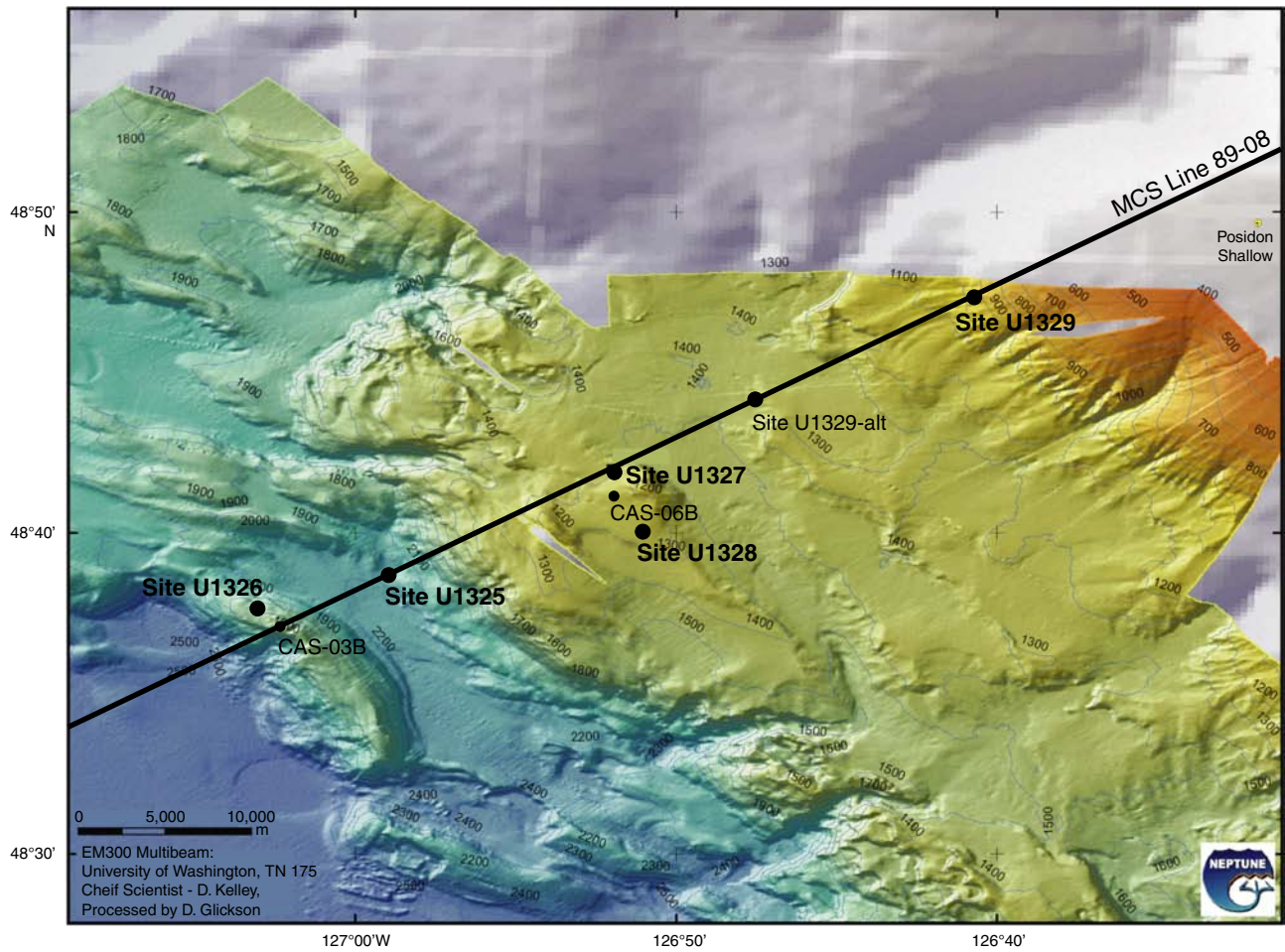


Figure F5. Multibeam bathymetry map over first accreted ridge documenting the collapse feature around Site U1326. Linear features, identified as faults from seismic data, cross the ridge in the central part (courtesy of D. Kelley, J. Delaney, and D. Glickson, University of Washington, and C. Barnes, C. Katnick, NEPTUNE Canada, University of Victoria; funded by the University of Washington and the W.M. Keck Foundation).

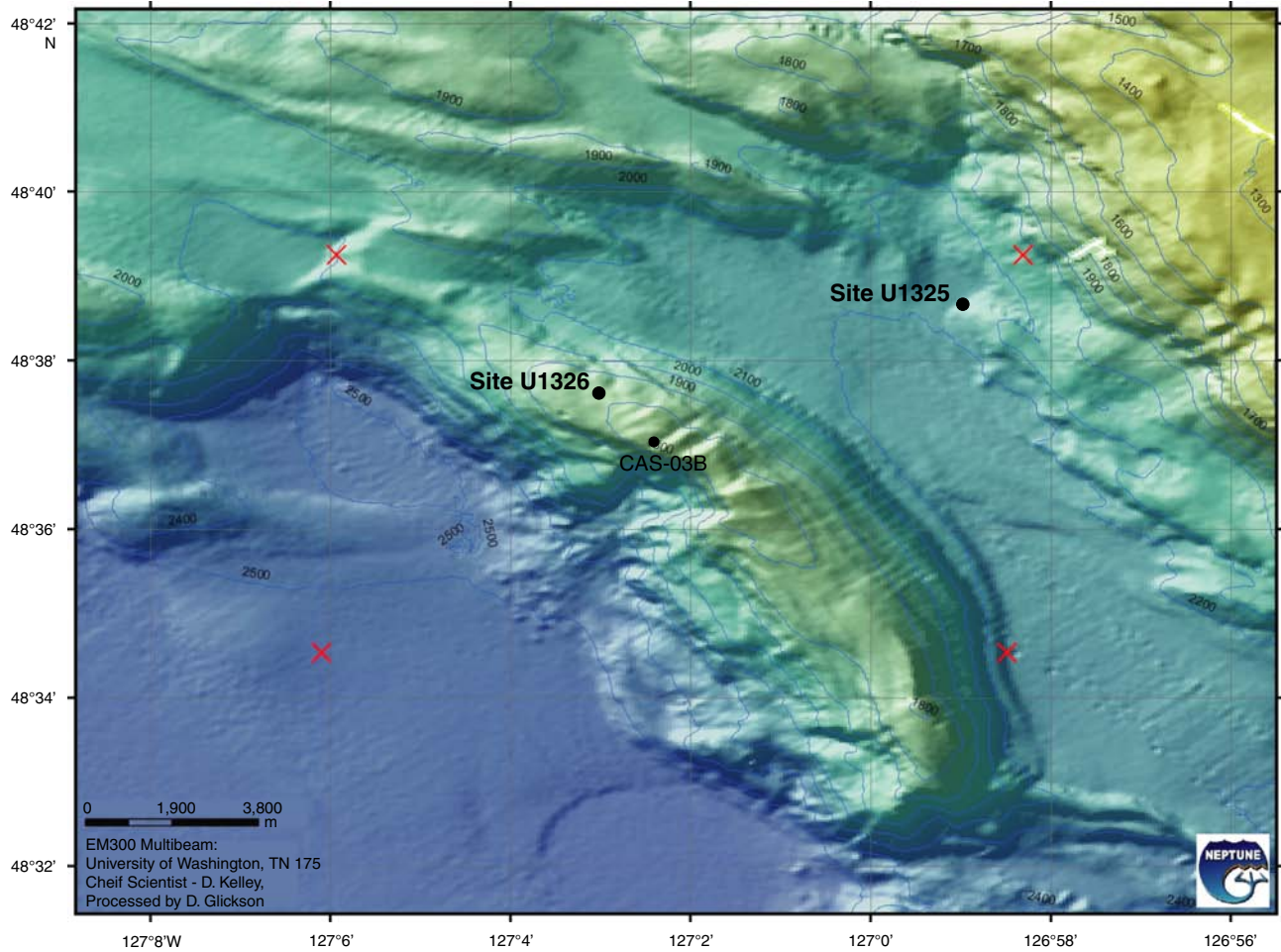


Figure F6. Summary of Expedition 311 measured borehole fluid pressure profiles with the best-fit linear trend subtracted. Annular pressure while drilling (APWD) data were the primary measurement used for gas monitoring during MWD/LWD drilling.

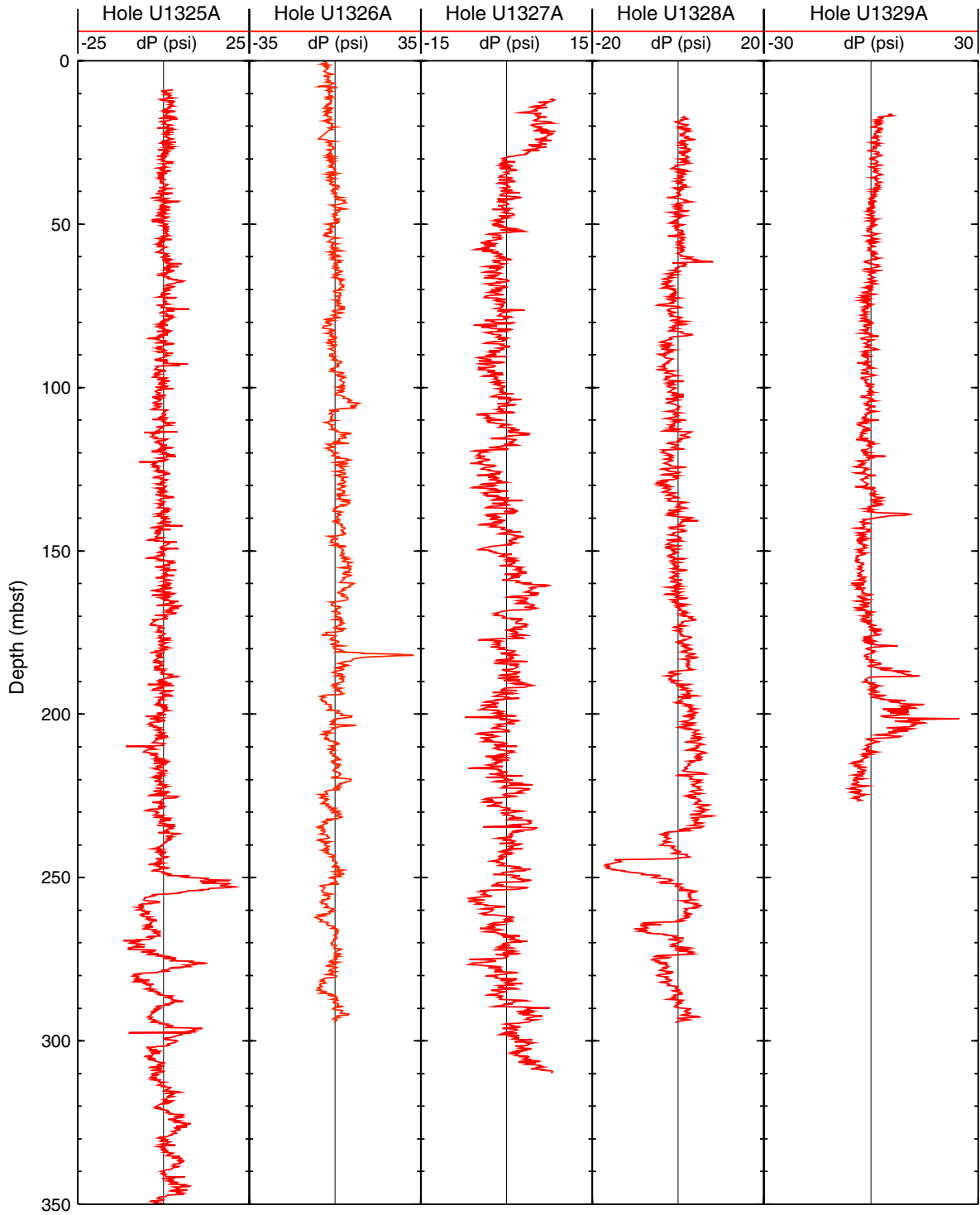


Figure F7. Section of MCS Line 89-08 over Site U1325. The data are of low seismic frequency content from 10 to 60 Hz. A buried ridge of accreted sediments can be identified between common depth points CDPs 1320 and 1440, covered by a thin interval of slope-basin-type sediments that form onlaps onto the ridge. Underneath the center of the ridge a bright spot is identified, potentially indicating the presence of free gas. A strong bottom-simulating reflector (BSR) is seen over much of the eastern section of the basin, whereas it disappears underneath the western part of the basin, filled with as much as 600 ms TWT thick slope-basin-type sediments. Site U1325 intersects an interval of 300 ms TWT thickness where sediments show strong eastward dip. Note the boundary between lithostratigraphic Subunits IA and IB is traceable for ~1.5 km around Site U1325. Several small-scale faults are visible, one of which was most likely penetrated at Site U1325 near the base of lithostratigraphic Unit II. Vertical exaggeration = 3.4.

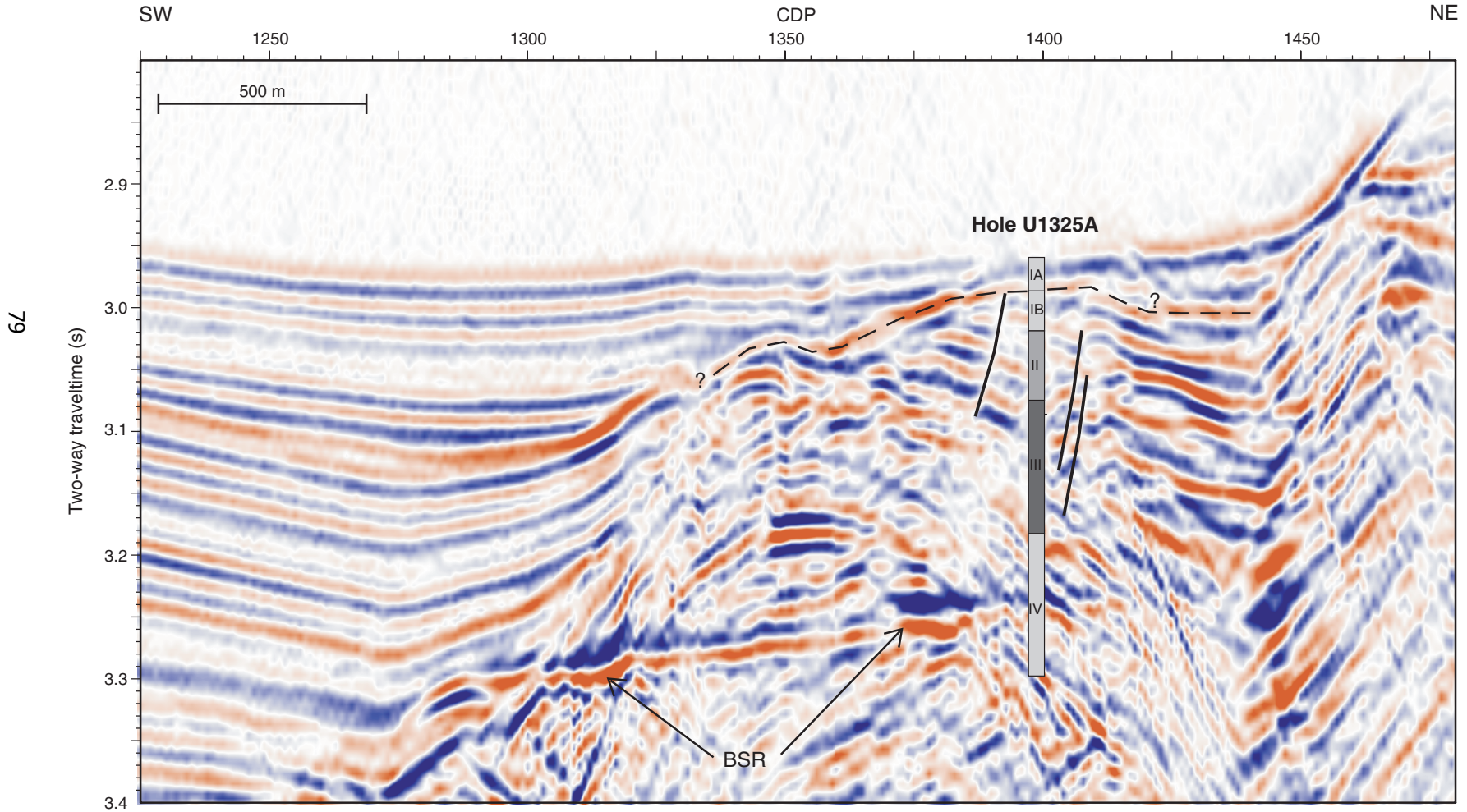
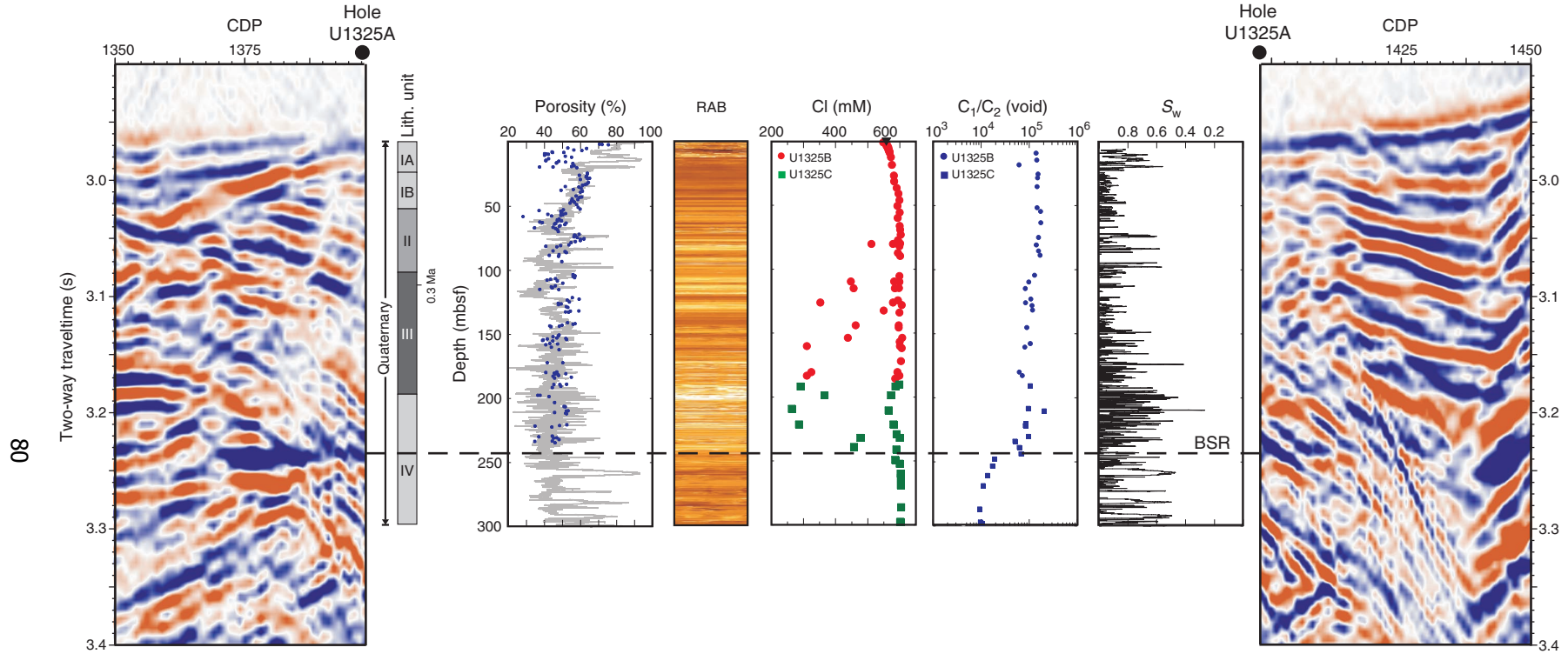


Figure F8. Comparison of seismic data from MCS Line 89-08 nearest to Site U1325 and main logging and core-derived measurements (porosity from LWD/MWD and MAD data, resistivity-at-the-bit [RAB] image from LWD/MWD deployment, chlorinity of interstitial water samples, C_1/C_2 ratio of the void gas, and pore water saturation [S_w] from LWD/MWD resistivity data). CDP = common depth point. BSR = bottom-simulating reflector.



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Figure F9. Comparison of seismic data from MCS Line PGC9902_ODP7 (twice the frequency content of MCS Line 89-08) nearest to Site U1325 and main logging- and core-derived measurements (porosity from LWD/MWD and MAD data, resistivity-at-the-bit [RAB] image from LWD/MWD deployment, chlorinity of interstitial water samples, C_1/C_2 ratio of the void gas, and pore water saturation [S_w] from LWD/MWD resistivity data). CDP = common depth point. BSR = bottom-simulating reflector.

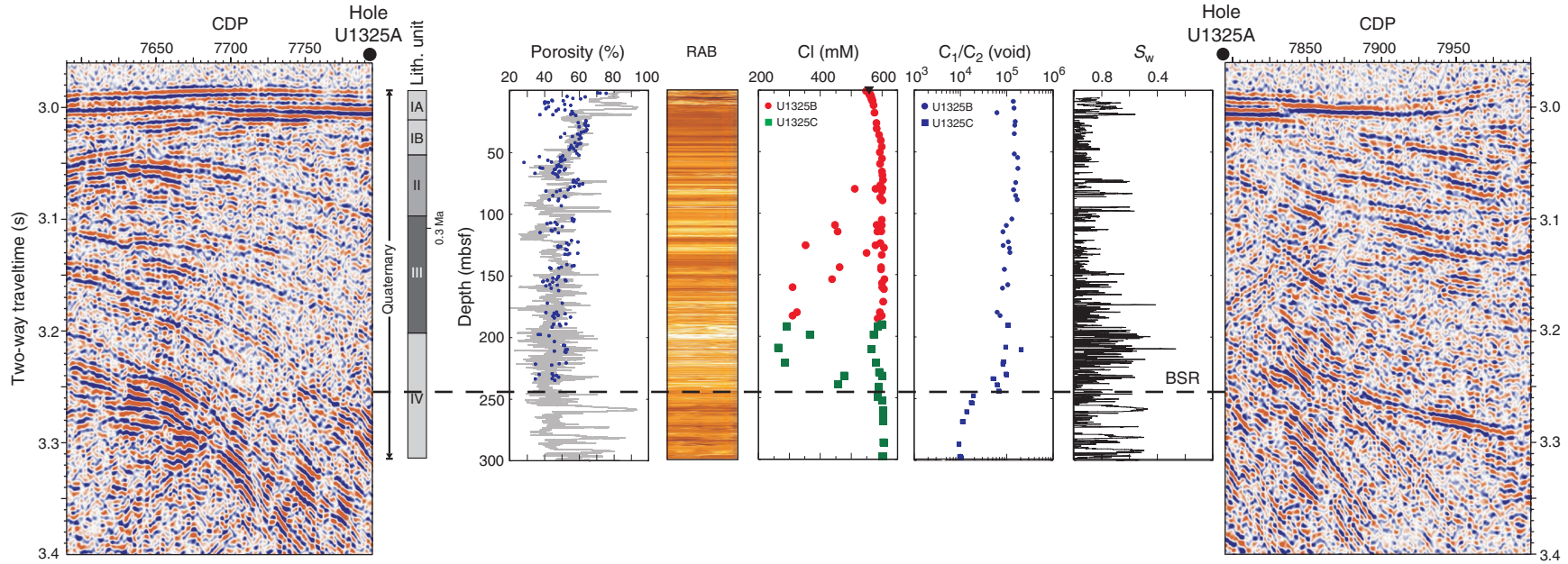


Figure F10. Section from MCS Line PGC9902_ODP7_MCS from the 1999 survey over Site U1326. The bottom-simulating reflector (BSR) is clearly visible. Also shown are deep faults related to apparent normal faulting at the surface. CDP = common depth point. Vertical exaggeration = 4.7.

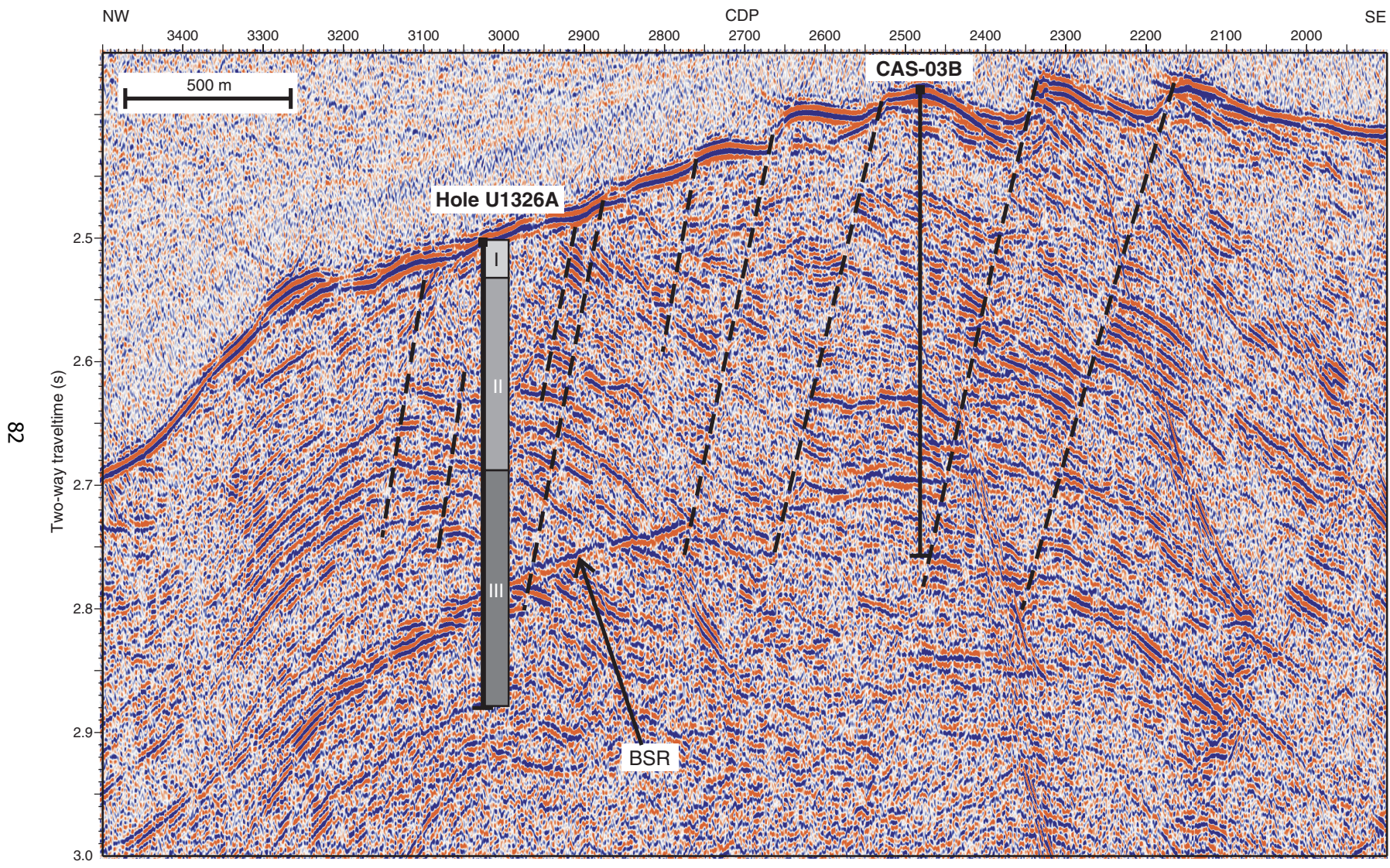


Figure F11. Section from SCS Line PGC0408_CAS03B_X03 crossing Site U1326. A bottom-simulating reflector (BSR) is visible, but almost no coherent reflectivity is identified. CDP = common depth point. Vertical exaggeration = 3.7.

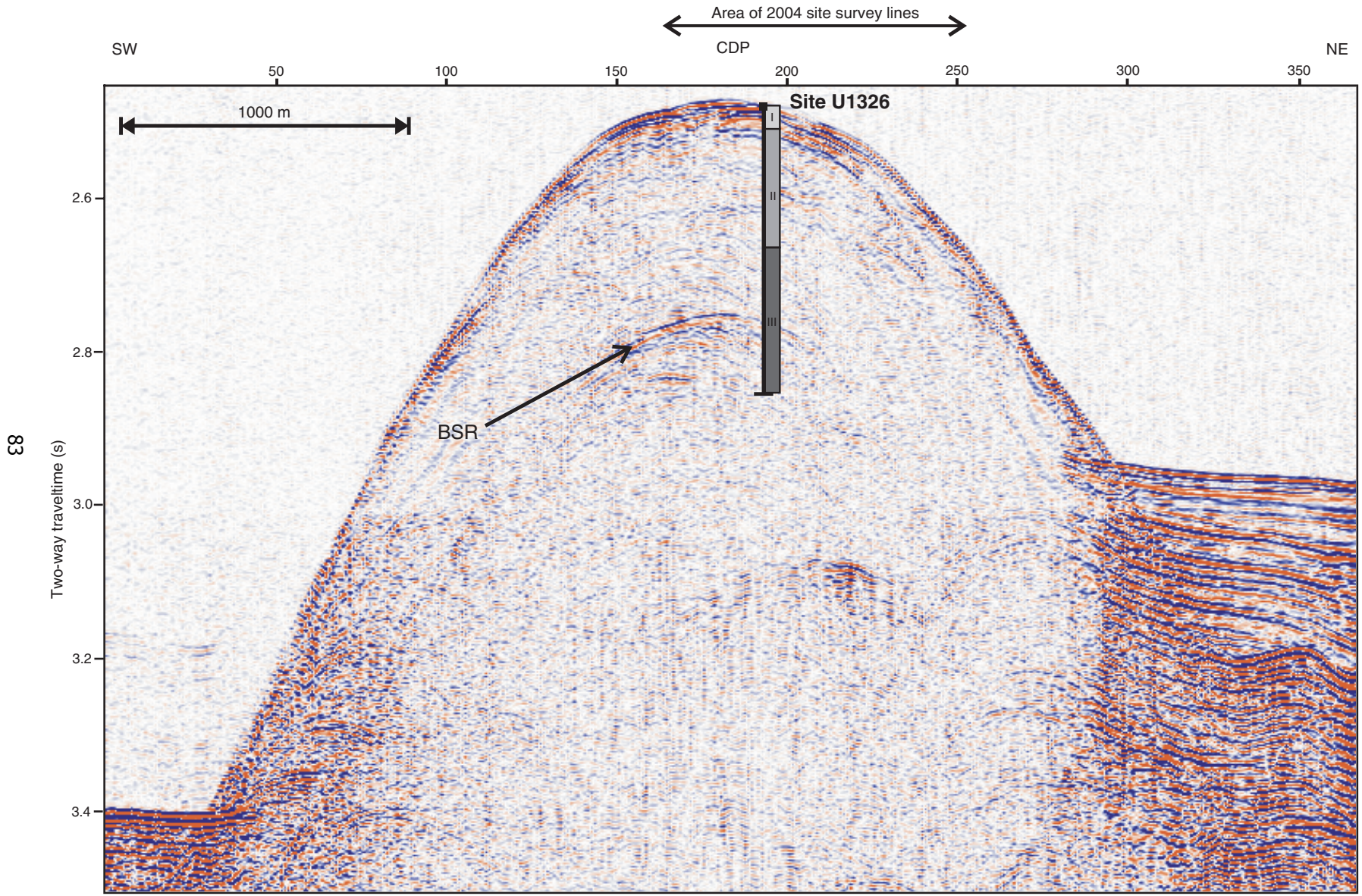


Figure F12. Section from MCS Line 89-08 over Site U1326 from Leg 146 presite survey showing in low-frequency data the strong bottom-simulating reflector (BSR) underneath the center of the ridge. The asymmetric deformation of the ridge is also seen by strong changes in coherent seismic reflectivity from southwest to northeast. CDP = common depth point. Vertical exaggeration = 4.4.

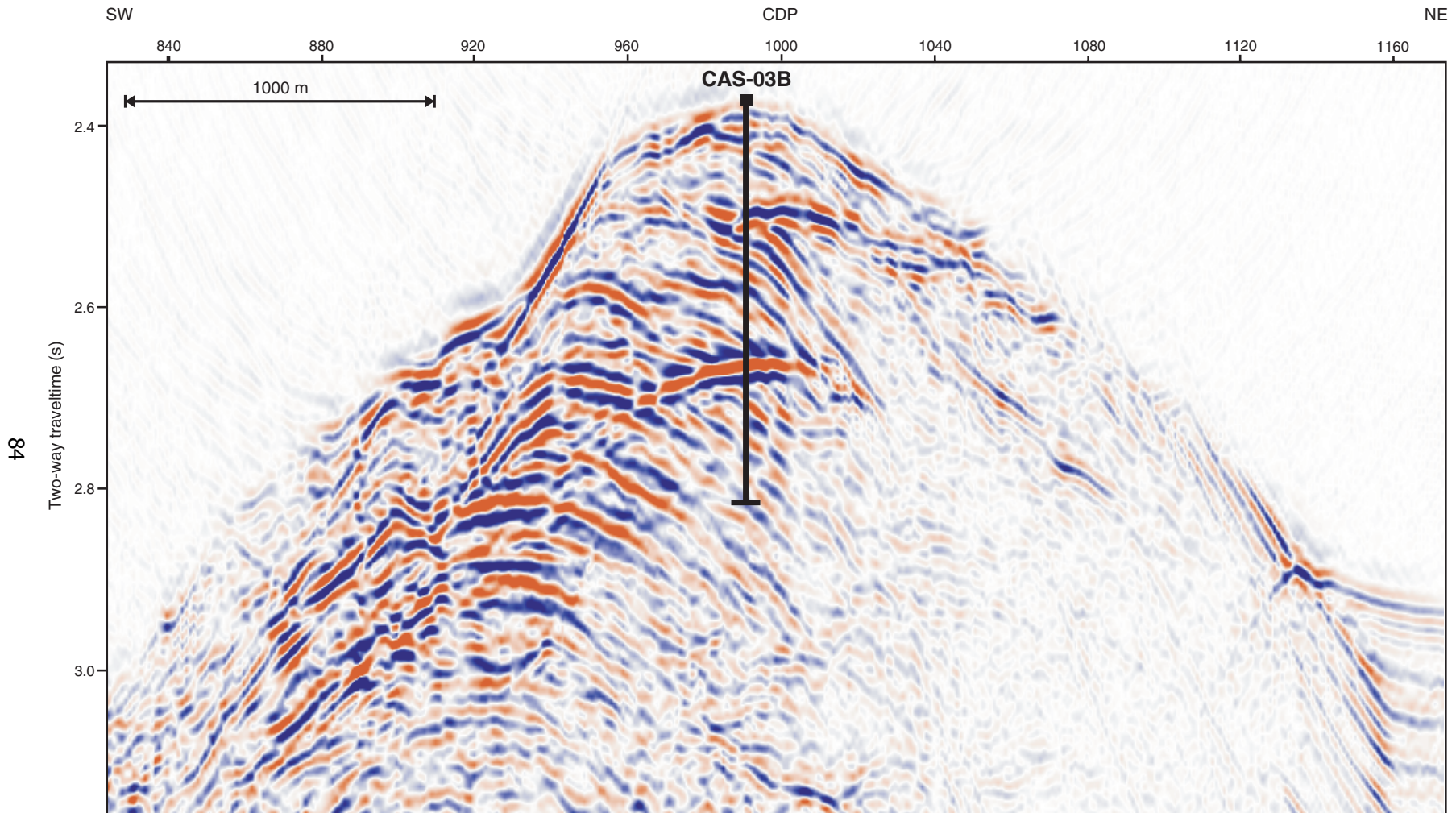


Figure F13. Compilation of seismic data from MCS Line PGC9902_ODP7_MCS and key data from Site U1326, including lithostratigraphic units (Holes U1326C and U1326D), porosity from LWD/MWD data (Hole U1326A) and MAD data (Holes U1326C and U1326D), RAB resistivity data from LWD/MWD (Hole U1326A), interstitial water chlorinity (Holes U1326C and U1326D), C_1/C_2 ratio (Holes U1326C and U1326D), and pore water saturation (S_w) from LWD data (Hole U1326A). The bright band of high electrical reflectivity as seen in the resistivity-at-the-bit (RAB) image (identified in Hole U1326A) does not correlate to a prominent seismic event. CDP = common depth point. GH = gas hydrate. PCS = pressure core sampler. BSR = bottom-simulating reflector.

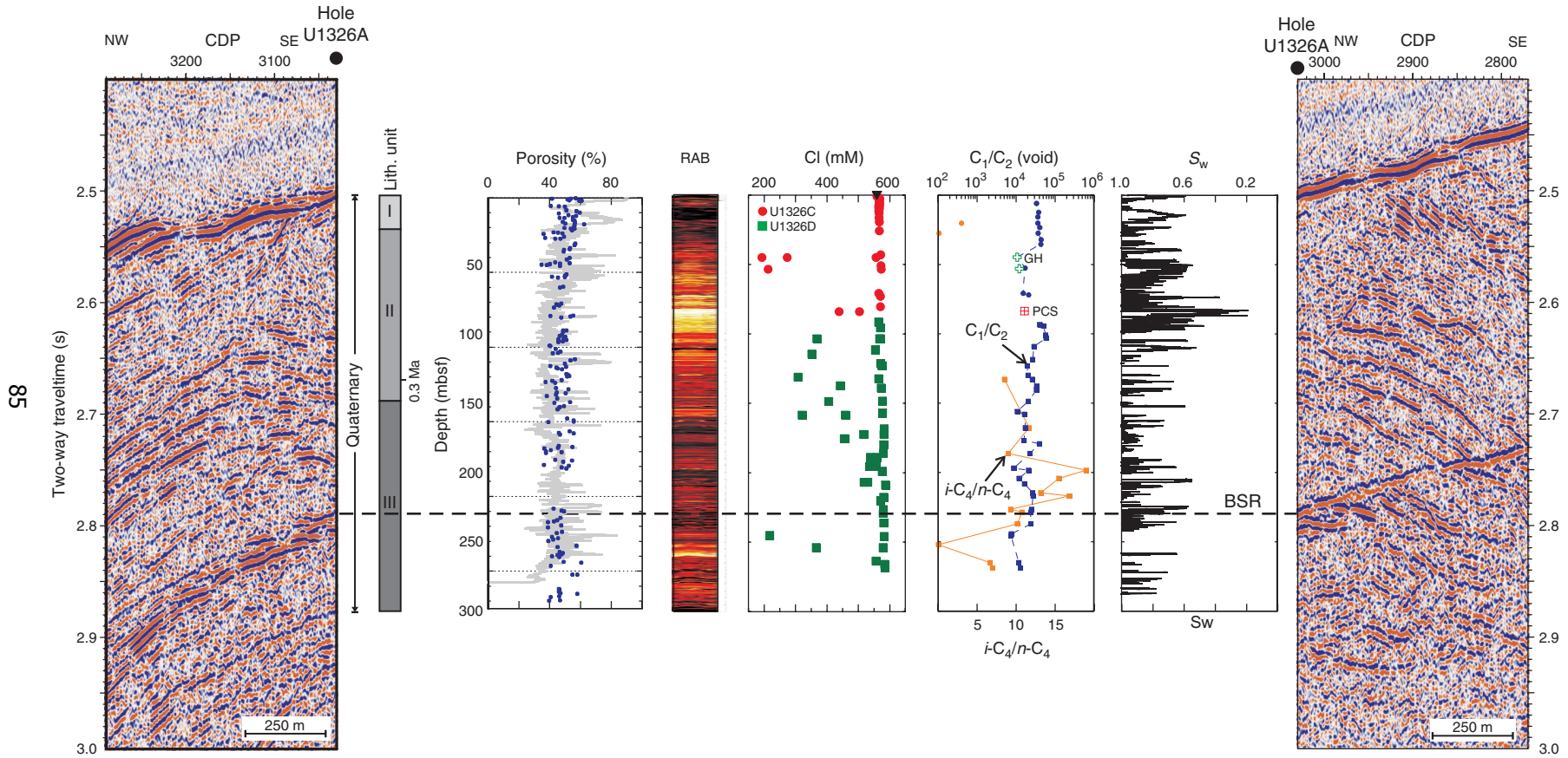


Figure F14. Bathymetric map of the seafloor near Site U1327 (courtesy by V. Spiess, *Sonne* cruise 111). The bathymetry is dominated by two topographic highs that rise to about 200 m above surrounding seafloor. MCS = multichannel seismic.

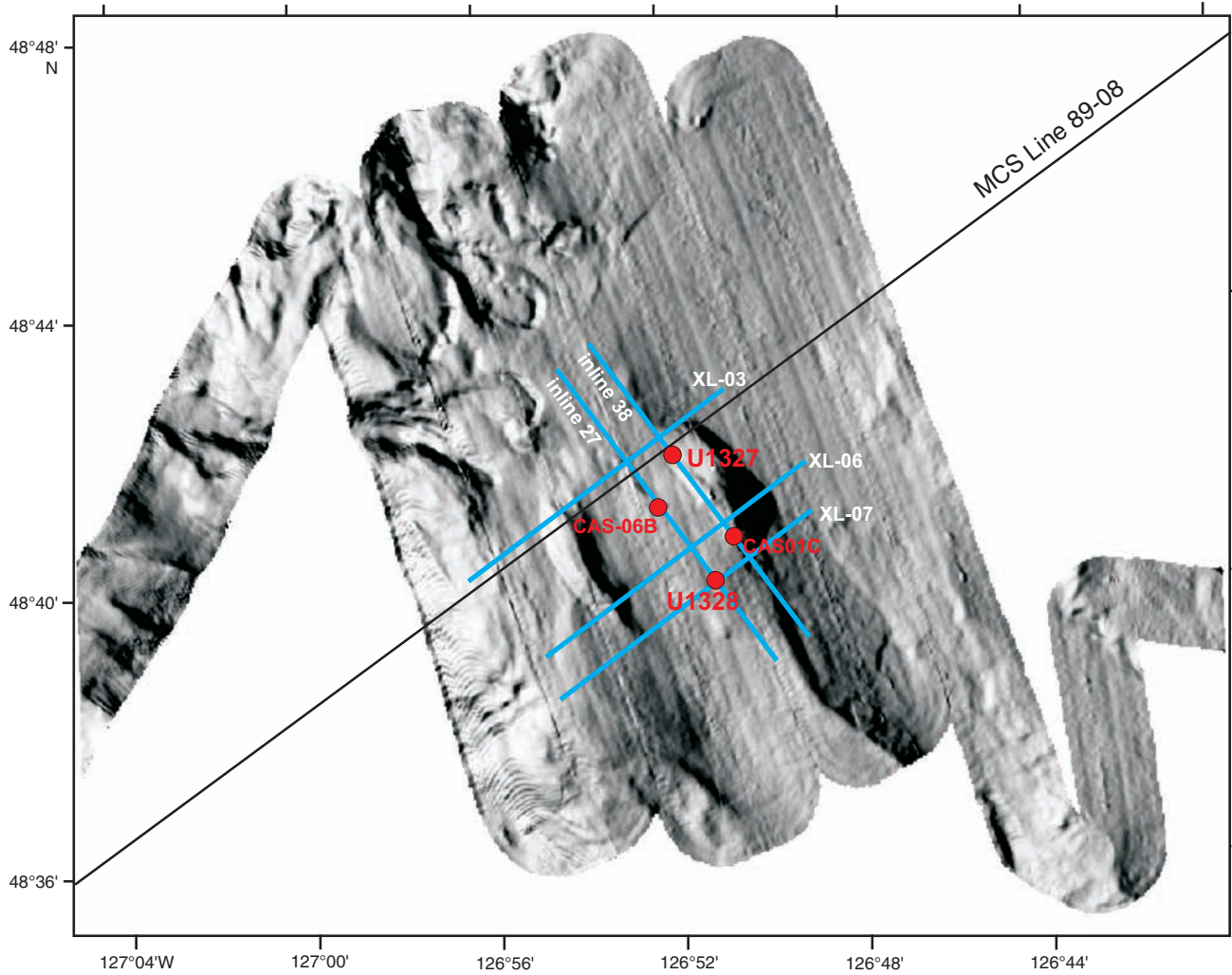


Figure F15. Section from Inline 38 from the 1999 3-D MCS survey over Site U1327. The bottom-simulating reflector (BSR) is clearly visible. The boundary between slope sediments and incoherent accreted sediments is indicated by a dashed line. This boundary is often gradational.

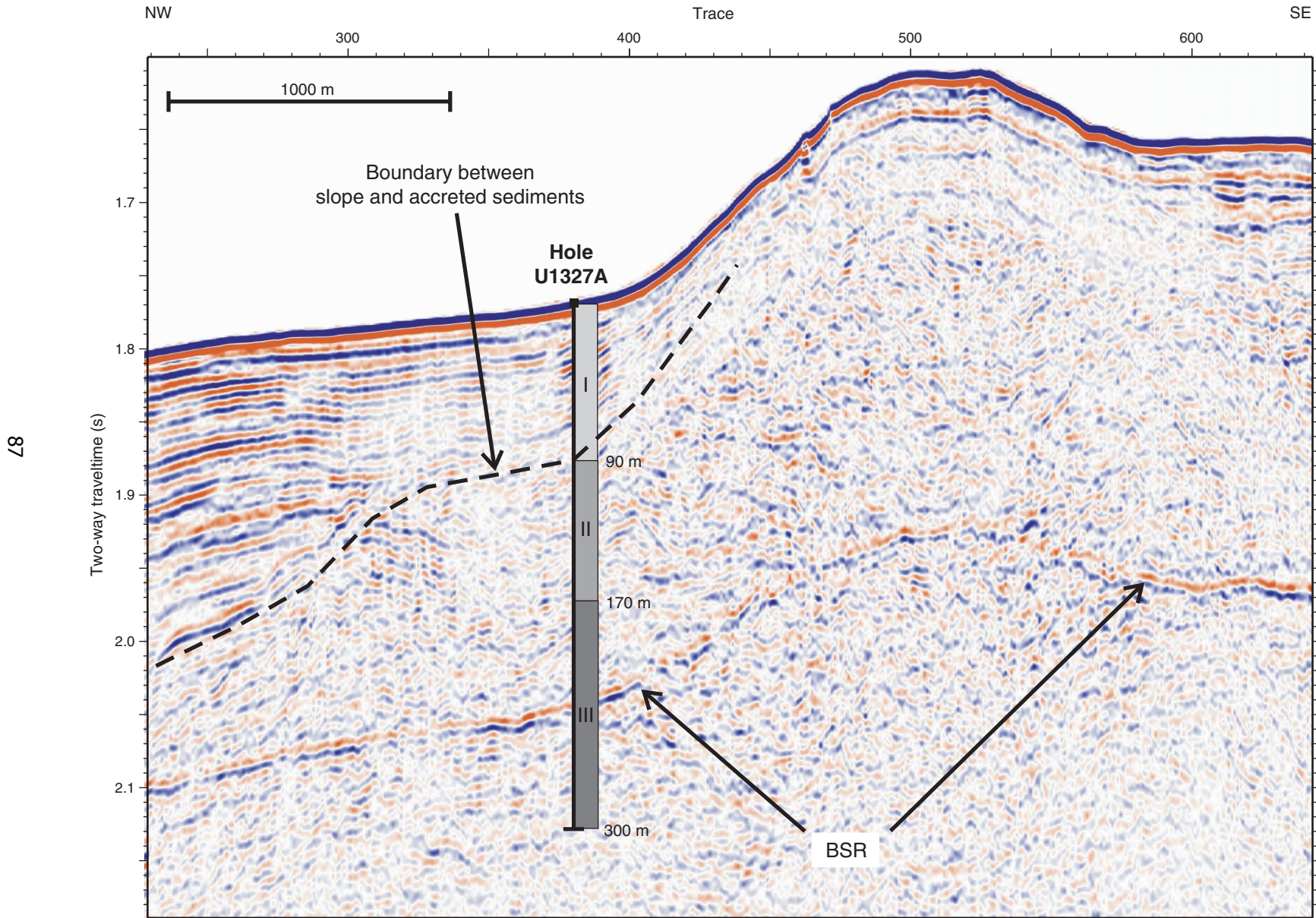


Figure F16. Compilation of seismic data from Inline 38 and key data from Site U1327, including lithostratigraphic units (Hole U1327C), porosity from LWD/MWD data (Hole U1327A) and MAD data (Hole U1327C), resistivity-at-the-bit (RAB) data from LWD/MWD (Hole U1327A), interstitial water chlorinity (Hole U1327C), C_1/C_2 ratio (Hole U1327C), and pore water saturation (S_w) from LWD data (Hole U1327C). Note that there is an offset of ~15 m between Holes U1327A and U1327C in a direction perpendicular to the orientation of Inline 38. The bright band of high electrical reflectivity as seen in the RAB image (identified in Hole U1327A) does not correlate to a prominent seismic event. PCS = pressure core sampler. BSR = bottom-simulating reflector.

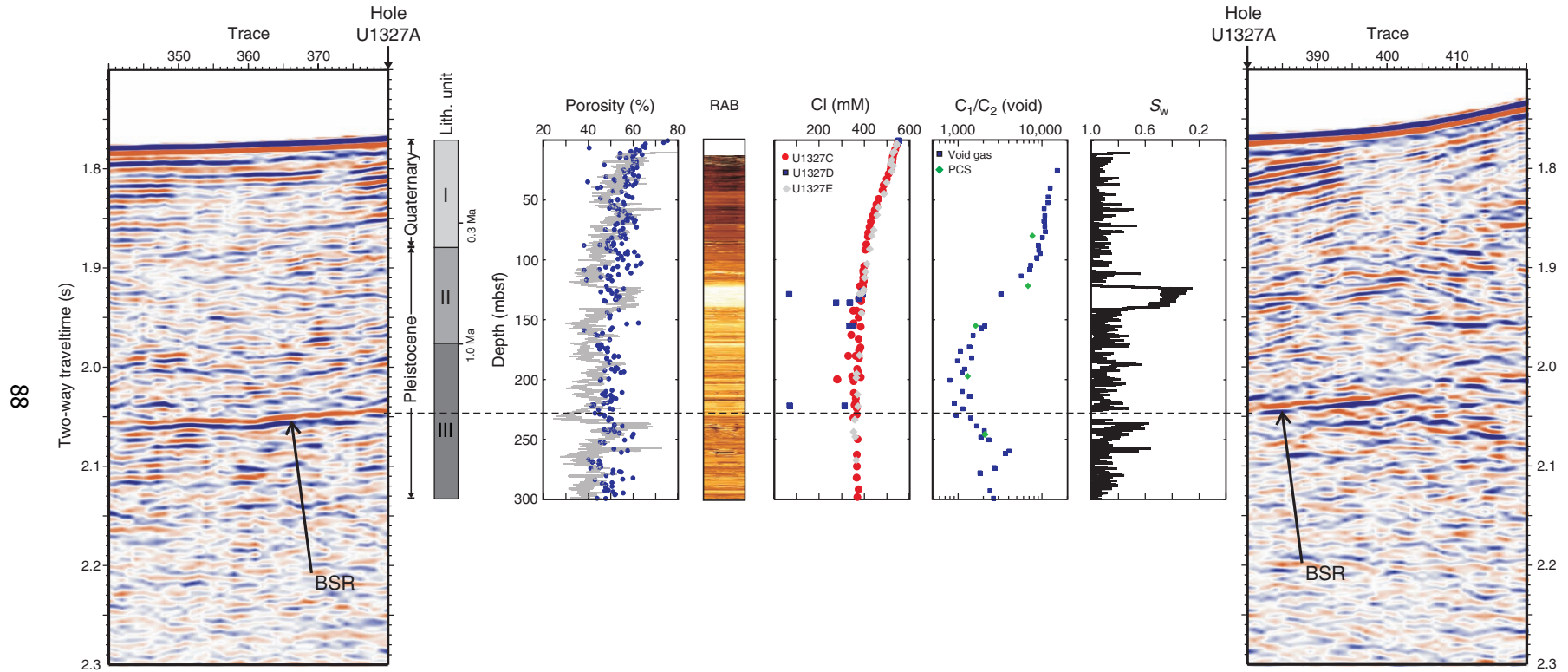


Figure F17. Section of MCS Line XL07 crossing at Site U1328. Shown are locations of bore holes and main lithostratigraphic units (corresponding to Hole U1328A). CDP = common depth point. BSR = bottom-simulating reflector.

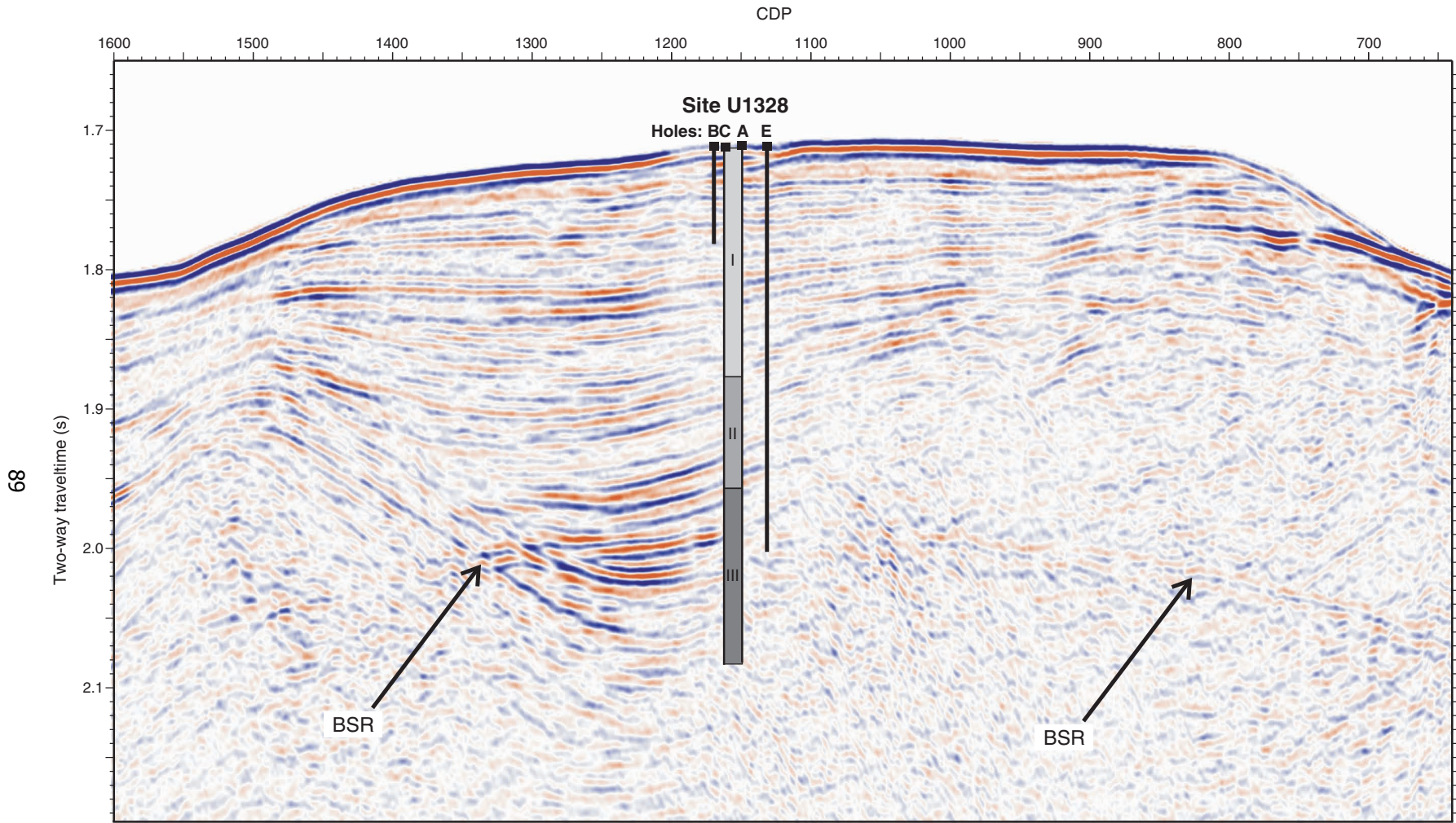


Figure F18. Comparison of major downhole measurements with seismic section from Line XL07 crossing at Hole U1328A. Shown are litho-stratigraphic units and age determination, porosity from logging while drilling (LWD) and moisture and density analyses, resistivity-at-the-bit (RAB) image of LWD resistivity, pore water chlorinity, C_1/C_2 ratio of the void gas, and pore water saturation (S_w) derived from LWD data using Archie's law. BSR = bottom-simulating reflector. CDP = common depth point. PCS = pressure core sampler.

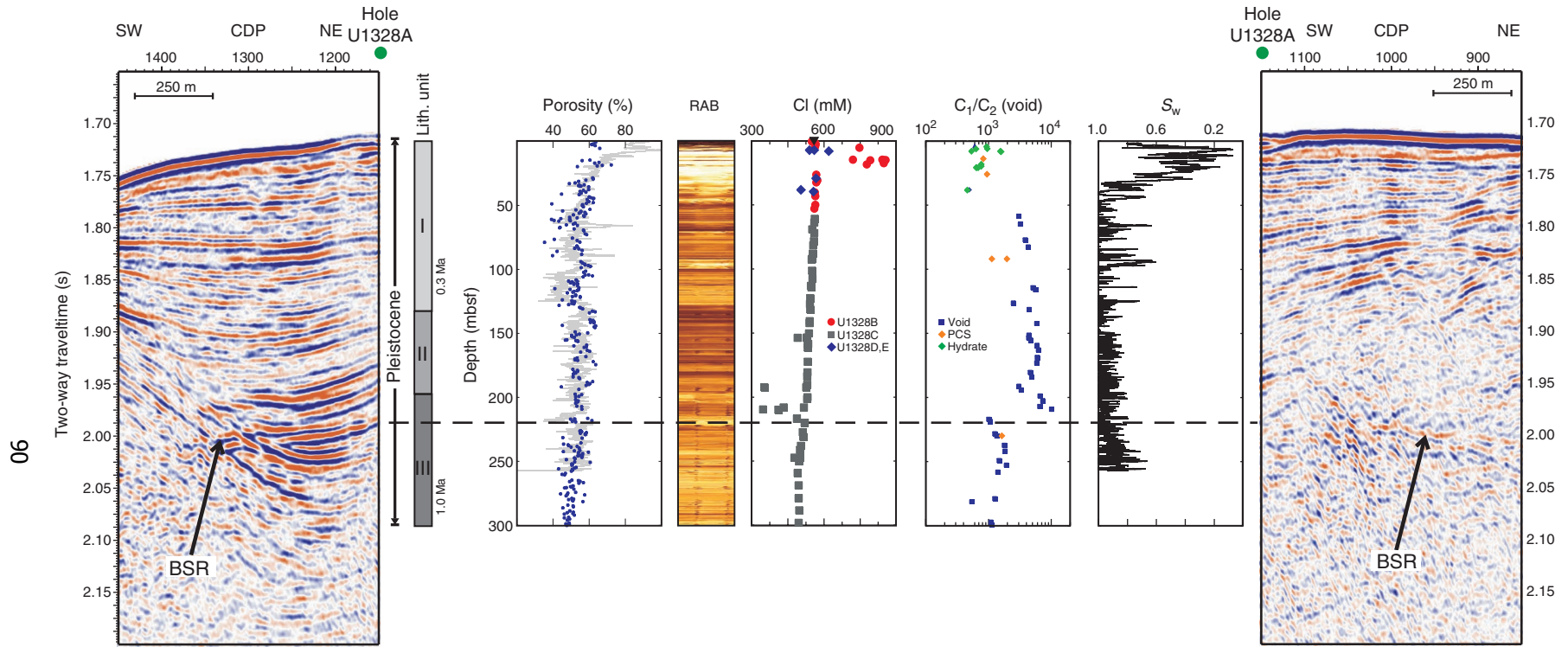


Figure F19. Comparison of major downhole measurements with seismic section from Inline 27 of the 1999 MCS pseudo-3-D data set crossing at Hole U1328A. Shown are lithostratigraphic units and age determination, porosity from logging while drilling (LWD) and moisture and density analyses, resistivity-at-the-bit (RAB) image of LWD resistivity, pore water chlorinity, C_1/C_2 ratio of the void gas, and pore water saturation (S_w) derived from LWD data using Archie's law. BSR = bottom-simulating reflector. PCS = pressure core sampler.

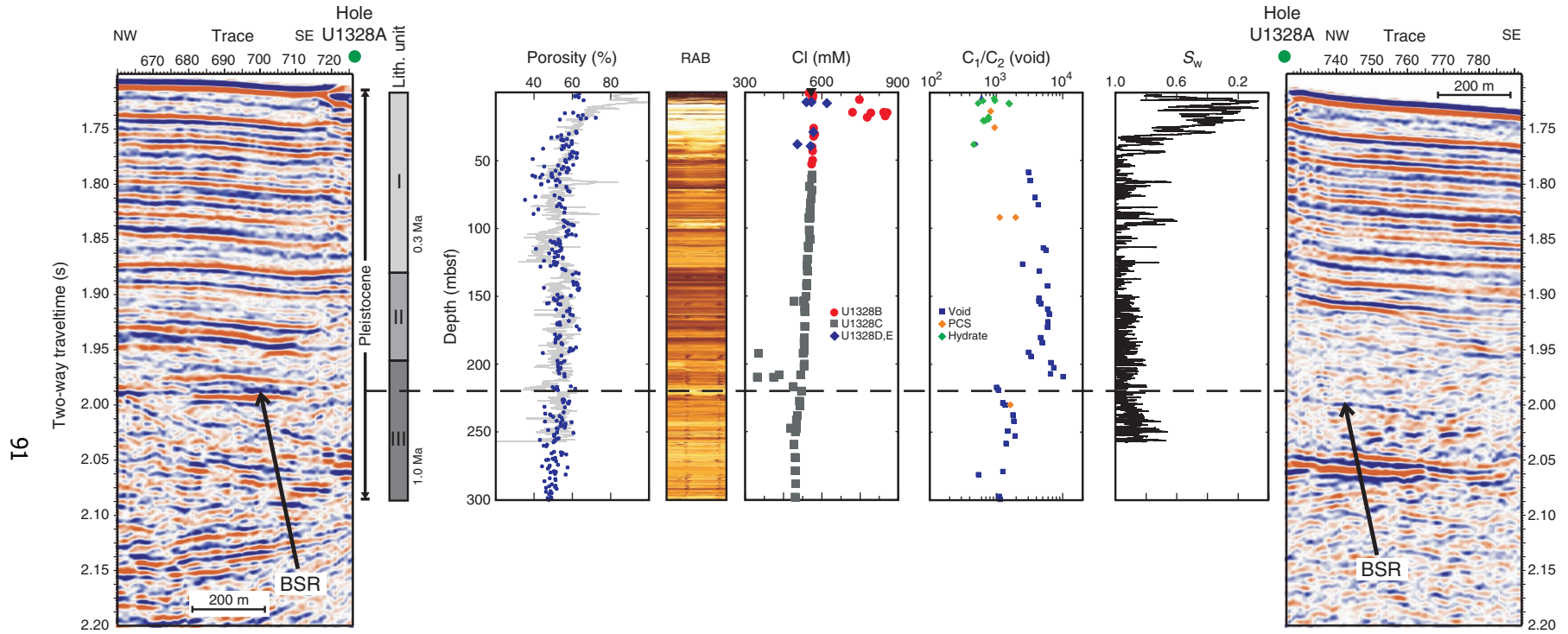


Figure F20. Seismic section from MCS Line 89-08 with overlay of lithostratigraphic units. The base of Unit I may be correlated to a shallow reflector that can be traced across the entire section shown. The boundary between Units II and III is an unconformity, which can be traced clearly between common depth points (CDPs) 3575 and 3650. Vertical exaggeration = 2.0.

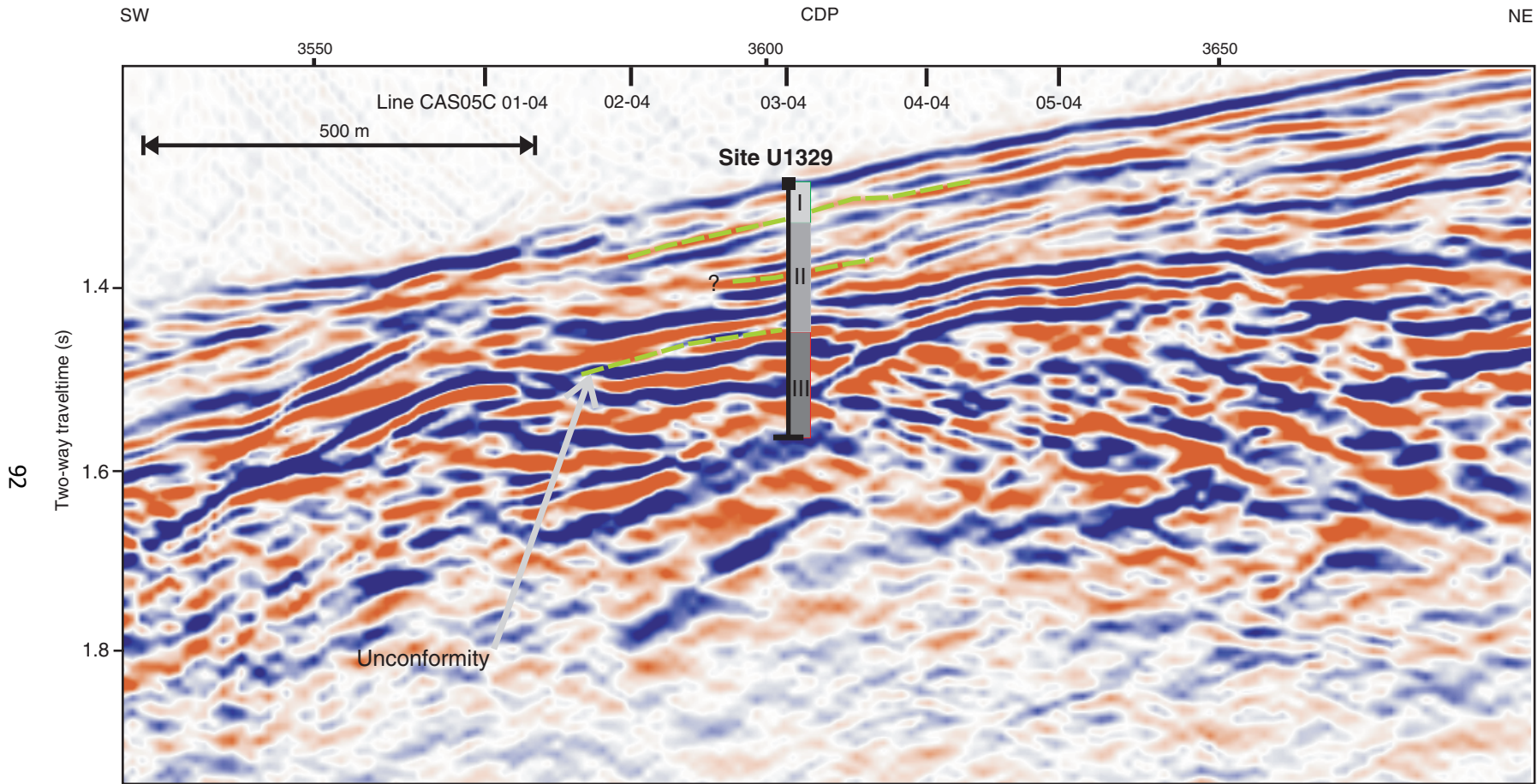


Figure F21. Seismic section from SCS Line CAS05C-01-04 showing some indication of a bottom-simulating reflector (BSR) at ~155 ms TWT (~125 mbsf). The boundary between lithostratigraphic Units II and III is ~10 m deeper than the predicted BSR depth and cannot be clearly identified. Vertical exaggeration = 3.5.

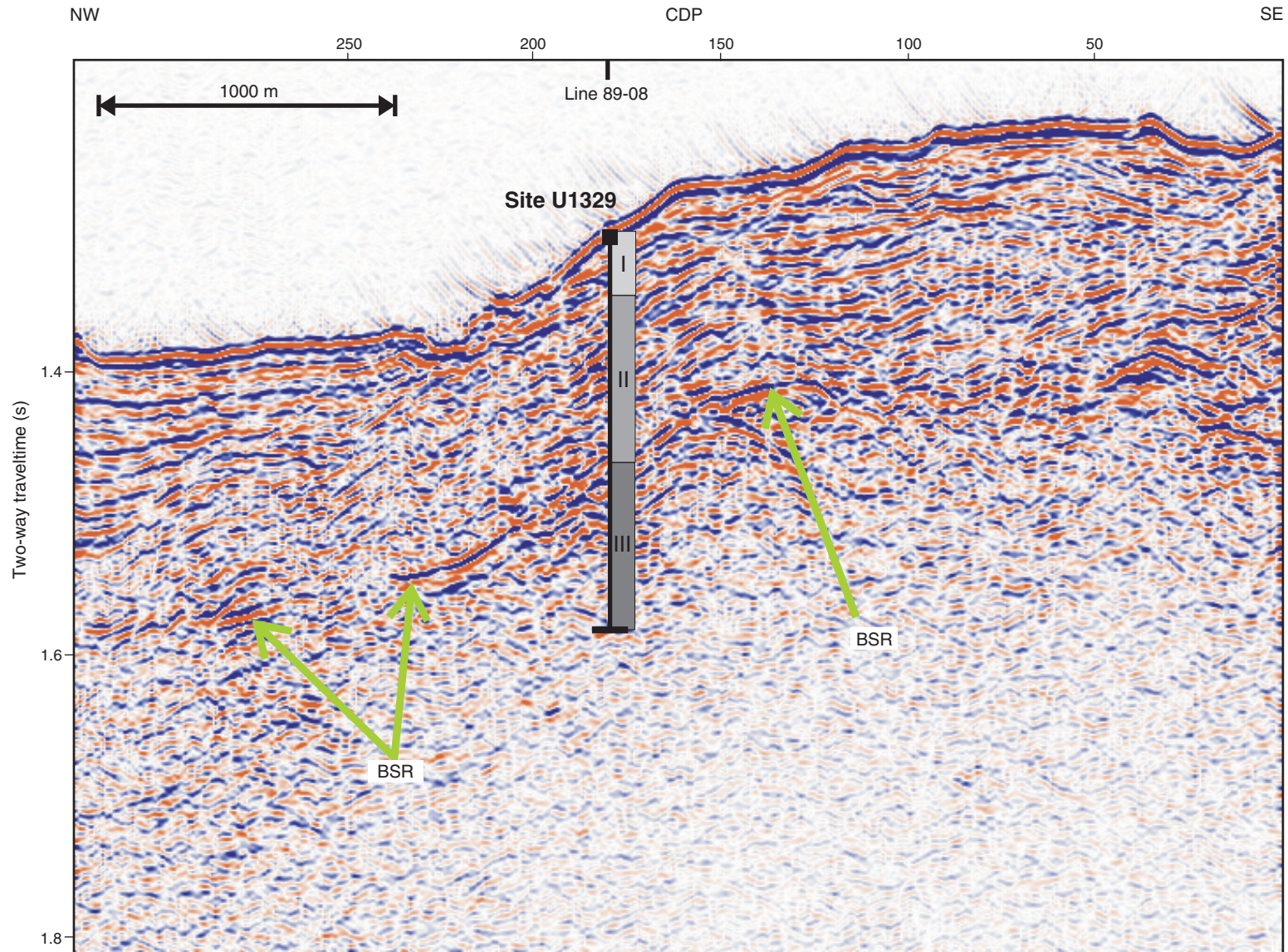


Figure F22. Summary of core and downhole logging data from Site U1329 in comparison to the MCS data from Line 89-08. Shown are litho-stratigraphic units and age determination, porosity from logging while drilling (LWD) and moisture and density (MAD) analyses, resistivity-at-the-bit (RAB) image of LWD resistivity, pore water chlorinity, C_1/C_2 ratio of the void gas, and pore water saturation (S_w) derived from LWD data using Archie's law. BSR = bottom-simulating reflector. CDP = common depth point. PCS = pressure core sampler.

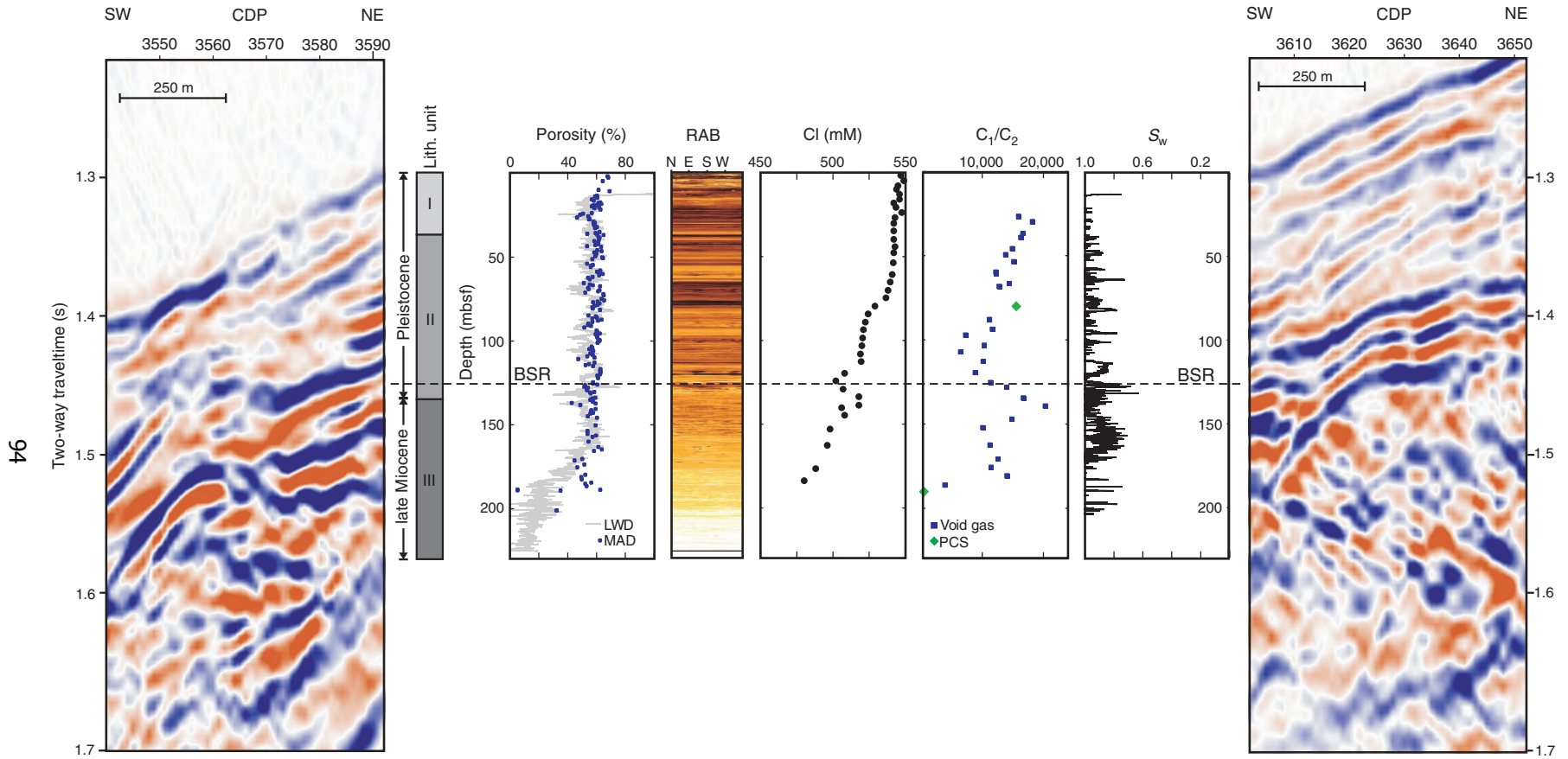
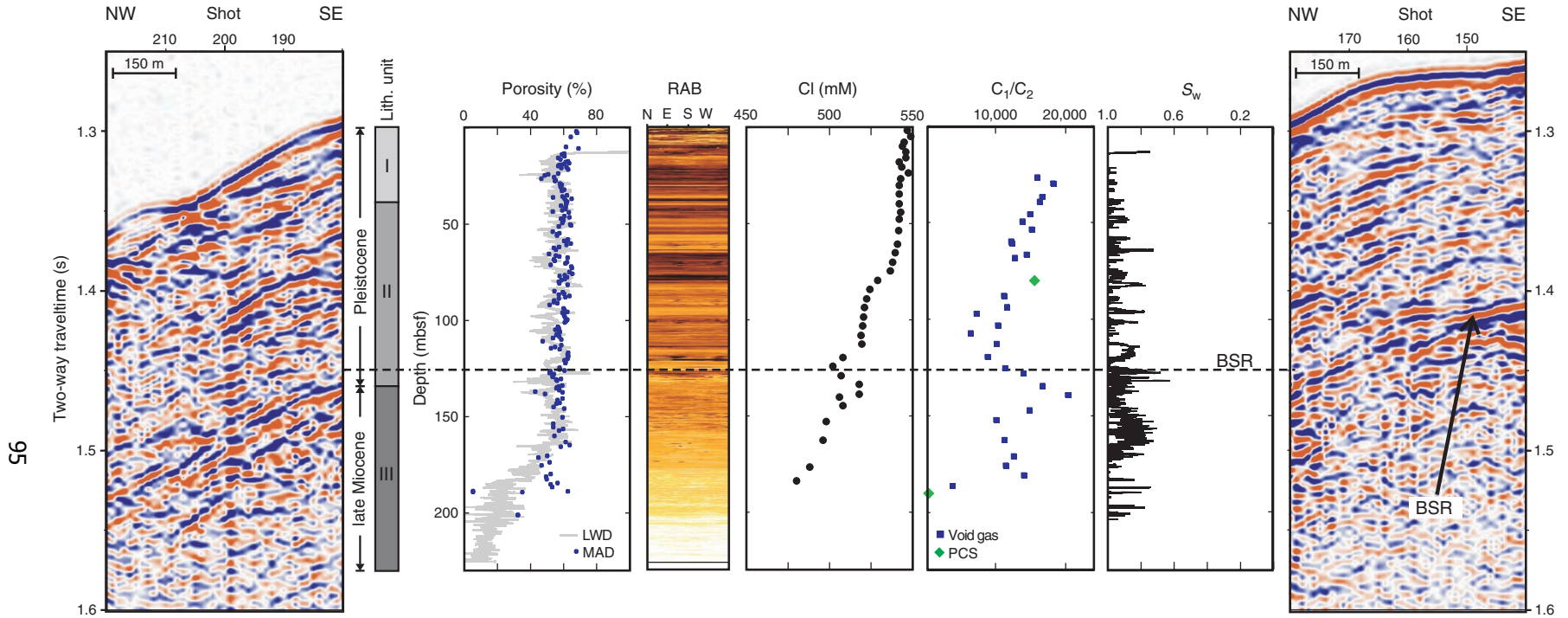


Figure F23. Summary of core and downhole logging data from Site U1329 in comparison to the SCS data from Line CAS05C-01-04. Shown are lithostratigraphic units and age determination, porosity from logging while drilling (LWD) and moisture and density (MAD) analyses, resistivity-at-the-bit (RAB) image of LWD resistivity, pore water chlorinity, C_1/C_2 ratio of the void gas and pore water saturation (S_w) derived from LWD data using Archie's law. BSR = bottom-simulating reflector. PCS = pressure core sampler.



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Figure F24. Images of pore-filling gas hydrate within coarse-grained sands. A–C. Interval 311-U1326C-6X-4, 83–96 cm (44.85 mbsf). D. Interval 311-U1326C-7X-3, 77–85 cm (53.27 mbsf).

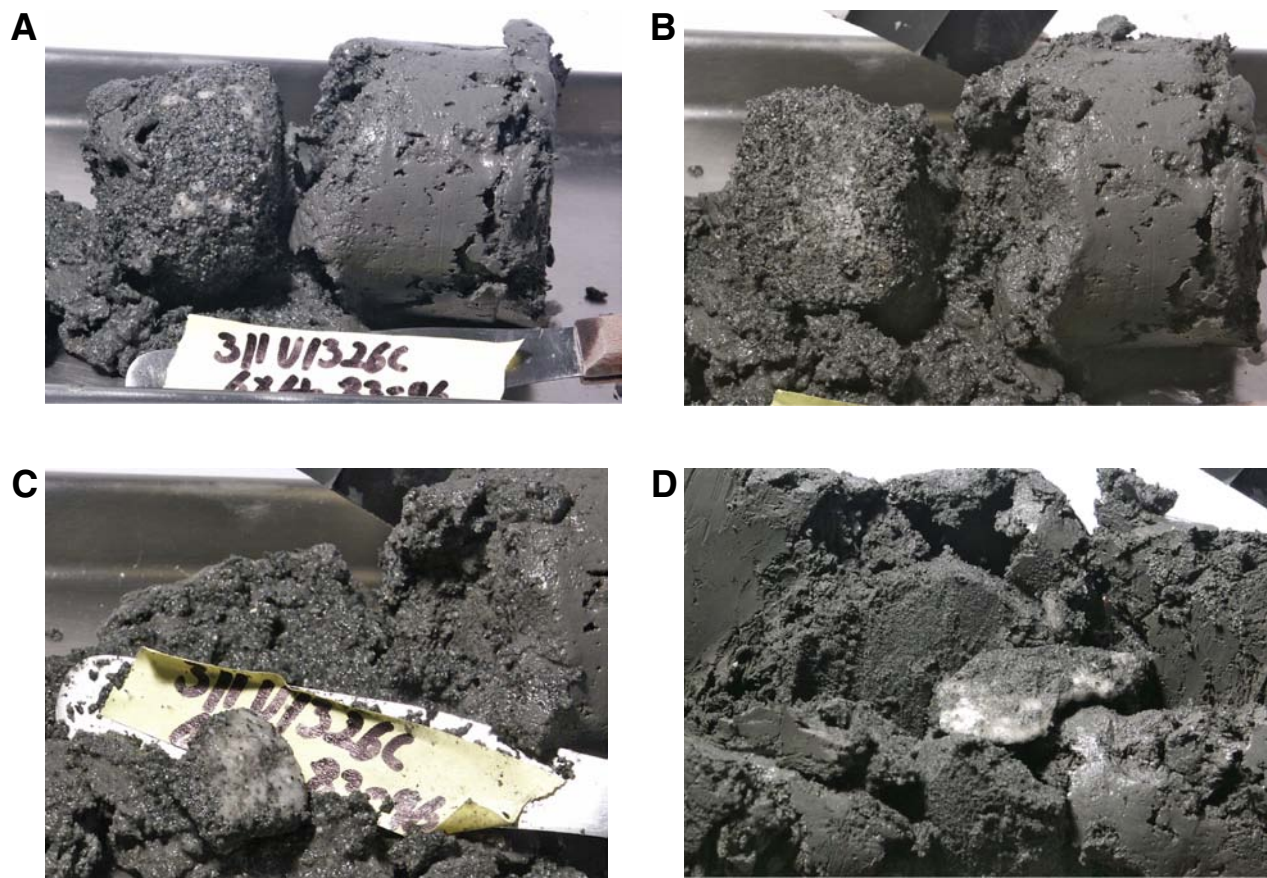


Figure F25. Images of gas hydrate samples recovered at cold vent Site U1328. **A, B.** Interval 311-U1328E-2X-2, 80–90 cm (8.69 mbsf). **C.** Section 311-U1328E-2X-CC.



Figure F26. All reliable in situ temperature measurements from Expedition 311. Data from Leg 146 Site 889 are also included. Two open squares = DVTP measurements at Site U1325 that were of marginal quality due to heave conditions. Postcruise work will attempt to estimate the in situ temperature by modeling heating of borehole by the surrounding rock.

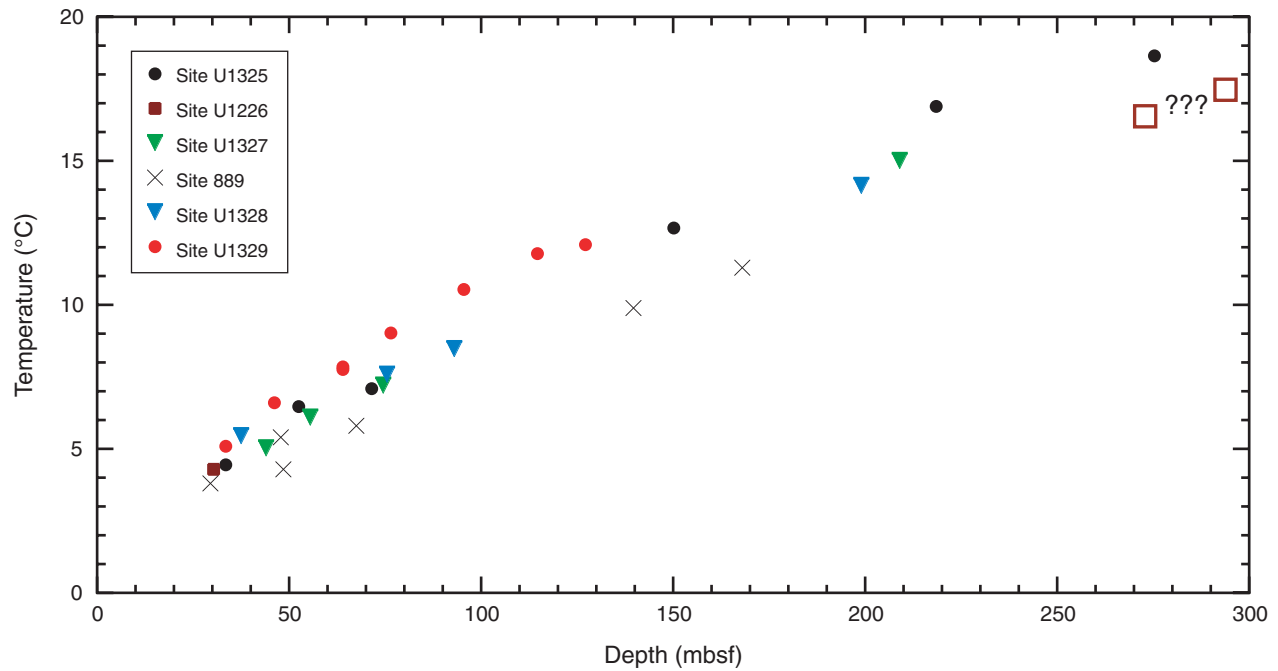


Figure F27. Summary of Expedition 311 interstitial water chlorinity profiles, including average interstitial water Cl from pressure core sampler (PCS) core. Sites 889/890 data are shown for comparison. Also shown are lithostratigraphic units, LWD RAB resistivity images, and location of the bottom-simulating reflector (BSR).

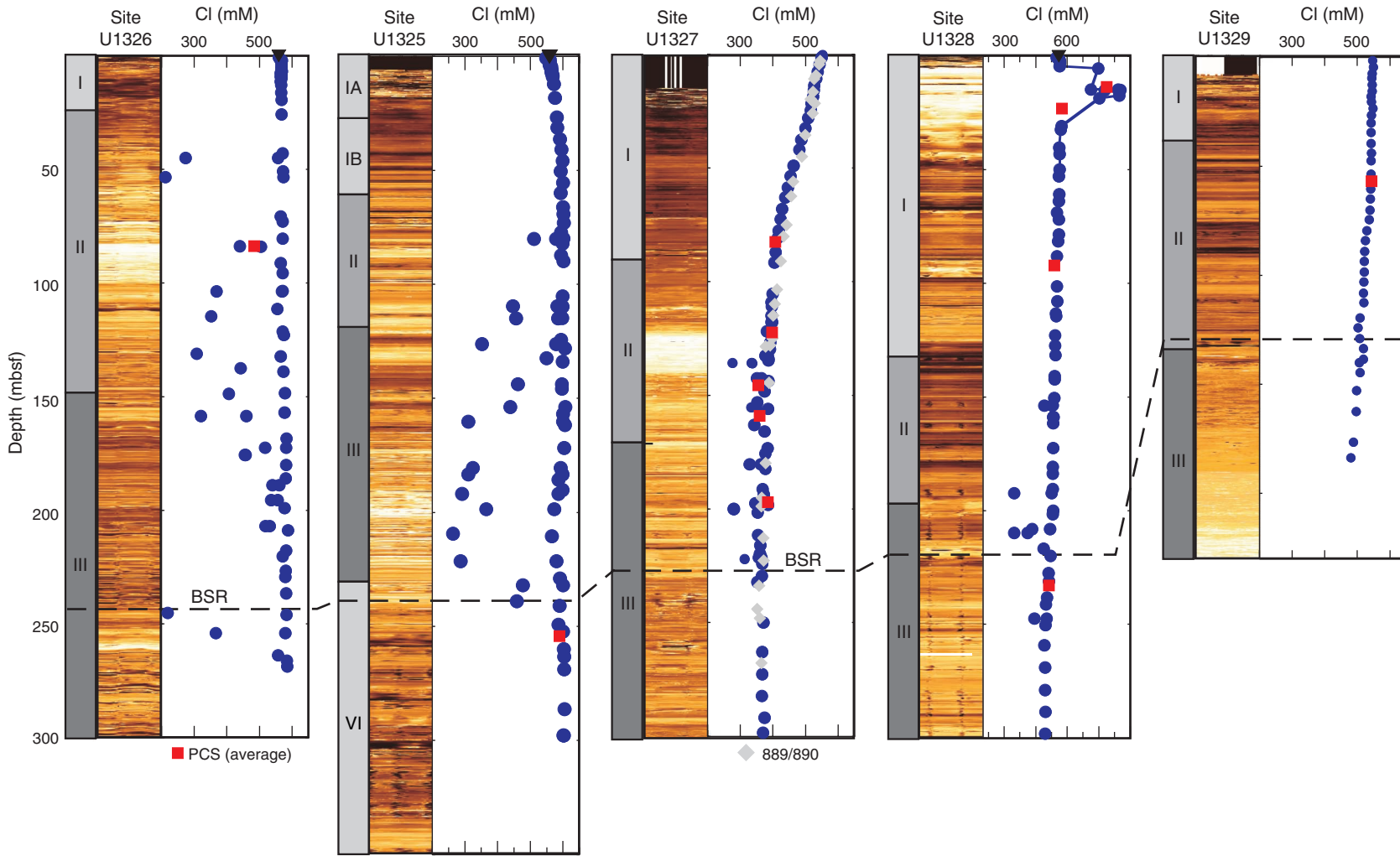


Figure F28. Summary of Expedition 311 compressed infrared images of core liner temperature. Also shown are lithostratigraphic units, LWD resistivity-at-the-bit (RAB) images, pore water saturation (S_w), and location of the bottom-simulating reflector (BSR). PCS = pressure core sampler.

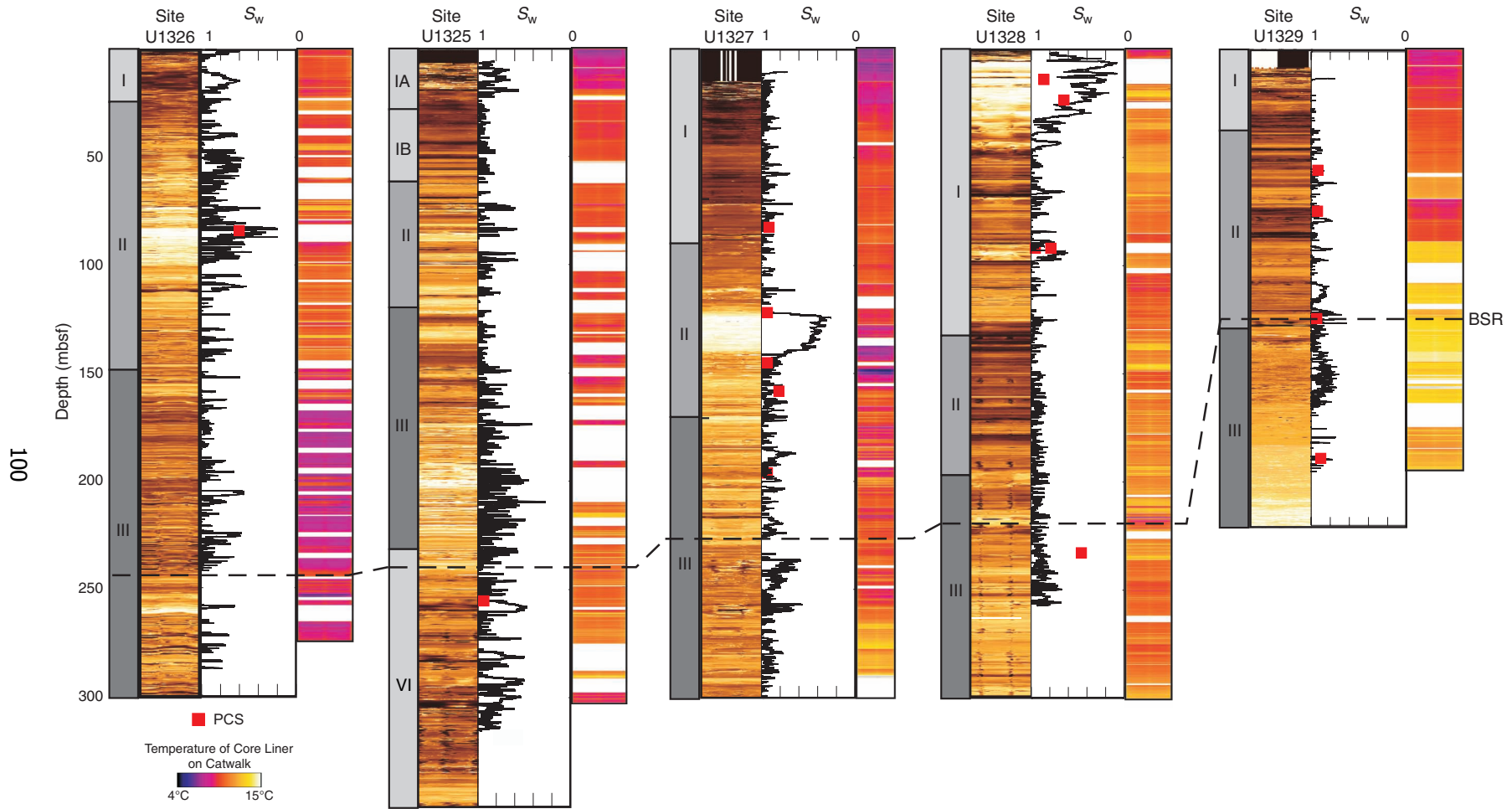
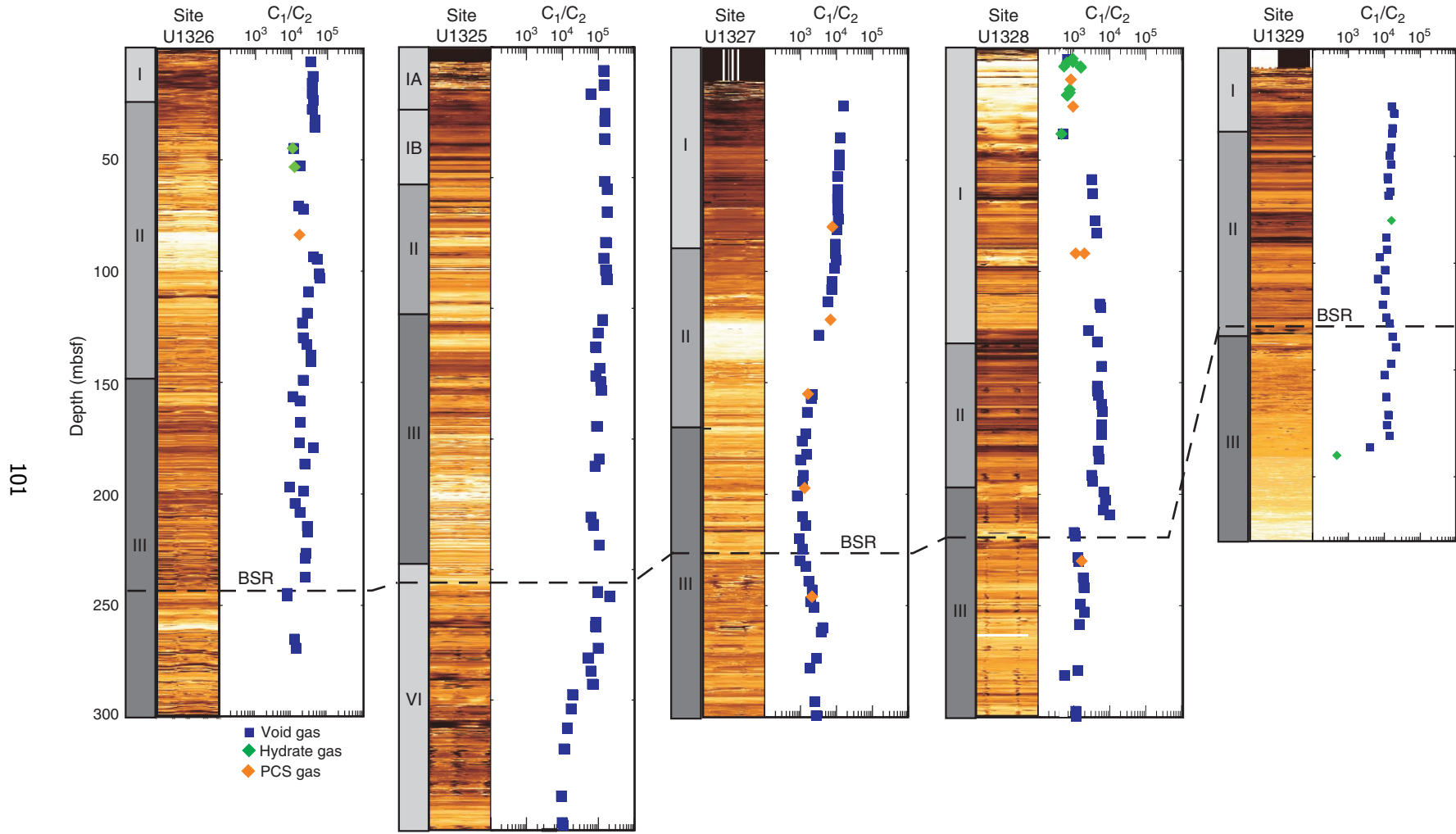


Figure F29. Summary of Expedition 311 C_1/C_2 gas profiles, including data from void, gas hydrate, and pressure core sampler (PCS) gas. Also shown are lithostratigraphic units and LWD RAB resistivity images and location of the bottom-simulating reflector (BSR).



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Figure F30. Summary of Expedition 311 pore water saturations derived from LWD data using Archie's relation and from PCS degassing experiments. Also shown are lithostratigraphic units and LWD RAB resistivity images and location of the BSR.

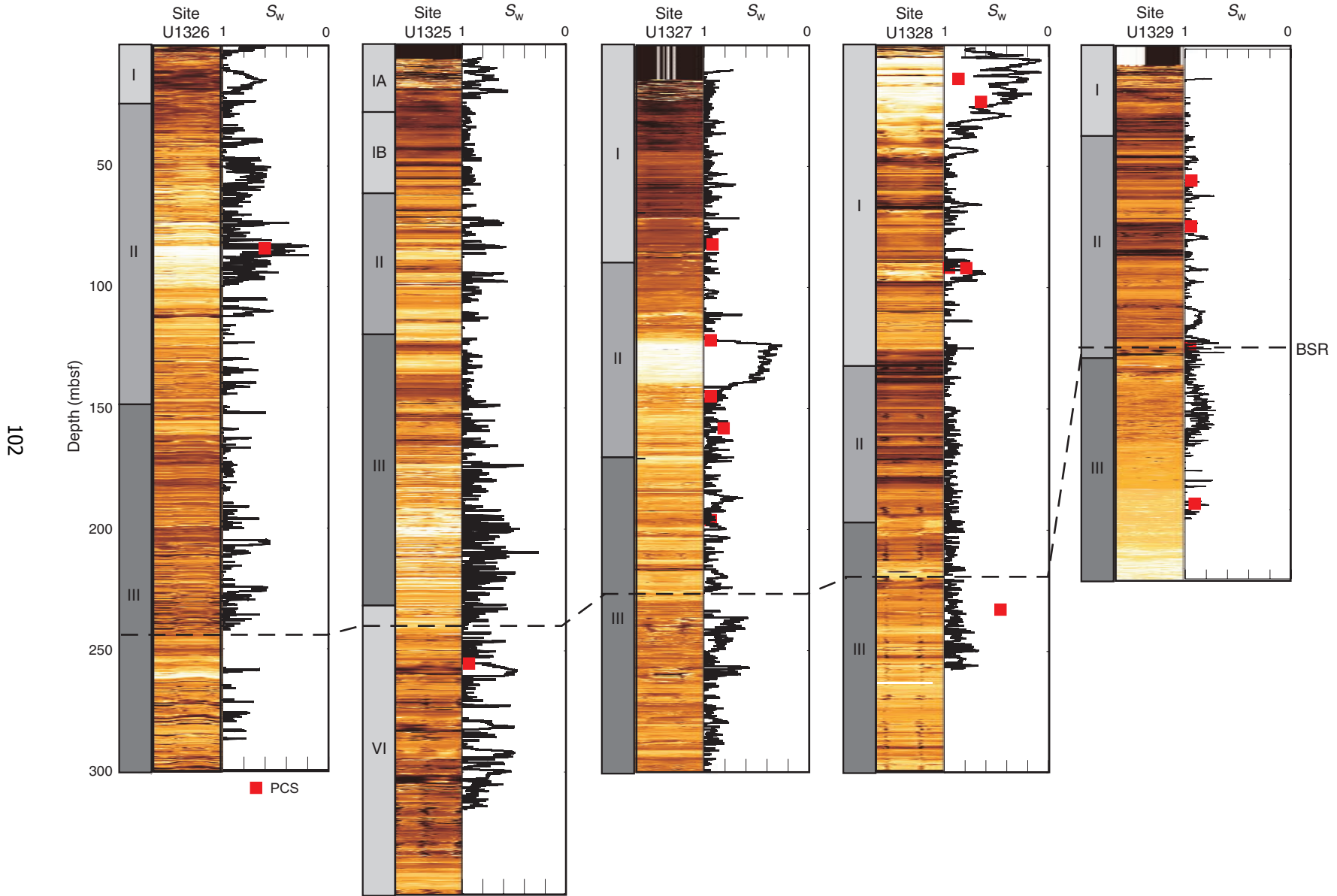


Figure F31. Methane phase diagram for Expedition 311 sites, with total methane concentration measured from depressurized PCS cores. The seafloor temperatures and the salinity profiles were measured, and methane saturation was calculated according to Xu (Xu, 2002, 2004).

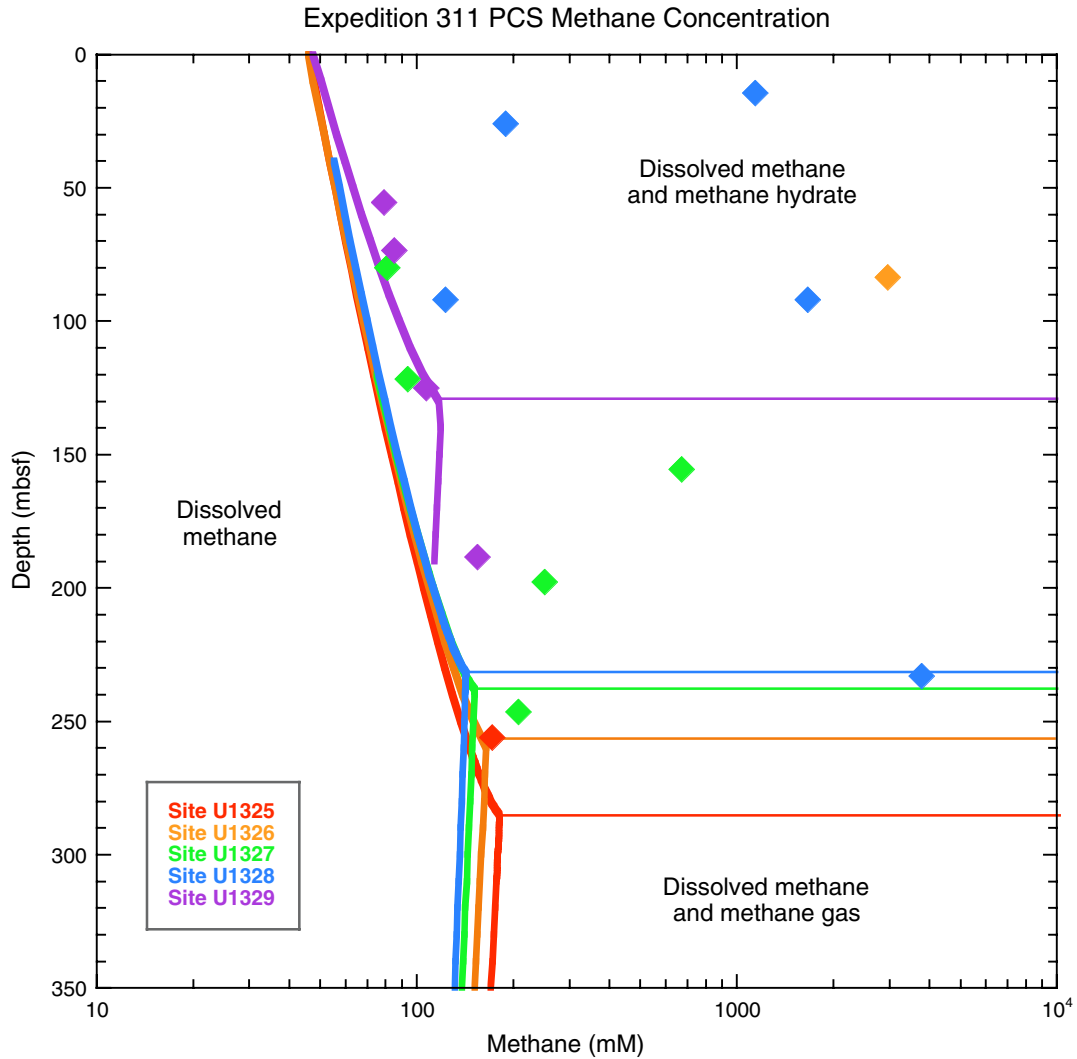


Figure F32. The sulfate/methane interface at Expedition 311 sites. Note the offset at Site U1325, where the mudline core at Hole U1325D showed that Hole U1325B was likely spudded at ~3 mbsf.

