

# Data report: clay mineral assemblages in the Shikoku Basin, NanTroSEIZE subduction inputs, IODP Sites C0011 and C0012<sup>1</sup>

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## Abstract

This report documents clay mineral assemblages outboard of the Nankai Trough at two sites in the Shikoku Basin. The sites make up part of the Kumano transect, offshore south-central Japan. Coring began during Integrated Ocean Drilling Program (IODP) Expedition 322, as part of the Nankai Trough Seismogenic Zone Experiment (NanTroSEIZE). IODP Site C0011 is located on the northwest flank of Kashinosaki Knoll, and IODP Site C0012 is located on the summit of the seamount. A total of 292 samples of hemipelagic mud and mudstone were analyzed by X-ray diffraction, using oriented aggregates of the clay-size fraction (<2 µm). Smectite varies the most among the clay-size constituents, ranging in abundance from 35% to 100% (where smectite + illite + chlorite + kaolinite + quartz = 100%). Specimens with smectite contents of 80% or more are classified as bentonites, and we identified 44 such samples. The estimated amount of smectite in bulk mudstone ranges from 24% to 87% (where total clay minerals + quartz + feldspar + calcite = 100%), and a substantial number of bulk samples contain >45% smectite. On average, the expandability of illite/smectite mixed-layer clay is equal to 77%, and the average proportion of illite in illite/smectite is 11%. There is no systematic progression of clay mineral diagenesis over the depths sampled. Most values of illite crystallinity index are consistent with a detrital source that contains sedimentary rocks that were exposed to anchizone conditions of incipient metamorphism. The smectite-rich clay mineral assemblage is typical of early to middle Miocene deposits within the lower part of the Shikoku Basin.

## Introduction

The Nankai Trough subduction zone is the product of convergence between the Philippine Sea plate and the Eurasian plate (Fig. F1). Many sites have been drilled and cored in this region over the past four decades, including sites from Deep Sea Drilling Project (DSDP) Legs 31 and 87 (Karig, Ingle, et al., 1975; Kagami, Karig, Coulbourn, et al., 1986) and Ocean Drilling Program (ODP) Legs 131, 190, and 196 (Taira et al., 1992; Moore et al., 2001, 2005). Integrated Ocean Drilling Program (IODP) Expeditions 314, 315, and 316 focused on a new transect—the Kumano transect—during Stage 1 of the Nankai Trough Seismogenic Zone Ex-

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periment (NanTroSEIZE) (Ashi et al., 2009; Screaseon et al., 2009; Tobin et al., 2009). Subsequent drilling during IODP Expedition 322 concentrated on inputs to the subduction zone by coring two sites in the Shikoku Basin (see the “[Expedition 322 summary](#)” chapter [Underwood et al., 2010]). IODP Sites C0011 and C0012 are located on the northwest flank and at the summit of a subducting basement high known as Kashinosaki Knoll (Ike et al., 2008) (Fig. F2).

Previous investigations of clay minerals in the vicinity of the Nankai Trough and the Shikoku Basin demonstrated that the hemipelagic mud(stones) change in composition largely as function of depositional age (Cook et al., 1975; Chamley, 1980; Chamley et al., 1986; Fagel et al., 1992; Underwood et al., 1993a, 1993b; Steurer and Underwood, 2003; Underwood and Steurer, 2003). Underwood (2007) summarized findings from the so-called reference sites along the Muroto and Ashizuri transects (ODP Sites 1173 and 1177) (Fig. F1), which are of particular interest for comparisons with Sites C0011 and C0012. Miocene strata throughout the Nankai-Shikoku region tend to contain higher percentages of smectite, whereas Pliocene and Pleistocene deposits are more enriched in illite and chlorite. This temporal trend also exists in the shallow accretionary prism and forearc basin of the Kumano transect (Guo and Underwood, 2012).

The abundance and hydration state of expandable clay minerals are important factors to consider during the NanTroSEIZE project because of their influence on fluid production as depth increases along the plate interface (Saffer et al., 2008). Clay diagenesis (particularly the smectite-to-illite reaction) is more advanced along the Muroto transect (Site 1173), where proximity to the paleospreading center of the subducting Shikoku Basin is responsible for higher heat flow (Underwood and Pickering, 1996; Masuda et al., 1996, 2001; Steurer and Underwood, 2003; Spinelli and Underwood, 2005; Saffer et al., 2008). In contrast, there is no evidence for presubduction diagenesis at Site 1177 (Steurer and Underwood, 2003). To document clay composition in the Kumano transect area, particularly for those stratigraphic units that eventually affect conditions near the subduction megathrust, we analyzed the clay mineral assemblages from 292 samples of hemipelagic mud and mudstone using X-ray diffraction (XRD). This report documents how the common clay minerals (smectite, illite, chlorite, and kaolinite) change in relative abundance as a function of depositional age and lithostratigraphy. We also test whether or not smectite-to-illite diagenesis has progressed to any measurable extent prior to the arrival of sedimentary strata at the subduction front.

## Methods

### Calculations of mineral abundance

Marine sediment samples can be analyzed by XRD using a variety of techniques. For example, the presence of a specific detrital and/or authigenic mineral can be detected easily through visual recognition of characteristic (hkl) peak positions. It is more challenging, however, to estimate the relative abundance of a mineral in bulk sediment or in the clay-size fraction with meaningful accuracy (e.g., Moore, 1968; Heath and Pisias, 1979; Johnson et al., 1985). The most common approach for analyzing clays in marine geology has been to multiply the Biscaye (1965) weighting factors by the peak areas of basal reflections and normalize to 100% (McManus, 1991). Errors can be substantial, however, and accuracy of calculated values is affected by the absolute abundance by weight of each mineral in the mixture (Underwood et al., 2003). XRD results also change depending on sample disaggregation technique, chemical pretreatments, particle size separation, crystallinity and chemical composition of minerals, peak-fitting algorithms, and the degree of preferred orientation of crystallites (e.g., Moore and Reynolds, 1989; Ottner et al., 2000). Even though data reproducibility might be very good, accuracy is usually no better than  $\pm 10\%$  unless the analytical methods include calibration with internal standards, use of single-line reference intensity ratios, and some fairly elaborate sample preparation steps to create random particle orientations (Środoń et al., 2001; Omotoso et al., 2006).

One goal of NanTroSEIZE is to obtain internally consistent, semiquantitative estimates of mineral abundance in the clay-size fraction for a large number of samples. To accomplish this, we use a matrix of singular value decomposition (SVD) normalization factors, as documented in full detail by Underwood et al. (2003). Figure F3 shows representative X-ray diffractograms for two clay-size aggregates from the Shikoku Basin. The matrix of SVD factors (Table T1) is applied to the integrated areas of a broad smectite (001) peak centered at  $\sim 5.3^\circ 2\theta$  (d-value = 16.5 Å), the illite (001) peak at  $\sim 8.9^\circ 2\theta$  (d-value = 9.9 Å), the composite chlorite (002) + kaolinite (001) peak at  $12.5^\circ 2\theta$  (d-value = 7.06 Å), and the quartz (100) peak at  $20.85^\circ 2\theta$  (d-value = 4.26 Å). Average errors for the standard mineral mixtures used to calibrate this method are approximately 3% for smectite, 1% for illite, 2% for chlorite, and 1.4% for quartz (Underwood et al., 2003). Because of the nearly total overlap between the kaolinite (001) and chlorite (002) reflections, we first calculate that relative abundance as undifferentiated chlorite + kaolinite, and then



solve for the proportion of each mineral using the double peak at  $\sim 25^\circ 2\theta$  (Fig. F3) and a refined version of the Biscaye (1964) method, as documented fully by Guo and Underwood (2011). Analysis of standard mineral mixtures shows that the average error for the chlorite/kaolinite ratio is 2.6%. To provide an estimate of the abundance of individual clay minerals in the bulk mudstone, we also multiply each relative percentage among the clay minerals (i.e., excluding quartz) by the weight percent of total clay minerals from shipboard bulk powder XRD analyses of co-located “cluster” specimens (e.g., see the “[Site C0011](#)” chapter [Expedition 322 Scientists, 2010a]). To facilitate comparisons with many of the other published data sets from the region, data tables include weighted peak area percentages for smectite, illite, chlorite, and kaolinite using Biscaye (1965) weighting factors. These values are relative percentages and should be regarded as semiquantitative.

To characterize the extent of clay diagenesis, we used the saddle/peak method of Rettke (1981) to calculate the percent expandability of smectite and illite/smectite (I/S) mixed-layer clay. This method is sensitive to the proportions of discrete illite (I) versus I/S mixed-layer clay. Our calculations follow a curve for 1:1 mixtures of I and I/S. A complementary way to calculate the proportion of illite in the I/S mixed-layer phase is based on the position (d-value) of the (002/003) peak (following Moore and Reynolds, 1989) after correcting the diffractogram peaks for misalignment of the detector and sample holder. We also report values of illite crystallinity index as the peak width measured at half height ( $\Delta^\circ 2\theta$ ) for the (001) reflection.

### Sample preparation

Isolation of clay-size fractions starts with air drying and gentle hand-crushing of the mud/mudstone with a mortar and pestle, after which specimens are immersed in 3%  $H_2O_2$  for at least 24 h to digest organic matter. We then add  $\sim 250$  mL of Na hexametaphosphate solution (concentration of 4 g/1000 mL distilled  $H_2O$ ) and insert the beakers into an ultrasonic bath for several minutes to promote disaggregation and deflocculation. These steps are repeated until disaggregation is complete. Washing consists of two passes through a centrifuge (8200 revolutions per minute [rpm] for 25 min;  $\sim 6000$  g) with resuspension in distilled-deionized water after each pass. After transferring the suspended sediment to a 60 mL plastic bottle, each sample is resuspended by vigorous shaking and a 2 min application of a sonic cell probe. The clay-size splits ( $< 2 \mu m$  spherical equivalent settling diameter) are then separated by centrifugation (1000 rpm for 2.4 min;  $\sim 320$  g). Oriented

clay aggregates are prepared using the filter-peel method (Moore and Reynolds, 1989) and  $0.45 \mu m$  membranes. The clay aggregates are saturated with ethylene glycol vapor for at least 24 h prior to XRD analysis, using a closed vapor chamber heated to  $60^\circ C$ .

### X-ray diffraction parameters

The XRD laboratory at the University of Missouri (USA) utilizes a Scintag Pad V X-ray diffractometer with  $CuK\alpha$  radiation ( $1.54 \text{ \AA}$ ) and a Ni filter. Scans of oriented clay aggregates are run at 40 kV and 30 mA over a scanning range of  $3^\circ$ – $26.5^\circ 2\theta$ , a rate of  $1^\circ 2\theta/\text{min}$ , and a step size of  $0.01^\circ 2\theta$ . Slits are 0.5 mm (divergence) and 0.2 mm (receiving). The digital data are processed using MacDiff software (version 4.2.5) to establish a baseline of intensity, smooth counts, correct peak positions offset by misalignment of the detector (using the quartz [100] peak at  $20.95^\circ 2\theta$ ; d-value =  $4.24 \text{ \AA}$ ), determine peak intensity (counts/step), and calculate integrated peak areas (total counts). This program also calculates peak width at half height ( $\Delta^\circ 2\theta$ ).

## Results

Almost all of the samples analyzed in this study were selected from co-located clusters immediately adjacent to the whole-round samples used for shipboard analyses of interstitial water chemistry and shore-based tests of frictional, geotechnical, and hydrogeological properties. Each cluster includes a specimen for shipboard bulk powder XRD analysis, which provided estimates of the relative abundance of total clay minerals (see the “[Site C0011](#)” and “[Site C0012](#)” chapters [Expedition 322 Scientists, 2010a, 2010b]). All of the values of XRD peak-area (total counts) for minerals in the clay-size fraction are tabulated in Table T2. Table T3 lists the calculated values of mineral abundance (weight percent) using SVD normalization factors, as well as area percent using the Biscaye (1965) peak-area weighting factors. Specimens with 80% or more smectite in the clay-size fraction are classified in this report as bentonites. For smectite, we calculated its abundance in the bulk mudstone by multiplying the percentage of smectite within each clay mineral assemblage (where smectite + illite + chlorite + kaolinite = 100%) by the percentage of total clay minerals in the same bulk mudstone (where total clay minerals + quartz + feldspar + calcite = 100%) (for bulk powder data, see the “[Site C0011](#)” and “[Site C0012](#)” chapters [Expedition 322 Scientists, 2010a, 2010b]). Indicators of clay diagenesis (I/S expandability, percent illite in I/S, and illite crystallinity index) are tabulated in Table T4.



The descriptions below are meant to highlight temporal variations in clay composition organized by lithostratigraphic unit, and Table T5 provides a statistical comparison among the units.

### Site C0011

Shipboard scientists during Expedition 322 divided the stratigraphic column at Site C0011 into five lithologic units (see the “[Site C0011](#)” chapter [Expedition 322 Scientists, 2010a]). Because of time limitations, sampling began at 340 m core depth below seafloor (CSF) rather than the mudline, and the hole was abandoned ~170 m above the basement total depth target because of premature destruction of the drill bit. The provisional lithologic characterization of Unit I (hemipelagic/pyroclastic facies) was based on logging-while-drilling results, but the upper part of the sedimentary section was eventually verified by coring during IODP Expedition 333 (Expedition 333 Scientists, 2011).

The first cores recovered during Expedition 322 fall within the upper Miocene epoch (~7.6 Ma) and define the top of Unit II (Fig. F4). This volcanic turbidite facies consists of hemipelagic mudstone, sandstone/siltstone turbidites with abundant pyroclastic debris, and mass transport deposits. Smectite is the most abundant clay-size mineral in Unit II (average = 65%) followed by illite (average = 24%). Both of these clay minerals show considerable amounts of scatter. The chlorite + kaolinite content averages 8% within this unit, and clay-size quartz averages 2%. We sampled 11 bentonites within Unit II, and our estimates for the proportion of smectite in the bulk mudstones range from 29% to 87% (average = 42%) (Fig. F5; Table T5).

The dominant lithology of Unit III (hemipelagic facies) is bioturbated mudstone (silty clay to clayey silt); these strata range in age from 9.1 to 12.2 Ma (Fig. F4). Smectite is the dominant clay-size mineral within this unit, making up an average of 59% of the clay-size fraction. Illite content ranges from 10% to 33% (average = 27%). Values of chlorite + kaolinite average 10%, and clay-size quartz content averages 4%. We found two identifiable bentonites within Unit III. Figure F5 shows that smectite abundance in bulk mudstone ranges from 31% to 73% (average = 42%).

Unit IV (silty turbidite facies) is composed of silty claystone, fine-grained siliciclastic sandstone, and siltstone ranging in age from 12.2 to 14.0 Ma (Fig. F4). Core recovery within this unit was incomplete. Smectite (average = 68%) is the most abundant mineral in the clay-size fraction, followed on average by illite (20%), chlorite + kaolinite (7%), and quartz

(5%). Unit IV contains five bentonite samples, and smectite abundance in bulk mudstone ranges from 25% to 85% (Fig. F5).

Recovery of Unit V (volcaniclastic facies) was very poor because of technical difficulties, and drilling was terminated at 881 m drilling depth below seafloor (DSF). The estimated depth to the top of basaltic basement is 1050 m seismic depth below seafloor (SSF). Based on limited recovery, the lithology of Unit V includes silicic tuff and tuffaceous sandstone (Fig. F4). Mudstone interbeds contain between 45% and 100% smectite in the clay-size fraction (average = 76%). Average values for illite, chlorite + kaolinite, and quartz are 13%, 4%, and 6%, respectively. The volcaniclastic facies includes seven identifiable bentonites, and smectite abundance in bulk mudstone ranges from 31% to 91% (average = 57%) (Fig. F5).

Figure F6 shows that the expandability of I/S mixed-layer clay minerals does not change significantly from the top of Unit II to base of the cored interval. The average value of I/S expandability is 77%. The proportion of illite in the I/S averages 11% and ranges from 0% to 27%. These plots do not reveal any meaningful progression of smectite-to-illite diagenesis as a function of depth, although it is important to remember that coring terminated ~170 m above the top of igneous basement. Heat flow data collected during Expedition 333 indicate the temperature is equal to ~80°C at the sediment/basalt interface (Expedition 333 Scientists, 2011), so diagenesis could advance to higher proportions of illite in I/S below the deepest core recovery. Values of illite crystallinity index typically range between 0.42 and 0.25 Δ°2θ. Such values are consistent with erosion of the illite as a detrital constituent from sedimentary source terrains that had been exposed to anchizone conditions of incipient metamorphism (e.g., Blenkinsop, 1988; Kisch, 1990).

### Site C0012

Six sedimentary units rest above igneous basement at Site C0012 (Fig. F7), which is positioned at the summit of Kashinosaki Knoll (Fig. F2). The overall trend displayed by clay mineral assemblages shows progressive downsection decreases in illite content coupled with consistent increases in smectite (Fig. F7). Values of smectite also display considerable scatter, particularly within the lower half of the section. The contents of chlorite, kaolinite, and quartz are subsidiary to smectite throughout.

The principal lithology of Unit I (hemipelagic/pyroclastic facies) is hemipelagic mud (silty clay to clayey silt) with numerous interbeds of volcanic ash. Recovery of this unit was poor, however, and drilling dis-



turbance was severe. Consequently, the upper part of the sedimentary section was cored again during Expedition 333 (Expedition 333 Scientists, 2011). Smectite is the most abundant clay-size mineral in Unit I (average = 53%), followed by illite (average = 32%), chlorite + kaolinite (average = 14%), and quartz (average = 1%). We found only one bentonite at the top of this unit (cored during jet-in). Estimates for the proportion of smectite in the bulk mud range from 25% to 56% (average = 36%) (Fig. F8; Table T5).

The age of Unit II (volcanic turbidite facies) ranges from 7.8 to 9.4 Ma, and the sediment consists of hemipelagic mudstone, sandstone/siltstone turbidites with abundant pyroclastic material, and mass transport deposits (Fig. F7). The most abundant clay-size mineral in Unit II is smectite (average = 63%), followed by illite (average = 26%), chlorite + kaolinite (average = 9%), and quartz (average = 2%). We found only one bentonite within this unit. Estimates for the proportion of smectite in the bulk mudstones range from 26% to 62% (average = 43%) (Fig. F8; Table T5).

The dominant lithology of Unit III (hemipelagic facies) is bioturbated mudstone (silty clay to clayey silt); these homogeneous strata range in age from 9.4 to 12.7 Ma (Fig. F7). Smectite is the dominant clay-size mineral within this unit, making up an average of 64% of the clay-size fraction. Illite content ranges from 13% to 35% (average = 25%). Values of chlorite + kaolinite average 8%, and clay-size quartz content averages 5%. We found four identifiable bentonites in the lower portion of Unit III. Smectite abundance in bulk mudstone ranges from 38% to 62% (average = 49%).

Unit IV (silty turbidite facies) is composed of silty claystone, fine-grained siliciclastic sandstone, and siltstone ranging in age from 12.7 to 13.5 Ma (Fig. F7). Smectite (average = 66%) is the most abundant mineral in the clay-size fraction, followed on average by illite (21%), chlorite + kaolinite (8%), and quartz (5%). Unit IV contains two bentonites. Smectite abundance in bulk mudstone ranges from 39% to 87% (Fig. F8).

Recovery of Unit V (siliciclastic/volcaniclastic turbidite facies) was relatively good at Site C0012. The lithology consists of silicic tuff, volcaniclastic sandstone, siliciclastic sandstone, and hemipelagic mudstone (Fig. F7). There is a significant unconformity near the base of the unit, and the maximum age is >18.9 Ma. Mudstone interbeds within Unit V contain between 38% and 97% smectite in the clay-size fraction (average = 74%). Illite content averages 18%, chlorite + kaolinite averages 3%, and quartz averages 5%. This facies includes 11 identifiable ben-

tonites, and Figure F8 shows that smectite abundance in bulk mudstone ranges from 24% to 82% (average = 50%).

Unit IV is ~10 m thick and consists of reddish brown pelagic claystone with high concentrations of calcium carbonate. Samples from that unit were not analyzed for their clay-size mineral assemblages. Igneous basement (basalt) was also omitted from our XRD study, but information on its composition and alteration can be found in Kameda et al. (2011).

Figure F9 demonstrates that the expandability of I/S mixed-layer clay minerals does not change significantly from the top of Unit I to the base of Unit V. The proportion of illite in the I/S averages 12% and ranges from 3% to 48%. The highest values of illite occur near the base of the sedimentary section, but there is no consistent pattern of progressive smectite-to-illite diagenesis as a function of depth. The estimated temperature at the top of basement is ~65°C (Expedition 333 Scientists, 2011). Values of illite crystallinity index typically range between 0.42 and 0.25 Δ<sup>20</sup>, which is similar to what was detected at Site C0011. When compared to Site C0011, however, the results from Site C0012 display more scatter in illite crystallinity and a larger number of values fall within the diagenetic zone. This is likewise consistent with erosion of the illite as a detrital constituent from sedimentary source terrains that had been exposed to anchizone conditions of incipient metamorphism (e.g., Blenkinsop, 1988; Kisch, 1990).

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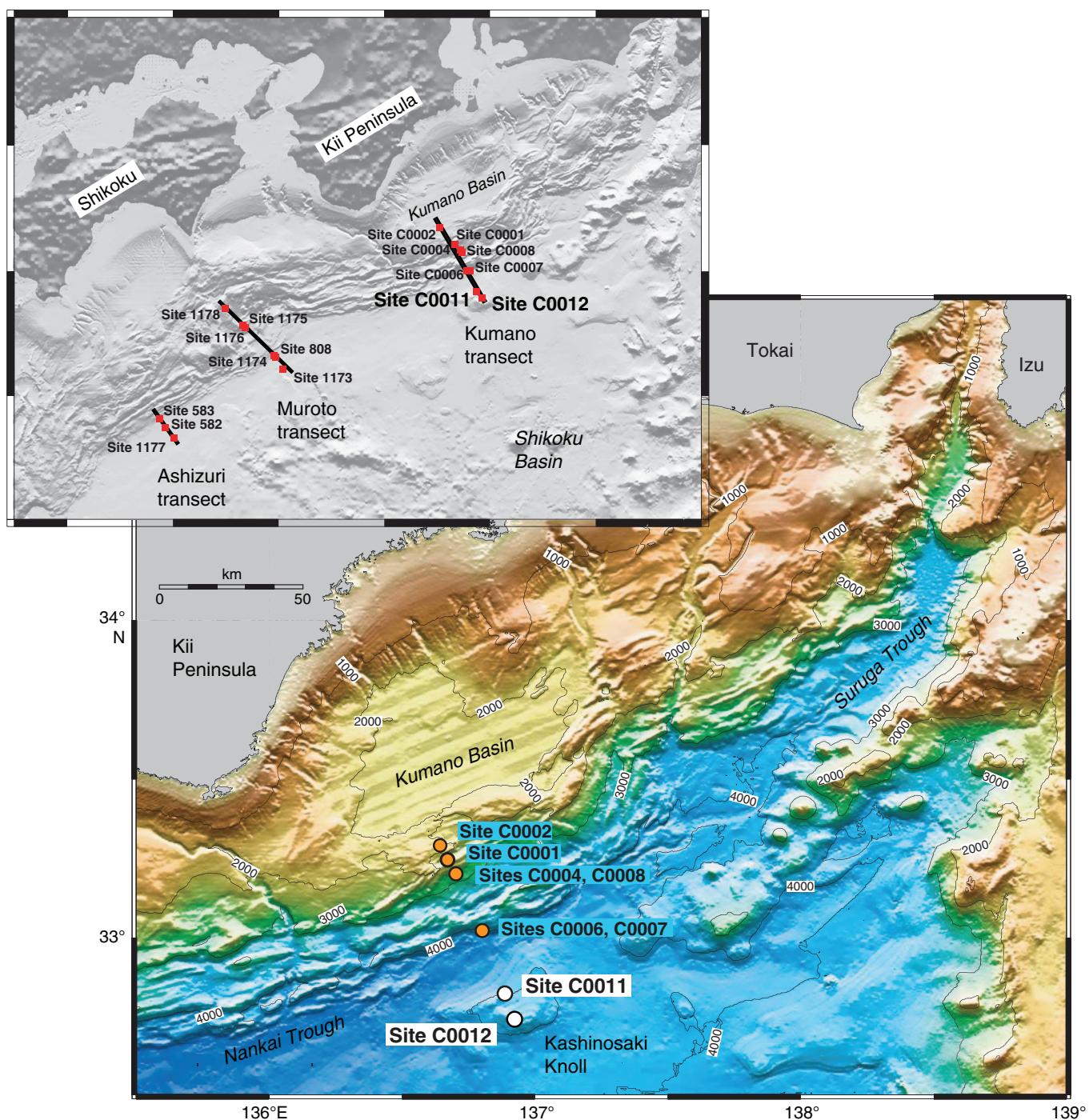
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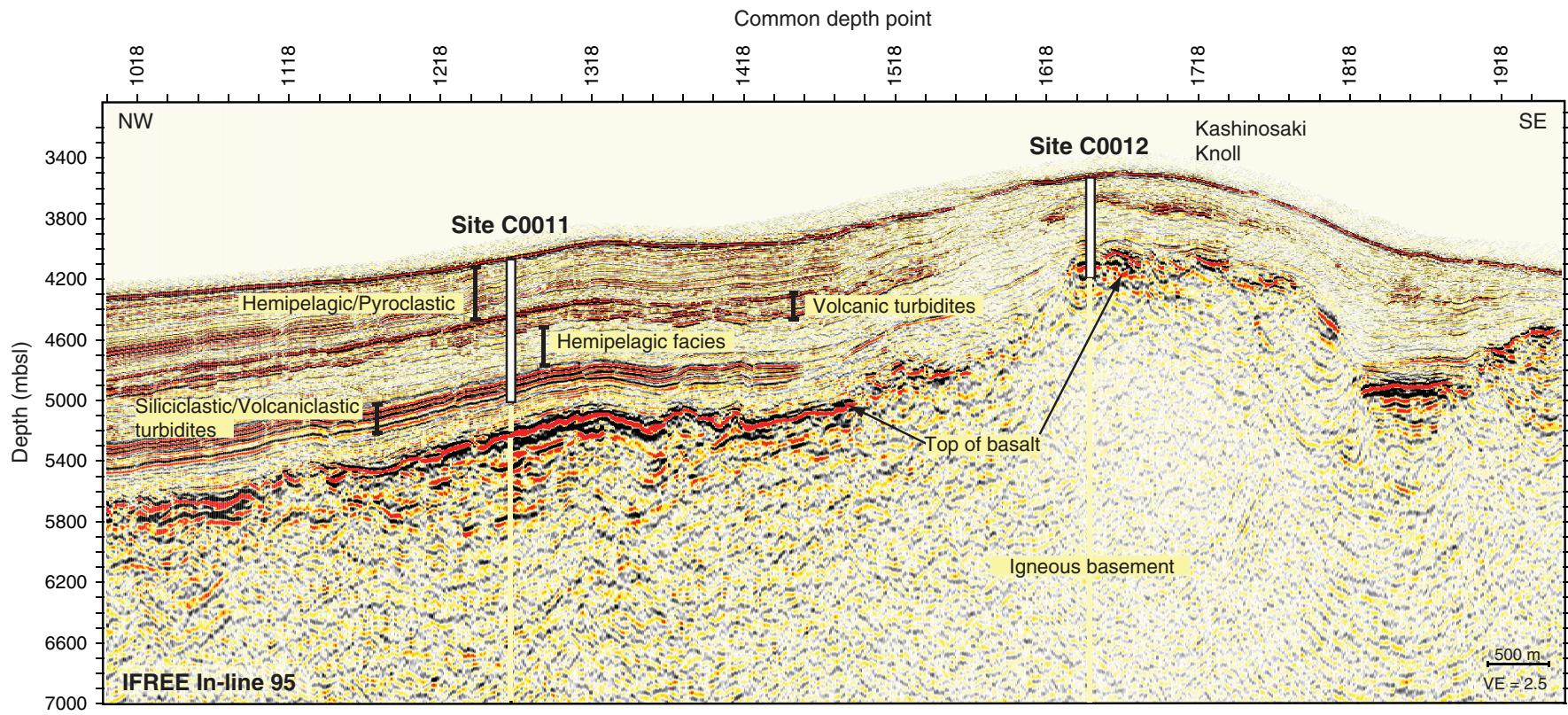
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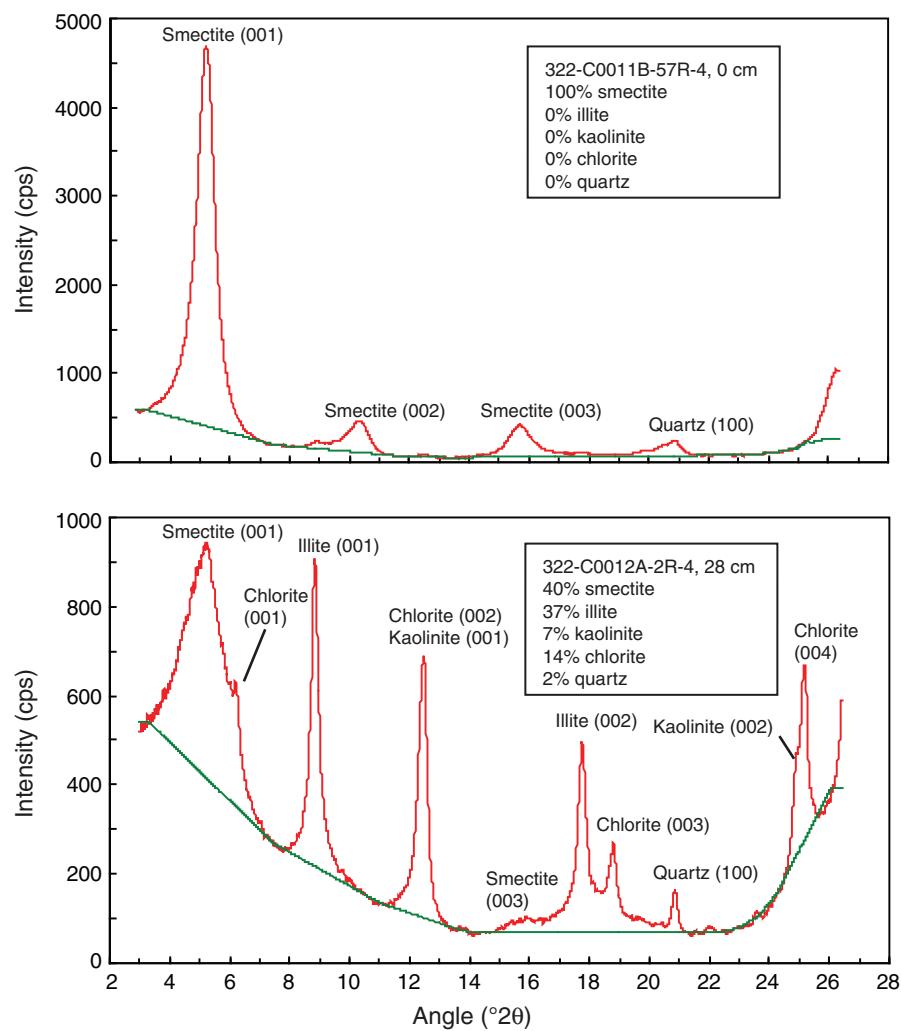
**Figure F1.** Map of the Nankai Trough and Shikoku Basin study area with IODP NanTroSEIZE drill site locations. Also shown are locations of DSDP and ODP transects and drill sites.



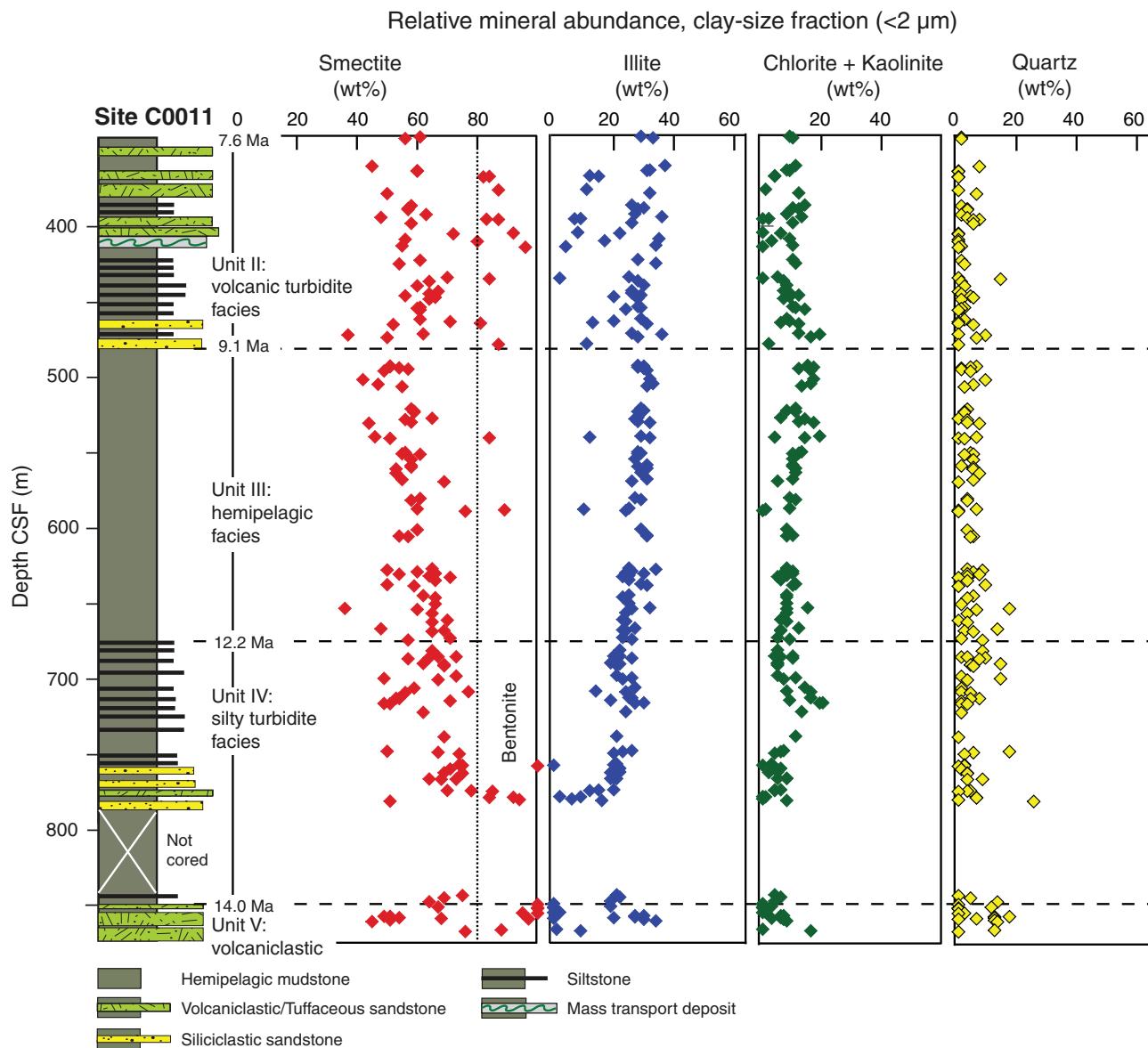
**Figure F2.** Seismic in-line section crossing the Kashinosaki Knoll showing locations of Sites C0011 and C0012 and interpretation of acoustic units based on core recovery. IFREE = Institute for Research on Earth Evolution. VE = vertical exaggeration.



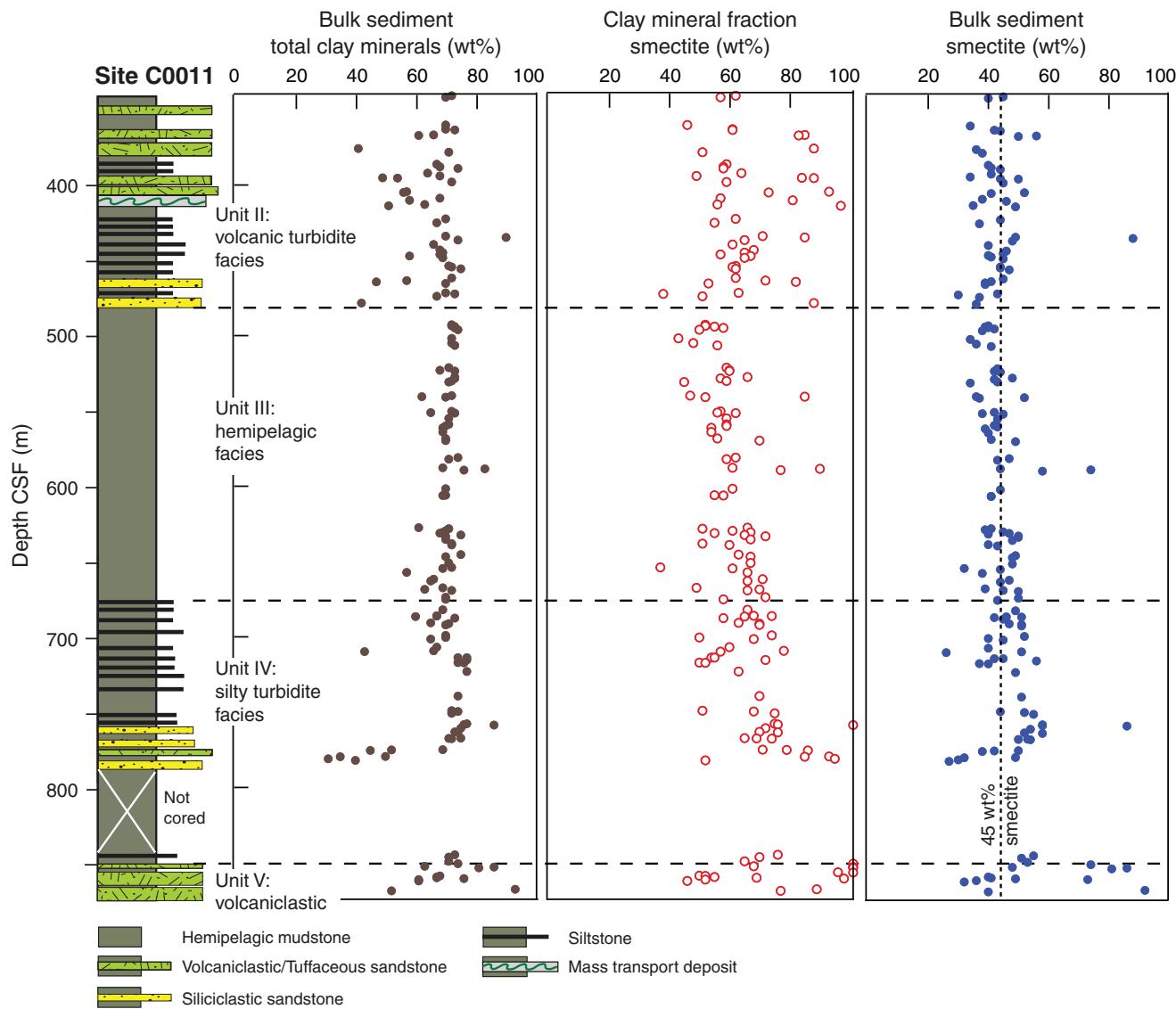
**Figure F3.** Representative examples of X-ray diffractograms showing peaks for smectite, illite, chlorite, kaolinite, and quartz, Sites C0011 and C0012. Values of relative mineral abundance (weight percent) were calculated using the SVD normalization factors of Underwood et al. (2003) and the following peak areas: smectite (001), illite (001), chlorite (002) + kaolinite (001), and quartz (100). Proportion of kaolinite to chlorite was calculated using the equations of Guo and Underwood (2011).



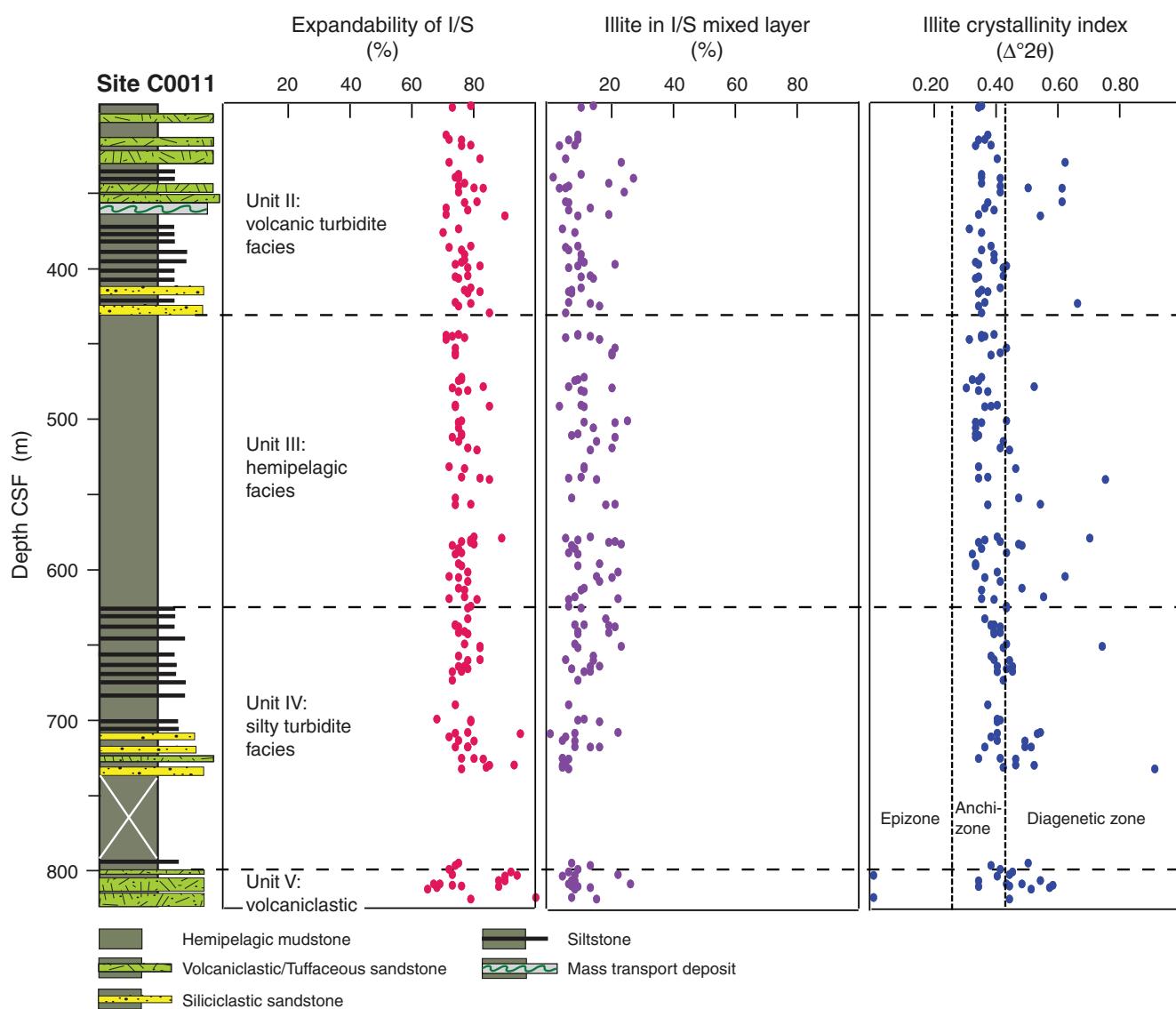
**Figure F4.** Lithostratigraphy of Site C0011 with relative abundances of smectite, illite, chlorite, kaolinite, and quartz in clay-size fractions of mud(stone). Ages at the boundaries are based on an integrated age-depth model (see the “Site C0011” chapter [Expedition 322 Scientists, 2010a]). Calculations were made using SVD normalization factors.



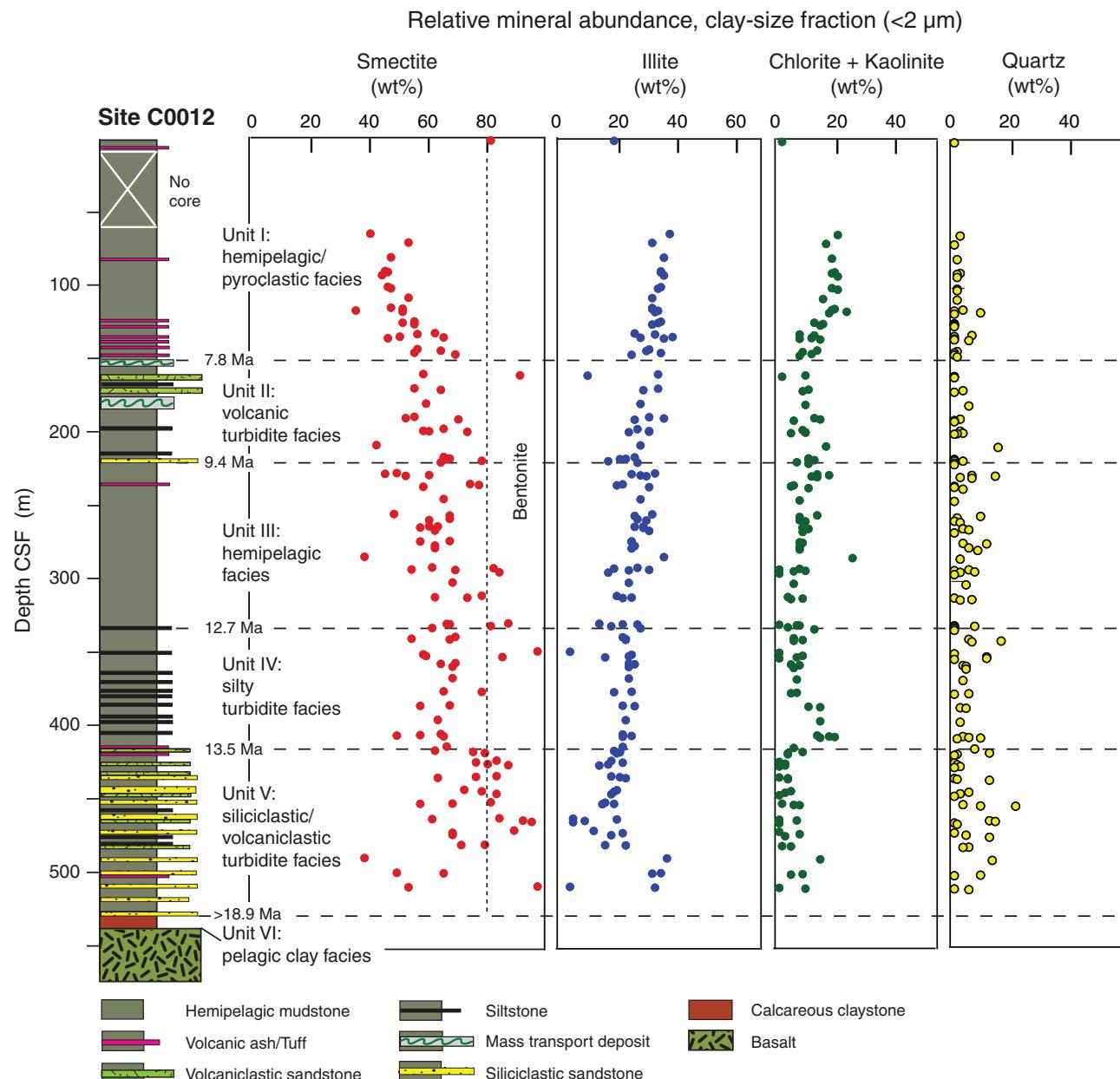
**Figure F5.** Lithostratigraphy of Site C0011 with relative abundances of total clay minerals in bulk mudstone relative to quartz + feldspar + calcite (see the “[Site C0011](#)” chapter [Expedition 322 Scientists, 2010a]), smectite in the clay mineral fraction (relative to illite + chlorite + kaolinite), and smectite in bulk mudstone. Calculations were made using SVD normalization factors.



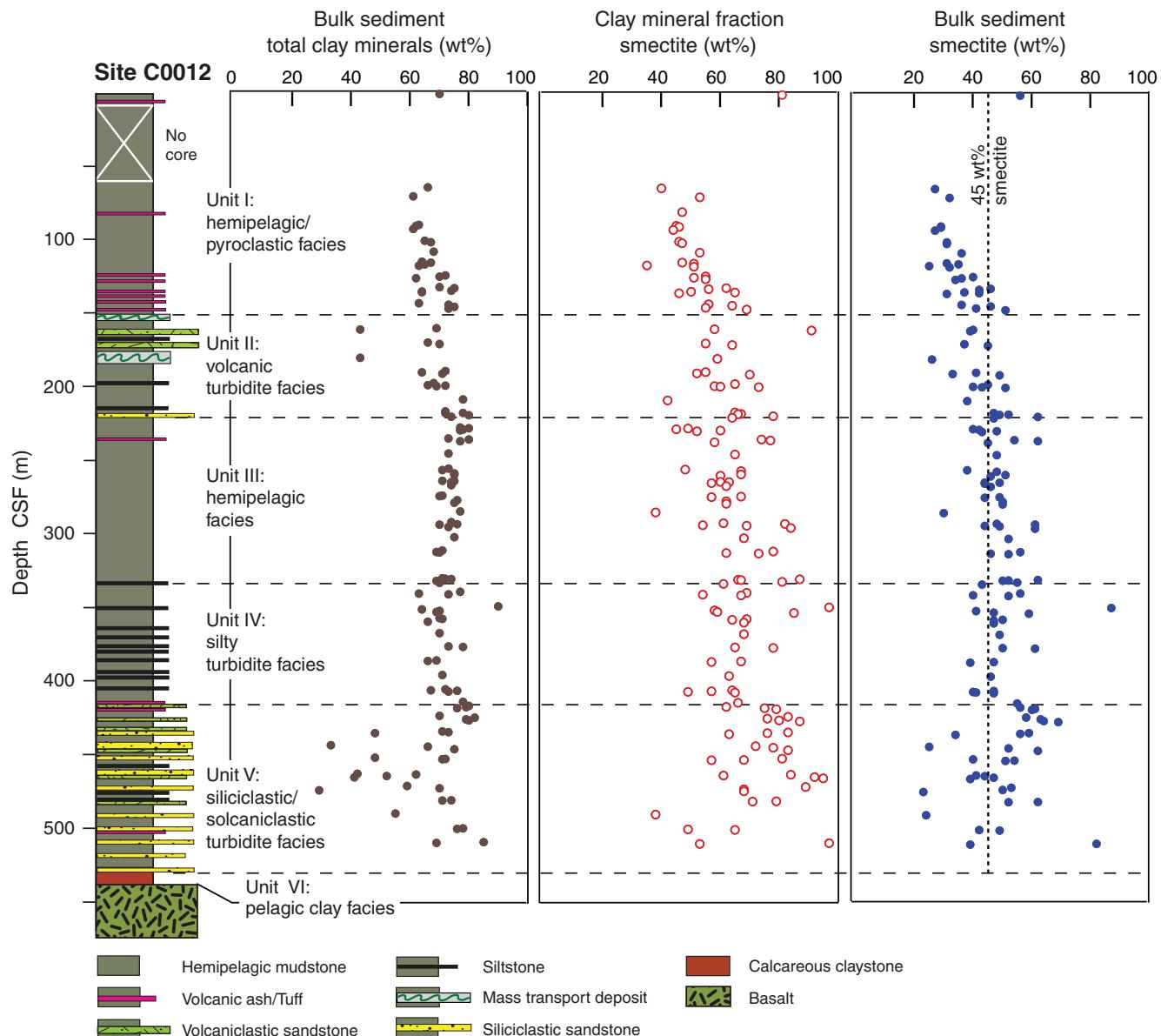
**Figure F6.** Lithostratigraphy of Site C0011 with values of illite/smectite (I/S) mixed-layer clay expandability, proportion of illite in I/S mixed-layer clay, and illite crystallinity index.



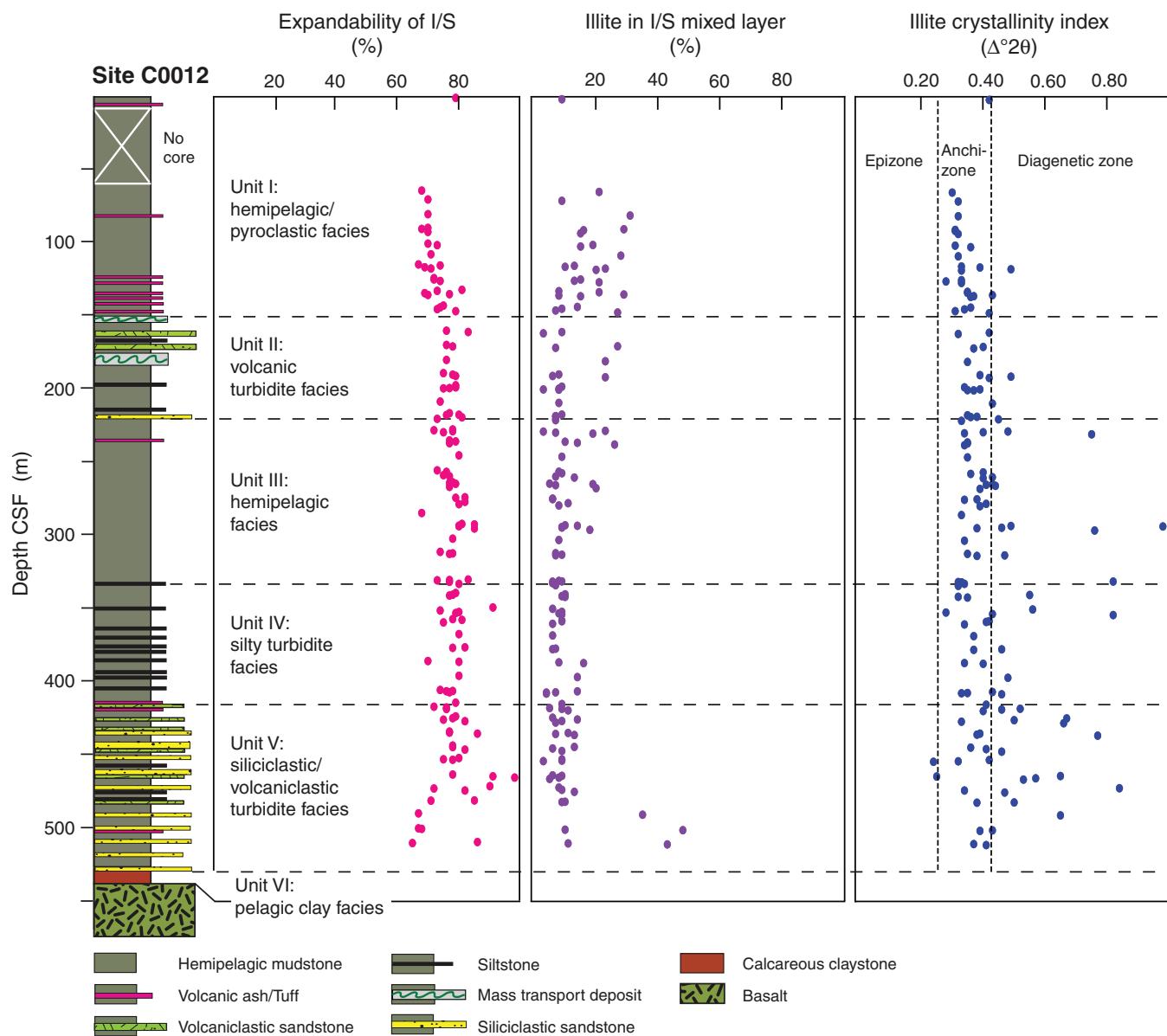
**Figure F7.** Lithostratigraphy of Site C0012 with relative abundances of smectite, illite, chlorite, kaolinite, and quartz in clay-size fractions of mud(stone). Ages at unit boundaries are based on an integrated age-depth model (see the “Site C0012” chapter [Expedition 322 Scientists, 2010b]). Calculations were made using SVD normalization factors.



**Figure F8.** Lithostratigraphy of Site C0012 with relative abundances of total clay minerals in bulk mudstone relative to quartz + feldspar + calcite (see the “Site C0012” chapter [Expedition 322 Scientists, 2010b]), smectite in the clay mineral fraction (relative to illite + chlorite + kaolinite), and smectite in bulk mudstone. Calculations were made using SVD normalization factors.



**Figure F9.** Lithostratigraphy of Site C0012 with values of illite/smectite (I/S) mixed-layer clay expandability, proportion of illite in I/S mixed-layer clay, and illite crystallinity index.



**Table T1.** Singular value decomposition normalization factors (from Underwood et al., 2003) used to calculate relative mineral abundances in clay-size aggregates (<2 µm size fraction).

<2 µm fraction	Affected mineral standard mixture			
	Smectite	Illite	Chlorite	Quartz
<b>Influencing mineral:</b>				
Smectite	3.7398559E-04	-2.8994615E-05	-3.4377535E-05	-7.4421238E-05
Illite	4.2720105E-05	1.2499784E-03	-2.836388E-05	3.3838456E-05
Chlorite	-6.7662186E-05	-2.008419E-07	7.6974847E-04	5.240881E-05
Quartz	2.4368789E-03	9.2311541E-04	8.195109E-04	3.7061975E-03











**Table T2 (continued).**

Core, section, interval (cm)	Depth CSF (m)	Integrated peak area (total counts)					
		Smectite (001)	Illite (001)	Chlorite (002) + kaolinite (001)	Quartz (100)	Half peak chlorite (004)	Kaolinite (002) + chlorite (004)
37R-2, 0	386.30	128,247	20,278	20,269	3,226	5,651	14,169
38R-1, 95	395.95	181,501	22,271	26,263	3,572	6,362	17,145
39R-1, 118	405.68	156,991	18,203	21,974	3,390	4,480	14,419
39R-2, 37	406.23	134,398	17,715	24,660	3,526	5,896	15,101
39R-2, 73	406.59	98,657	17,004	23,726	3,627	6,600	15,286
39R-3, 35	407.25	214,283	24,656	31,539	3,857	7,310	18,269
40R-1, 21	414.21	144,437	17,005	9,054	4,564	1,987	4,937
Unit V - siliciclastic/volcaniclastic turbidite facies							
40R-3, 26	416.98	45,502	4,578	3,984	1,956	1,452	3,063
40R-4, 30	417.71	137,445	13,756	6,340	2,854	1,441	3,468
40R-5, 45	418.55	204,764	18,266	9,883	3,344	2,624	5,637
41R-1, 23	423.73	178,663	14,848	5,053	1,762	915	2,802
41R-2, 0	424.88	130,852	13,344	5,421	2,601	1,092	2,761
41R-3, 0	426.07	154,714	12,002	5,485	3,308	1,506	3,535
41R-4, 38	426.89	293,703	18,570	5,508	3,463	1,521	3,469
42R-2, 29	434.24	244,518	19,600	5,374	3,988	1,284	2,750
42R-3, 18	434.84	167,010	16,342	8,674	3,225	2,682	5,621
42R-3, 76	435.42	148,784	18,294	5,095	6,336		
43R-1, 114	443.64	143,535	13,527	7,597	3,873	2,084	4,746
43R-2, 85	444.75	176,114	15,191	6,991	3,895	1,595	3,865
43R-4, 41	446.40	299,999	24,185	9,795	4,613	2,366	5,100
44R-1, 16	452.16	230,179	16,253	6,520	5,356	1,510	3,245
44R-1, 94	452.94	134,345	13,089	7,253	4,624	2,063	4,729
44R-2, 0	453.26	74,335	6,023	4,421	5,502	1,033	2,305
45R-2, 22	463.03	1,611,840	17,210	3,457	61,145		
45R-2, 74	463.55	119,016	12,815	7,580	5,688	1,272	3,208
45R-3, 46	464.55	531,869	22,074	6,446	8,912	883	2,244
45R-4, 76	465.55	749,785	15,059	4,914	15,941		
46R-2, 0	471.36	596,795	30,980	7,581	10,500		
46R-3, 109	473.02	123,286	13,577	9,478	3,157	2,271	5,325
47R-2, 30	474.45	73,364	6,590	1,940	2,987	480	1,438
47R-3, 10	481.07	109,248	7,579	2,757	2,879	447	1,123
48R-1, 10	481.22	123,017	14,033	6,763	2,895	1,727	3,856
48R-2, 0	490.10	18,408	6,803	4,421	1,111	1,314	3,236
49R-2, 37	500.30	62,062	15,713	6,835	2,482	2,029	4,961
49R-3, 22	500.54	119,816	20,638	7,955	1,935	1,849	4,164
50R-2, 40	509.69	420,922	11,231	5,156	4,125	1,188	2,762
50R-3, 50	510.21	58,677	12,888	7,164	1,689	1,973	4,168























