Data report: sediment major element, minor element, and reactive iron and manganese data from the Okinawa Trough: IODP Expedition 331 Sites C0014 and C0017¹

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Abstract

We present sediment solid-phase chemical data from Integrated Ocean Drilling Program Expedition 331 Sites C0014 and C0017, two sites with contrasting sediment chemistries. Results presented here include a suite of major and trace elements from whole-sediment digestions and sediments extracted using a selective leaching procedure for Fe and Mn. The sediment package at Site C0014 is chemically variable, displaying a wide range of concentrations for a number of major and trace elements. This sediment package has intervals rich in solid-phase Mg, with values as high as 10–15 wt% between ~30 and 38 meters below seafloor (mbsf) and lower values at shallower depths and deeper than ~40 mbsf. This pattern is consistent with other expedition data, suggesting a sedimentary zone of active high-temperature alteration centered roughly between 30 and 40 mbsf. The upper sedimentary package is also rich in dithionate extractable iron, which decreases to low values by ~30 mbsf. Unlike iron, dithionate extractable Mn is relatively low in the upper 20–30 m of sediment, with some enrichment in the uppermost sediment package for Holes C0014G and C0014B. Below this region, however, a zone of enriched reactive and total Mn is centered just above ~40 mbsf. In contrast to Site C0014, Site C0017 generally lacks the more pronounced chemical signatures apparent within the sediments, and many of the chemical signatures likely reflect the regional background hemipelagic sediment cover. There are, however, elevated levels of reactive Mn both in the upper \sim 30 m and the bottom \sim 50–75 m of sediment.

Introduction

During Integrated Ocean Drilling Program (IODP) Expedition 331, several sites were drilled within the Okinawa Trough region where the Iheya North hydrothermal field is located. This report focuses on two sites located east of the main hydrothermal vent field, which is an active hydrothermal venting region (Takai et al., 2011; see the "Expedition 331 summary" chapter [Expedition 331 Scientists, 2011a]).

Site C0014 is located ~450 m east of the main vent field and is a dynamic environment with pore fluid chemistry suggestive of a hemipelagic sediment cover overlying a diagenetically and hydro-thermally altered sediment package containing volcanic sediments, hydrothermally altered muds, and pumiceous sediments

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(Takai et al., 2011). Temperatures increase from bottom water values to ~145° at 47 mbsf with an abrupt increase to >200°C at 50 mbsf (see the "Site C0014" chapter [Expedition 331 Scientists, 2011b]). For Hole C0014G, the hole with the most continuous sampling through the sediment column and the hole for which we have the most complete data set, the upper ~20 m is rich in dissolved sulfate with a decrease to near zero at ~30 mbsf (Takai et al., 2011). Within this upper, sulfate-rich package, dissolved Mg is close to the seawater value with a decrease to values <5 mM by ~40 mbsf (Takai et al., 2011).

Site C0017 is located farther east, roughly 1.6 km from the Iheya hydrothermal vent field. This site has the deepest penetration of the holes drilled during the expedition. Similar to Site C0014, Site C0017 is primarily a site of mixed sediment types with thick hemipelagic mud and various volcaniclastic sediment types. Temperature measurements indicate that the sediment column reaches ~90°C by ~150.7 mbsf (see the "Site C0017" chapter [Expedition 331 Scientists, 2011c]). This site, however, lacks the dramatic and larger scale hydrothermal alteration seen at Site C0014. Consistent with this contention, the pore fluid chemistry of the uppermost 25-35 m is dominated by more typical hemipelagic diagenetic reactions (Takai et al., 2011). Below this depth, lowtemperature reactions continue to dominate the pore fluid chemistry (Takai et al., 2011).

For both of these sites, lateral fluid flow has been inferred (Takai et al., 2011). From the perspective of this report, it is important to note that Site C0014 exhibits clear evidence for active high-temperature alteration, whereas Site C0017 does not. Both of these systems also have physical and chemical features that imply a dynamic and perhaps nonsteadystate environment where any possible fluid flow is likely to be three dimensional (Takai et al., 2011).

Analytical methods Total digestion methods

The details of our sediment digestion methods are outlined in Muratli et al. (2012). Briefly, dried and ground sediment samples were digested using a combination of inorganic acids (HCl, HNO₃, and HF) in a CEM MARS-5 microwave oven (CEM Corp., Matthews, NC). Postdigestion evaporation within the system was accomplished utilizing the MicroVap accessory. Following sample dilution with 5% HNO₃ into preweighed 10 mL HDPE bottles (this dilution is somewhat modified from Muratli et al., 2010a, 2010b, 2012), samples were heated in an analog block heater for ~24 h. Prior work identified the necessity of this step to redissolve any remnant fluoride-metal complexes that, although they might not be visually apparent in the digestion matrix, were found to impact analytical results (e.g., Muratli et al., 2010b, 2012).

ICP-OES

Samples for major elements (Tables T1, T2) were run on a Leeman Laboratories Prodigy inductively coupled plasma-optical emission spectrometer (ICP-OES) at the W.M. Keck Collaboratory at Oregon State University (USA). This instrument is capable of two viewing modes, axial and radial, and the various elements were run in one mode or the other. Table T1 also presents the specific emission wavelengths utilized for each element. Samples and standards were diluted twenty-fold with 1% quartz-distilled nitric acid. Repeated runs of the standard curve monitored instrumental drift, and for those elements influenced by drift we applied a correction factor. The reported data represent the mean of three replicate analyses for each element in each tube except in instances where we have results from multiple sample digestions. In this case, the mean is the average of the two results. Uncertainties for samples that were not digested in duplicate are derived from two sources: (1) the regression uncertainty, which is calculated using the standard error of the regression, and (2) the internal uncertainty calculated from the standard deviation of the three replicate analyses. These two uncertainties are combined as the square root of the sum of squares.

ICP-MS

Trace constituents were analyzed on a Thermo X-Series II inductively coupled plasma-mass spectrometer (ICP-MS), also at the W.M. Keck Collaboratory. IPC-MS results can drift over the course of the day's run, and drift is corrected by spiking each sample with an internal standard solution. This solution consisted of 250 ng/mL Be and 50 ng/mL each of Rh and Bi or just 50 ng/mL of Rh and Bi. Measured counts throughout the run were corrected to the value of the first blank run of the day. Because of the varied matrix in these samples, Rh was used as the lone internal reference standard for the calculation of results. Although using only a single element for drift correction is less than optimal, samples had varying amounts of Be and Bi, thus making reliable concentration results difficult to obtain. Sample dilutions varied with concentration, but samples were diluted with 1% HNO₃. Sample uncertainties were calculated using the same approach employed for the ICP-OES results.



We used the following isotope masses to calculate concentrations: ⁶⁶Zn, ⁹⁸Mo, ¹¹¹Cd, ¹¹⁴Cd, ¹³³Cs, ¹⁴⁶Nd, ¹⁸⁵Re, ¹⁸⁷Re, and ²³⁸U. Multiple analyte masses have interferences; thus, some of the analytes above were used to estimate these interferences. For ¹¹⁴Cd we measured ¹¹⁷Sn to assess the level of interference of ¹¹⁴Sn:

$$^{114}Cd_{corr} = (^{114}Cd_{meas}) - 0.0859375 \times ^{117}Sn.$$

An additional interference from ¹¹⁴MoO on ¹¹⁴Cd is estimated from the Mo results and the estimated MoO/Mo formation ratio. For this work, we only report results from the ¹¹⁴Cd analyses because we are more confident in the correction for this Cd mass. For ¹⁸⁷Re, a series of solutions containing Yb were analyzed and a YbO/Yb ratio was calculated. Counts of ¹⁷¹Yb were then used to estimate the interference on ¹⁸⁷Re. ¹⁸⁵Re has an interference by TmO; for this reason, we generally use ¹⁸⁷Re to calculate Re concentrations. However, at higher Re concentrations (and higher Re dilutions) this interference is small, and we use ¹⁸⁵Re for estimating concentrations.

Dithionite extractions of Fe and Mn

To determine "reactive" iron and manganese, we employed a single-step dithionite extraction at 60°C (Mehra and Jackson, 1960; Kostka and Luther, 1994; McManus et al., 2012; Roy et al., 2013). We used approximately 0.25 g of dried ground sediment to which we added ~10 mL of dithionite reagent. pH is buffered using a sodium acetate, sodium citrate solution (Roy et al., 2013). These extractions are generally thought to dissolve amorphous Fe oxides, some crystalline Fe oxides, and acid volatile sulfides (Kostka and Luther, 1994; Poulton and Canfield, 2005). However, for many of the samples digested here it is possible that the sediments are sufficiently reducing to the point where sulfides are the dominant reactive phase; thus, any future interpretations of the data presented here require caution regarding the specific phase being extracted using this procedure. Unlike Fe, this procedure is not well calibrated for Mn (McManus et al., 2012).

Quality control

To access accuracy and precision, we report values from replicate digestions of a PACS-2 standard and an in-house laboratory sediment standard (Tables **T2**, **T3**, **T4**). This latter standard is a continental margin sediment standard from the Chile margin (Muratli et al., 2010a, 2010b). We also report our own laboratory's long-term average for these materials (Tables **T2**, **T3**). Finally, we redigested/reanalyzed a number of samples, which are indicated within the data tables. Because the samples that we analyzed as part of this study are quite unlike either of the reference material matrixes, the reproducibility of our reference materials may not be an accurate measure of the true reproducibility of the samples. Generally speaking, however, the samples did indeed reproduce quite well, with one exception. As noted in Table T3, the Zn and Cd results from one duplicate sample (331-C0014B-3H-5, 20-30 cm) did not agree, and the results for these two analytes (for one of the duplicates) were higher than all other samples measured during this study. We did not obtain an accurate concentration estimate for these two analytes because the concentrations were above our highest standard and we did not reanalyze this sample. We report here a lower concentration for the duplicate, for which we have a more analytically robust measure of these two analytes. We suspect that either the sample was somehow contaminated or it was not homogeneous. We do note, however, that other elements do show agreement between these two duplicates.

In the case of our Fe and Mn extractions, we also extracted these same sediment standards and report the data in Table T4. These values are not standardized and can thus only be used as a measure of technique precision. Furthermore, many of the reported results are the result of duplicate sediment extractions and analyses, and for these samples the average values are reported (\pm 1 standard deviation).

Results

The solid phase Mg from Site C0014 (Fig. F1A) is highly variable, with low values in the uppermost sediment package and higher values between 30 and 40 mbsf. We show here the dissolved Mg profiles from the cores in which we have corresponding solid-phase analyses and which show the expected depletion of Mg relative to the seawater value within the zone of high Mg concentrations (Fig. F1B). The Ca/Al ratios generally show lower values than those expected for average upper continental crust, with the exception of the uppermost samples (Fig. F2A). The K/Al ratios are quite variable, with low values near the zone of high Mg/Al and at slightly shallower depths for some of the samples (Fig. F2B, F2C). Generally speaking, the upper sediment package from core to core can be variable, and this variability was noted for the pore fluid results from this site (Takai et al., 2011). Below ~40 mbsf, K/Al values are slightly enriched relative to average upper continental crust (Fig. F2B). Ba/Al is also generally lower than upper continental crust values, with the exception of



the uppermost sediment package and at ~12–15 mbsf (Fig. F2B). Fe/Al is generally enriched in the uppermost 40 mbsf, whereas Fe_R/Fe_T enrichments are generally confined to the upper 20 mbsf (Fig. F3A, F3B). Below these depths, Fe/Al is depleted relative to average upper crust values and the Fe_R/Fe_T ratio approaches zero, meaning that very little of the iron is extractable at these depths. The upper sediment package is depleted in Mn, but a broad elevated Mn/Al maximum is centered between ~30 and 50 mbsf (Fig. F3C, F3D). There is some suggestion of elevated Mn_R/Mn_T in this region, particularly with respect to the upper sediment package; however, much of the data set suggests Mn_R concentrations are ~10% or less of the total Mn.

In contrast to the data from Site C0014, the data from Site C0017 has much smaller variability in the sedimentary Mg contents, with few samples significantly enriched over the value for the upper continental crust (Fig. F4A). This observation is consistent with the dissolved profile, which exhibits little deviation from the ambient seawater value (Fig. F4B). Ca/Al ratios in the upper ~40 m are elevated over the background, and deeper samples have values that hover near the upper continental crust ratio (Fig. F5A). The K to Al ratios are relatively constant, as are the Ba to Al ratios (Fig. F5B, F5C). Fe to Al ratios are generally slightly enriched relative to the upper continental crust value but vary little with depth (Fig. **F6A**). The Fe_R to Fe_T ratios are higher on average relative to Site C0014, but no particular pattern is apparent in the data (Fig. F6B). Mn to Al ratios also do not vary significantly compared to Site C0014 and generally scatter about the crustal ratio, with the average value slightly below that ratio (Fig. F6C). The Mn_{R} to Mn_T ratio has a distinct minimum below 50 mbsf with values increasing with depth (Fig. F6D).

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Figure F1. Plots of (A) Mg/Al and (B) pore fluid dissolved Mg, Holes C0014B, C0014D, and C0014G. Dashed lines = (A) upper continental crust ratio (Rudnick and Gao, 2004), (B) seawater value.





Figure F2. Plots of (A) Ca/Al, (B) K/Al, and (C) Ba/Al, Holes C0014B, C0014D, and C0014G. Dashed lines = upper continental crust ratio (Rudnick and Gao, 2004).





Figure F3. Plots of (A) Fe/Al, (B) Fe_R/Fe_T , (C) Mn/Al, and (D) Mn_R/Mn_T , Holes C0014B, C0014D, and C0014G. Dashed lines = upper continental crust ratio (Rudnick and Gao, 2004).





Figure F4. Plots of (A) Mg/Al and (B) pore fluid dissolved Mg, Holes C0017A–C0017D. Dashed lines = (A) upper continental crust ratio (Rudnick and Gao, 2004), (B) seawater value.





Figure F5. Plots of (A) Ca/Al, (B) K/Al, and (C) Ba/Al, Holes C0017A–C0017D. Dashed lines = upper continental crust ratio (Rudnick and Gao, 2004).





Figure F6. Plots of (A) Fe/Al, (B) Fe_R/Fe_T , (C) Mn/Al, and (D) Mn_R/Mn_T , Holes C0017A–C0017D. Dashed lines = upper continental crust ratio (Rudnick and Gao, 2004).





Table T1. ICP-OES run specifications.

Element	Line (nm)	View
Al	396.152	Radial
Ва	455.403	Radial
Ca	317.933	Radial
Cu	324.754	Axial
Fe	259.940	Radial
К	766.491	Radial
Mg	279.078	Axial
Mn	257.610	Radial
Na	589.592	Radial
Sr	407.771	Radial
Ti	334.941	Radial
V	292.401	Axial
Zn	206.200	Axial
Zr	343.823	Axial

Table T2. ICP-OES sediment results, Expedition 331. This table is available in an oversized format.

Table T3. ICP-MS sediment results, Expedition 331. (Continued on next page.)

Core, section, interval (cm)	Depth (mbsf)	Zn (µg/g)	±	Mo (µg/g)	±	Cd (µg/g)	±	Cs (µg/g)	±	Nd (µg/g)	±	Re (ng/g)	±	U (µg/g)	±
331-C0014B-															
1H-1, 25–35*	0.25	222	9	3.7	0.6	0.24	0.01	6.4	0.1	22.4	0.4	2.6	0.4	3.9	0.2
1H-2, 100–110	2.32	184	1	27.7	0.3	0.78	0.02	12.7	0.1	32.4	0.9	16.2	0.8	8.5	0.5
1H-3, 92–105	3.54	108	1	2.10	0.04	0.29	0.02	12.0	0.1	30.2	0.9	6.6	0.7	4.6	0.5
1H-4, 100–110	4.95	110	1	1.21	0.03	0.28	0.01	17.6	0.1	29.5	0.8	7.8	0.7	5.5	0.5
1H-5, 65–75	5.91	101	1	0.93	0.02	0.31	0.02	20.0	0.1	28.0	0.9	8.7	0.9	5.5	0.5
2H-3, 10–20	8.46	106	1	140	2	3.58	0.03	8.5	0.1	21.6	0.9	55 [‡]	6	37.4	0.9
2H-7, 40–50	12.18	981 [†]	12	193	3			10.1	0.1	7.2	0.9	75 [‡]	5	247	25
2H-10, 20–30	14.26	790 [†]	13	289	3	3.99	0.05	12.3	0.1	7	1	195‡	6	51.5	0.9
3H-2, 52–62	17.21	42.7	0.8	28.4	0.3	0.12	0.02	12.0	0.1	119	1	11	1	27.0	0.4
3H-5, 20–30*	19.39	197	1	22.5	0.3	0.60	0.02	10.2	0.2	35	1	5.6	0.0	5.7	0.1
3H-7, 89–99	21.74	1003†	13	6.19	0.08	4.31	0.02	11.2	0.1	50.6	0.9	2.5	0.7	9.8	0.5
3H-9, 67.5–77.5	23.52	139	1	4.66	0.02	0.48	0.02	14.5	0.1	42.6	0.9	4	1		
331-C0014D-															
1H-1, 2–5	0.02	186	1	76.6	0.7	4.43	0.03	8.8	0.1	22.3	0.9	96‡	4	17.6	0.5
1H-2, 108–110*	2.48	122	2	130.4	0.9			4.5	0.1	11.1	0.2	155‡	3	26.1	0.6
1H-3, 80–82	3.61	78	1	4.85	0.07	0.07	0.02	15.8	0.1	35.6	0.9	3.0	0.7	88	1
1H-4, 80–82	5.03	74	1	11.2	0.1	0.18	0.02	1.9	0.1	29	1				
2H-1, 30–32	6.80	96	1	170	2	4.17	0.04	3.4	0.1	8.1	0.9	114‡	7	29	1
2H-2, 30–32	8.18	381†	13	810	12			2.8	0.1	26.1	0.9	536‡	15	80	1
2H-3, 37–39	9.66	330†	13	546	2	2.97	0.10	3.4	0.1	5	1	433 [‡]	6	201	4
2H-6, 60–62	13.21	649†	13	39.9	0.3	0.64	0.02	18.6	0.1	23.8	0.9	22.5	0.8	19.0	0.4
331-C0014G-															
1H-1, 0–2	0.00	239	2	5.34	0.03	0.57	0.01	10.8	0.2	30.8	0.6	13.4	0.5	4.1	0.3
1H-2, 20–22	1.58	119	1	10.51	0.09	0.31	0.02	13.9	0.1	33.0	0.6	7.6	0.6	6.3	0.3
1H-5, 10–12	5.56	127	1	78.4	0.6	0.78	0.01	23.1	0.1	29.1	0.6	91	2	30.7	0.6
2H-5, 45–47	15.05	196	1	125.4	0.7	0.48	0.03	7.3	0.2	6.7	0.6	83	2	181	8
2H-7, 80–82	18.17	100	1	0.87	0.04	0.24	0.02	16.8	0.2	25.5	0.6	7.8	0.7	5.2	0.3
3H-2, 26–28	19.55	232	2	42.5	0.3	0.40	0.01	5.2	0.2	88.4	0.7	11	1	6.1	0.3
3H-5, 27–29	22.55	114	1	13.1	0.1	0.12	0.01	15.5	0.2	54.8	0.6	5.3	0.7	18.2	0.4
3H-8, 116–118	25.85	93	1	2.14	0.04	0.08	0.02	6.3	0.2	25.7	0.5	BDL		2.1	0.3
4H-5, 80–82*	31.21	50	1	13.2	0.2	BDL		2.3	0.1	52	1	6.2	0.5	5.3	0.0
4H-7, 70–72	33.13	95	1	5.55	0.05	0.09	0.02	2.3	0.2	21.2	0.5	2.6	0.8	4.9	0.3
4H-10, 60–62	36.05	321†	12	9.29	0.05	0.59	0.01	2.1	0.2	63.3	0.5	BDL		14.8	0.3
5H-3, 45–47	38.49	139	1	27.0	0.3	BDL		1.1	0.2	48.5	0.6	BDL		5.1	0.3
6H-2, 38–40	47.64	49	1	4.48	0.03	BDL		1.8	0.2	44.5	0.5	BDL		3.4	0.3
6H-3, 80–82	48.51	92	1	5.87	0.05	BDL		2.0	0.2	44.5	0.7	BDL		3.2	0.3
9X-2, 64–66	56.25	45	1	6.44	0.05	BDL		1.1	0.2	31.0	0.5	3.0	0.6	5.7	0.3
13T-1, 28–30	66.98	89	1	18.1	0.1	BDL		1.3	0.2	40.0	0.5	9.4	0.8	22.6	0.4
14T-2, 51–53*	72.19	64	2	24	1	BDL		1.4	0.0	40.0	0.4	8.5	0.3	19.1	0.1
16T-1, 7–9	75.97	79	1	10.2	0.2	BDL		1.5	0.2	35.1	0.7	5.9	0.7	8.2	0.3
17T-2, 20–22	80.81	78	1	4.16	0.05	BDL		1.9	0.2	84.6	0.6	5.4	0.7	8.3	0.3



Table T3 (continued).

Core, section, interval (cm)	Depth (mbsf)	Zn (µg/g)	±	Mo (µg/g)	±	Cd (µg/g)	±	Cs (µg/g)	±	Nd (µg/g)	±	Re (ng/g)	±	U (µg/g)	±
20T-1, 30–32	94.00	26	1	4.95	0.07	BDL		0.8	0.2	31.9	0.6	2.9	0.7	3.1	0.3
21H-3, 79-81	99.90	48	1	9.5	0.1	BDL		1.1	0.2	63	1	3.3	0.9	6.4	0.3
23X-1, 5–8	104.75	43	1	3.92	0.03	BDL		0.6	0.2	31.3	0.6	BDL		2.3	0.3
24T-2, 90–92	110.60	43	1	3.42	0.04	BDL		0.8	0.2	36.9	0.7	BDL		2.5	0.3
24T-4, 49.5–51.5	112.18	48	1	3.55	0.03	BDL		0.9	0.2	39.0	0.6	2.2	0.8	3.2	0.3
25T-1, 28.5–30.5	113.99	61	1	4.05	0.07	BDL		0.8	0.2	30.9	0.6	2.8	0.7	3.4	0.3
331-C0017A-															
1H-1, 80–90	0.80	104	1	1.02	0.03	0.32	0.01	11.1	0.1	28.7	0.9	4.6	0.5	4.1	0.5
1H-5, 88–98	6.52	108	1	0.53	0.03	0.35	0.01	10.4	0.1	28.8	0.9	11.5	0.6	5.0	0.5
331-C0017B-															
1H-2, 85–95	11.03	134	1	0.63	0.02	0.31	0.01	10.6	0.1	29.0	0.9	5.7	0.7	3.9	0.5
1H-5, 70–80	15.01	98	1	0.64	0.02	0.46	0.01	11.2	0.1	27.9	0.9	20	1	5.5	0.5
331-C0017C-															
1H-2, 60–70	20.26	81	1	1.66	0.02	0.22	0.01	6.7	0.1	28.9	0.9	5.1	0.5	3.3	0.5
1H-5, 115–125	24.87	72	1	1.27	0.02	0.60	0.01	5.6	0.1	21.0	0.9	25.6	0.9	7.9	0.5
2H-1, 65-75	28.45	70	1	4.06	0.04	0.13	0.01	3.1	0.1	29.3	0.9	BDL		2.4	0.5
2H-2, 65–75	29.85	73	1	0.28	0.02	0.12	0.01	4.6	0.1	24.0	0.9	BDL		1.9	0.5
331-C0017D-															
1H-3, 80–90*	63.55	117	0	1.84	0.16	0.19	0.00	3.2	0.1	34.4	1.1	BDL		1.8	0.1
1H-5, 80–90	66.30	116	1	2.13	0.03	0.18	0.01	3.3	0.2	34.8	0.6	BDL		1.7	0.3
1H-6, 110–120	67.98	73	1	0.51	0.03	BDL		6.7	0.2	30.5	0.6	BDL		2.0	0.3
2H-5, 30–40	75.08	111	1	0.53	0.03	0.10	0.01	8.9	0.2	28.0	0.6	7.9	0.8	3.5	0.3
6X-1, 66–76	94.86	75	1	0.75	0.03	0.09	0.01	8.4	0.2	30.6	0.6	BDL		2.4	0.3
7H-2, 112–122	106.16	92	1	1.18	0.03	0.19	0.01	10.3	0.2	31.6	0.6	BDL		2.8	0.3
7H-4, 42–52	108.12	96	1	2.20	0.04	0.15	0.01	11.3	0.2	31.3	0.7	2.0	0.5	2.8	0.3
7H-5, 80–90*	109.85	55	1	1.57	0.20	0.10	0.00	5.4	0.1	28.0	1.9	2.8	0.6	1.6	0.2
9X-5, 25–35	125.72	83	1	0.91	0.03	0.15	0.01	9.1	0.2	35.1	0.6	2.1	0.6	2.5	0.3
9X-8, 75–90	130.04	83	1	1.39	0.04	0.12	0.01	9.5	0.2	33.6	0.6	BDL		2.8	0.3
10X-1, 88–102	131.58	94	1	1.23	0.03	0.15	0.01	12.2	0.2	35.4	0.6	2.3	0.6	3.1	0.3
10X-4, 80–90	135.95	104	1	0.87	0.03	0.16	0.01	13.8	0.2	38.9	0.6	BDL		3.2	0.3
11X-1, 80–90	141.00	80	1	1.16	0.03	0.15	0.01	8.5	0.2	33.0	0.6	BDL		2.3	0.3
Reference materials															
PACS-2		369	10	5.7	0.2	2.16	0.06	2.40	0.07	18.5	0.4	6.1	0.3	2.7	0.1
Reference value		364	12	5.4	0.1	2.11	0.08							(3)	
Long-term average		395	31	6.8	1.9	2.14	0.08					6.3	0.4	2.5	0.1
Laboratory standard		84	1	1.8	0.1	0.45	0.01	5.24	0.07	19.4	0.4	14.5	0.6	3.4	0.2
Long-term average	2	89	7	1.9	0.1	0.45	0.01					14.7	0.6	3.4	0.1
Detection limit (appr	oximate)					0.045						2.0			

* = digested in duplicate. \dagger = analyzed using ICP-OES; Zn reference material results are for ICP-OES. \ddagger = analyzed at a larger dilution factor than others; therefore, no oxide correction was applied. For Sample 331-C0014B-3H-5, 20–30 cm, we disregarded a high Zn and Cd value from one duplicate digestion. These samples were not rerun to ascertain accurate concentrations, but in each case they would have had the highest concentrations measured during this study (Zn ~ 3800 µg/g and Cd ~ 10 µg/g). For Sample 331-C0017D-7H-5, 80–90 cm, one Re value was below the detection limit (BDL) and one was just above; we report the value above the detection limit.



Table T4. Reactive iron and manganese results, Expedition 331. (Continued on next page.)

Core, section,	Depth	Reactive Fe		Reactive Mn		Fe		Mn		Mn			
interval (cm)	(mbsf)	(%)	±	(%)	±	(%)	±	(µg/g)	±	(mg/g)	±	Fe _R /Fe _T	Mn _R /Mn _T
331-C0014B-													
1H-1, 25–35	0.25	0.45	0.05	0.0096	0.0008	3.0	0.1	418	10	0.42	0.01	0.150	0.23
1H-2, 100–110	2.32	0.32		0.0089		4.44	0.03	492	9	0.492	0.009	0.071	0.18
1H-3, 92–105	3.54	0.29		0.0113		3.75	0.03	491	10	0.49	0.01	0.077	0.23
1H-4, 100–110	4.95	0.26		0.0102		3.61	0.02	411	8	0.411	0.008	0.072	0.25
1H-5, 65–75	5.91	0.29		0.0106		3.39	0.02	424	9	0.424	0.009	0.085	0.25
2H-3, 10–20	8.46	0.29		0.0030		2.37	0.02	343	9	0.343	0.009	0.121	0.09
2H-7, 40–50	12.18	0.06		0.0022		2.73	0.03	167	9	0.167	0.009	0.023	0.13
2H-10, 20–30	14.26	0.14	0.00	0.0034	0 0000	8.86	0.03	1051	10	1.05	0.01	0.016	0.03
3H-2, 52–62	17.21	0.10	0.00	0.0051	0.0003	5.96	0.03	339	9	0.339	0.009	0.018	0.15
3H-5, 20-30	19.39	0.04	0.00	0.0045	0.0001	3.85	0.01	184	1	0.184	0.001	0.011	0.25
3H-7, 69-99 3H 0 67 5 77 5	21.74	0.07		0.0033		5.09 1 28	0.02	208	0	0.208	0.008	0.022	0.16
30-9, 07.3-77.3	23.32	0.19		0.0032		4.20	0.05	235	9	0.235	0.009	0.043	0.15
331-C0014D-													
1H-1, 2–5	0.02	0.29		0.0027		3.01	0.03	453	10	0.45	0.01	0.096	0.06
1H-2, 108–110	2.48	0.11		0.0011		1.41	0.02	166	2	0.166	0.002	0.080	0.06
1H-3, 80–82	3.61	0.17		0.0007		2.13	0.02	252	10	0.25	0.01	0.078	0.03
1H-4, 80–82	5.03	0.10		0.0016		1.36	0.02	426	9	0.426	0.009	0.075	0.04
2H-1, 30–32	6.80	0.17		0.0012		2.00	0.02	160	9	0.160	0.009	0.085	0.08
2H-2, 30–32	8.18	0.37		0.0026		4.08	0.02	463	9	0.463	0.009	0.091	0.06
2H-3, 37–39	9.66	0.43	0.00	0.0041	0.0003	4.94	0.04	512	9	0.512	0.009	0.087	0.08
2H-6, 60–62	13.21	0.17		0.0037		4.07	0.03	817	12	0.82	0.01	0.041	0.05
331-C0014G-													
1H-1, 0–2	0.00	0.48	0.00	0.0083	0.0004	3.80	0.02	398	7	0.398	0.007	0.127	0.21
1H-2, 20–22	1.58	0.18	0.00	0.0064	0.0003	4.45	0.03	498	7	0.498	0.007	0.040	0.13
1H-5, 10–12	5.56	0.13	0.00	0.0026	0.0002	4.12	0.04	484	8	0.484	0.008	0.032	0.05
2H-5, 45–47	15.05	0.21		0.0015		4.98	0.02	308	7	0.308	0.007	0.041	0.05
2H-7, 80-82	18.17	0.27	0.00	0.0015	0.0002	4.43	0.03	360	6	0.360	0.006	0.062	0.04
3H-2, 26–28	19.55	0.49	0.00	0.0032	0.0000	5.52	0.02	560	7	0.560	0.007	0.088	0.06
3H-5, 27–29	22.55	0.17		0.0024		4.30	0.02	242	7	0.242	0.007	0.039	0.10
3H-8, 116–118	25.85	0.05		0.0025		2.92	0.03	251	6	0.251	0.006	0.017	0.10
4H-5, 80–82	31.21	0.01		0.0219		4.56	0.04	1326	14	1.33	0.01	0.002	0.16
4H-7, 70–72	33.13	0.00	0.00	0.0260	0.0010	2.12	0.02	1786	12	1.79	0.01	0.001	0.15
4H-10, 60–62	36.05	0.01		0.0316		5.80	0.06	2510	22	2.51	0.02	0.002	0.13
5H-3, 45–47	38.49	0.02		0.0286		6.97	0.04	2173	8	2.173	0.008	0.003	0.13
6H-2, 38–40	47.64	0.03		0.0261		2.68	0.02	1150	7	1.150	0.007	0.011	0.23
6H-3, 80–82	48.51	0.02		0.0127		3.28	0.02	1479	8	1.479	0.008	0.006	0.09
9X-2, 64–66	56.25	0.02		0.0064		2.17	0.02	763	6	0.763	0.006	0.011	0.08
13T-1, 28–30	66.98	0.08		0.0112		2.65	0.02	1007	8	1.007	0.008	0.031	0.11
14T-2, 51–53	72.19	0.02	0.00	0.0088	0.0002	2.99	0.08	853	20	0.85	0.02	0.008	0.10
16T-1, 7–9	75.97	0.04	0.01	0.0049	0.0002	2.85	0.02	625	11	0.63	0.01	0.015	0.08
17T-2, 20–22	80.81	0.05	0.00	0.0066	0.0005	2.42	0.02	894	10	0.89	0.01	0.022	0.07
20T-1, 30–32	94.00	0.01		0.0019		0.54	0.02	207	9	0.207	0.009	0.011	0.09
21H-3, 79–81	99.90	0.01		0.0053		2.47	0.02	422	9	0.422	0.009	0.003	0.13
23X-1, 5–8	104.75	0.02		0.0028		1.36	0.02	182	9	0.182	0.009	0.013	0.16
24T-2, 90–92	110.60	0.02		0.0034		1.67	0.02	376	9	0.376	0.009	0.013	0.09
24T-4, 49.5–51.5	112.18	0.03		0.0042		1.64	0.02	460	9	0.460	0.009	0.018	0.09
25T-1, 28.5–30.5	113.99	0.02		0.0057		1.45	0.02	490	9	0.490	0.009	0.016	0.12
331-C0017A-													
1H-1, 80–90	0.80	0.71	0.02	0.0176	0.0000	4.01	0.02	468	7			0.178	0.38
1H-5, 88–98	6.52	0.71	0.02	0.0190	0.0001	3.80	0.02	472	6			0.188	0.40
331-C0017B-													
1H-2 85_95	11 03	0.63	0.02	0.0155	0 0003	3 79	0.02	497	8			0 167	0.31
1H-5 70_80	15.01	0.67	0.02	0.0703	0.0003	3.86	0.02	434	6			0.107	0.47
111 3, 7 0 00	15.01	0.07	0.01	0.0205	0.0001	5.00	0.02	151	Ũ			0.171	0.17
331-C0017C-	00	• • • •		0.04		a =-	c		-			0.1.1-	
1H-2, 60–70	20.26	0.46		0.0441		2.75	0.02	821	7			0.167	0.54
1H-5, 115–125	24.87	0.34	0.02	0.0095	0.0002	2.68	0.03	384	7			0.127	0.25
2H-1, 65–75	28.45	0.43		0.0043		1.71	0.02	528	7			0.251	0.08
2H-2, 65–75	29.85	0.13		0.0014		2.67	0.02	233	7			0.049	0.06
331-C0017D-													
1H-3 <i>,</i> 80–90	63.55	0.38		0.0009		3.50	0.05	883	2			0.109	0.01
1H-5, 80–90	66.30	0.52		0.0013		3.66	0.03	816	9			0.141	0.02
1H-6, 110–120	67.98	0.58		0.0009		3.30	0.05	299	8			0.176	0.03
2H-5, 30–40	75.08	0.61		0.0021		5.23	0.03	567	7			0.116	0.04
6X-1, 66–76	94.86	0.53	0.02	0.0259	0.0000	3.47	0.03	552	7			0.151	0.47
7H-2, 112–122	106.16	0.56		0.0269		4.60	0.03	571	8			0.122	0.47



Table T4 (continued).

Core, section, interval (cm)	Depth (mbsf)	Reactive Fe (%)	±	Reactive Mn (%)	±	Fe (%)	±	Mn (µg/g)	±	Mn (mg/g)	±	Fe _R /Fe _T	Mn _R /Mn _T
7H-4, 42–52	108.12	0.51	0.02	0.0242	0.0002	4.87	0.03	549	7			0.106	0.44
7H-5, 80–90	109.85	0.27		0.0119		2.99	0.01	402	9			0.092	0.30
9X-5, 25–35	125.72	0.41		0.0235		3.84	0.04	534	9			0.106	0.44
9X-8, 75–90	130.04	0.47		0.0277		4.19	0.04	581	7			0.112	0.48
10X-1, 88–102	131.58	1.00	0.01	0.0455	0.0002	4.95	0.03	792	9			0.201	0.57
10X-4, 80–90	135.95	0.80	0.01	0.0515	0.0001	4.79	0.02	889	8			0.167	0.58
11X-1, 80–90	141.00	0.68		0.0426		4.33	0.03	826	7			0.158	0.52
PACS-2 sediment		0.81	0.03	0.0030	0.0002								
Laboratory standard		0.99	0.10	0.0025	0.0002								

