
Data report: grain size distribution of unconsolidated sands offshore Shimokita Peninsula, Japan (IODP Site C0020)¹

Stephen C. Phillips^{2, 3} and Joel E. Johnson³

Chapter contents

Abstract	1
Introduction	1
Materials and methods	2
Results	3
Acknowledgments	3
References	3
Figures	5
Table	9

Abstract

We report particle size analyses measured from 28 unconsolidated sand samples recovered during Integrated Ocean Drilling Program Expedition 337 from Hole C0020A. These samples span from 1277 to 2002 meters below seafloor (mbsf) across two lithostratigraphic units that record a transition from a nearshore estuarine/intertidal to an offshore hemipelagic paleoenvironment. Most recovered lithologies, including coarse-grained lithologies, are semi-consolidated to consolidated; however, some intervals of sand are unconsolidated and suitable for particle size analysis. Bulk sand samples were measured using a laser diffraction particle size analyzer. All samples fall within the range of sand to silty sand, with silty sand more abundant deeper than 1925 mbsf. Samples contain 49%–97% sand, 3%–42% silt, and 0%–13% clay. Median grain diameter ranges from 55 to 405 μm . Clay and silt content in these sands reach a minimum at ~1500 mbsf and then increase below 1925 mbsf in a coal-bearing unit containing shale, siltstone, sandstone, and unconsolidated sand.

Introduction

Integrated Ocean Drilling Program Expedition 337 explored the microbiology and hydrocarbon system associated with subsurface coal beds in the Hidaka Trough offshore the Shimokita Peninsula, Japan (see the “[Expedition 337 summary](#)” chapter [Expedition 337 Scientists, 2013a]) (Inagaki et al., 2015, 2016). Hole C0020A (41°10.5983'N, 142°12.0328'E; 1180 m water depth) was drilled to 2466 meters below seafloor (mbsf) as an extension of Japan Agency for Marine-Earth Science and Technology (JAMSTEC) Site C9001, which was originally drilled to 647 mbsf (Aoike, 2007). Spot coring began at 1276.5 mbsf at an interval of approximately one core every 100 m, and near-continuous coring occurred between 1919.0 and 2002.4 mbsf across the coal-bearing target. Cuttings were collected over the entire hole at 10 m intervals but were not used in this study.

The recovered sediment sequence revealed a transition from offshore marine to nearshore/terrestrial environments that was described in four lithostratigraphic units (Figure F1) (see the “[Site C0020](#)” chapter [Expedition 337 Scientists, 2013b]) that span the late Oligocene to early Pleistocene (Phillips et al., 2016). Unit I (647–1256.5 mbsf; no core recovery) is composed of a hemi-

¹Phillips, S.C., and Johnson, J.E., 2017. Data report: grain size distribution of unconsolidated sands offshore Shimokita Peninsula, Japan (IODP Site C0020). In Inagaki, F., Hinrichs, K.-U., Kubo, Y., and the Expedition 337 Scientists, *Proceedings of the Integrated Ocean Drilling Program, 337*: Tokyo (Integrated Ocean Drilling Program Management International, Inc.).

doi:10.2204/iodp.proc.337.203.2017

²Department of Earth Sciences, University of New Hampshire, Durham NH 03824, USA.

Correspondence author:
phillips.stephen@gmail.com

³Institute for Geophysics, Jackson School of Geosciences, University of Texas at Austin, Austin TX 78712, USA.



pelagic, diatom-bearing silty clay, representing a continental slope environment. Unit II (1256.5–1826.5 mbsf; 7% core recovery) is composed of primarily silty shale with common intervals of siltstone and sandstone/unconsolidated sands. Unit II contains an increasing frequency with depth of coarse-grained sediments, terrestrial plant debris, bioturbation, and cross-lamination, along with a decrease in diatom abundance, suggesting a continental shelf to intertidal environment. *Cruziana* ichnofacies were observed in Unit II. Well logging results indicate a massive 60–70 m thick sandstone in Unit II and three coal layers (0.3–0.6 m thick) that were not sampled by coring. Unit III (1826.5–2046.5 mbsf; 35% recovery) contains multiple coal beds with intervals of sandstone/sand, siltstone, and shale. Flaser bedding, cross bedding, and authigenic carbonate nodules are present in Unit III. Sharp increases in magnetic susceptibility driven by authigenic magnetite precipitation are present in Unit III (Phillips et al., 2017). Unit III represents a nearshore environment with coals derived from brackish wetlands at the top of Unit III and freshwater wetlands at the base of Unit III (Gross et al., 2015). Unit IV (2046.5–2466 mbsf; 8% core recovery) consists of shale, sandstone, and siltstone, representing a tidal flat to fluvial environment, and is largely devoid of coal except for a 30 cm coal bed at the very bottom of the hole. Unit IV also contains numerous authigenic carbonate nodules.

Particle size distribution in marine and terrestrial sediments can provide insight into paleoenvironmental conditions driven by climatic and tectonic changes (e.g., Orton and Reading, 1993; Goman and Wells, 2000; Ding et al., 2002; Warner and Domack, 2002; Asikainen et al., 2007). In addition, grain size can influence microbial biomass and microbial activity and diversity in a range of marine and terrestrial environments (e.g., Dale, 1974; Bender and Conrad, 1994; Sessitsch et al., 2001; Noffke et al., 2002). Thus, grain size analysis of sediments at Site C0020 has the potential to provide insight on paleoenvironmental and microbial processes offshore Shimokita Peninsula since the late Oligocene.

A complete downcore sampling strategy for grain size analysis from each unit and throughout the entire hole was not possible because most cored samples from Site C0020 were consolidated or semi-consolidated, including all samples of siltstone and finer grained lithologies. These consolidated or semi-consolidated sediments cannot be accurately analyzed because intense disaggregation imparts error into the grain size distribution (Gealy, 1971; Thayer et al., 1974). However, multiple intervals of uncon-

solidated sands were recovered across Units II and III. No cores were collected in Unit I (only cuttings), and all sediments in Unit IV were lithified. Accordingly, we report the results of particle size analysis of 28 unconsolidated sand samples from Units II and III.

Materials and methods

All measured samples were unconsolidated sands based on visual core description and required no crushing to loosen individual grains. We measured bulk, untreated samples containing all lithogenic, biogenic, and authigenic components. These 28 unconsolidated sand samples were selected out of a set of 144 samples originally collected for a separate rock magnetic and geochemical study (Phillips et al., 2017) that covers the full range of major lithologies. The samples measured here are representative of the unconsolidated sands recovered from Hole C0020A. Collection of these samples was conducted to avoid contamination of drilling mud (Figure F2).

Each sample was treated with 20 mL of a 5.0 g/L sodium hexametaphosphate dispersing agent. Each sample was then agitated using a vortex mixer, left for a minimum of 24 h, and then agitated again immediately before analysis.

All samples were measured using Malvern Mastersizer 2000 laser diffractometer particle size analyzer with a Hydro 2000 wet dispersion unit. Our method is based on the approach of Sperazza et al. (2004). Grain size measurements were made using a particle refractive index of 1.55, a dispersant refractive index of 1.33, and a particle absorption index of 0.01.

All samples were passed through a 2 mm (#10 mesh) sieve before entering the dispersion unit, and no material >2 mm was observed on the sieve in any sample. Each sample was completely dispersed in the dispersion unit to avoid subsampling errors, which generally have a higher impact on sample uncertainty than obscuration rate (Sperazza et al., 2004). Samples in sodium hexametaphosphate were added to the dispersion unit and diluted to obtain an obscuration between 8% and 22% (average = 14%). The pump speed was run at 2000 rpm, the stirrer was run at 770 rpm, and the sample was sonicated for 10 s prior to the sample analysis.

Grain size distributions are reported as frequencies in 89 diameter bins from 0.01 to 2187 μm . Median, tenth percentile ($d(0.1)$), and ninetieth percentile ($d(0.9)$) were calculated from this distribution using the accompanying Mastersizer software. Mean grain size was calculated as both a surface area moment mean (Sauter mean diameter or $D[3,2]$) and a vol-

ume moment mean (De Brouckere mean diameter or $D[4,3]$). We calculate sand ($>63\ \mu\text{m}$), silt ($4\text{--}63\ \mu\text{m}$), and clay ($<4\ \mu\text{m}$) from each sample's grain size distribution based on the classification of Wentworth (1922).

The Mastersizer 2000 was calibrated with a Malvern glass sphere standard with mean diameter ($d(0.5)$) of $61\ \mu\text{m}$, tenth percentile ($d(0.1)$) of $37\ \mu\text{m}$, and ninetyth percentile ($d(0.9)$) of $90\ \mu\text{m}$. A local, well-sorted sand from Wallis Sands Beach in Rye, New Hampshire (US), was run four times throughout the run of samples as a check standard for consistency. This check standard repeated to within $<1\ \mu\text{m}$ for all mean and percentile statistics: $D[3,2] = 242\ \mu\text{m}$, $D[4,3] = 262$, $d(0.1) = 175\ \mu\text{m}$, $d(0.5) = 251\ \mu\text{m}$, and $d(0.9) = 362\ \mu\text{m}$.

One sample (337-C0020A-5R-3, 36–39 cm) was run in duplicate as a measure of repeatability across the range of obscuration (9% and 17%).

Results

All samples are primarily sand sized with the primary peak of the distribution in the fine-to-medium sand range (Figures F3, F4; see SAMPLES in **Supplementary material**). Each sample has a low and variable amount of fine-grained material represented by a long tail on the left of the distribution. Each sample from Units II and III falls within the range of 49%–97% sand, 3%–42% silt, and 0%–13% clay (Table T1). Based on the classification system of Shepard (1954), all samples fall within the range of sand or silty sand, with silty sand samples more common in Unit III (Figure F5). Median grain diameter $d(0.5)$ ranges from 55 to $405\ \mu\text{m}$. Surface area-weighted mean diameter $D[3,2]$ ranges from 9 to $178\ \mu\text{m}$, and volume-weighted mean diameter $D[4,3]$ ranges from 77 to $400\ \mu\text{m}$.

The overall coarsest unconsolidated material recovered from Hole C0020A is found in the middle of Unit II between 1377 and 1630 mbsf, indicated by the highest percent sand, lowest percent silt and clay, and highest values of $d(0.1)$ (Figure F6). Within Unit III, the percent sand decreases downhole, and silt and clay contents increase relative to Unit II. Median grain size decreases in Unit III corresponding to the increase in clay content. Sand samples from the upper 30 m of Unit III (Figure F7) in general have lower clay content, higher silt content, and higher $d(0.1)$ compared to the lower intervals of Unit III.

Sample 337-C0020A-5R-3, 36–39 cm, run in duplicate, shows the sand-silt-clay percentages are repeat-

able to within 1% and $d(0.5)$ matches within $5\ \mu\text{m}$. $d(0.1)$ and $d(0.9)$ repeat within 6 and $20\ \mu\text{m}$, respectively. $D[4,3]$ and $D[3,2]$ repeat within 6 and $13\ \mu\text{m}$, respectively.

Our grain size measurement from the uppermost sample in Unit III (1925.21 mbsf) is consistent with the grain size distribution of one sample from 1925.38 mbsf measured via sieve analysis by Ijiri et al. (2017). Overall, these results show systematic changes downhole that may provide insight to paleoenvironmental conditions and the role of the host material for microbial activity.

Acknowledgments

We thank the Captain and crew of the D/V *Chikyu* and the IODP Expedition 337 scientists and technicians for core collection and initial onboard analysis. We are grateful for the support of IODP and the Ministry of Education, Culture, Sports, Science and Technology of Japan. Samples and/or data were provided by IODP. We appreciate the review and comments from John Jaeger.

References

- Aoike, K. (Ed.), 2007. *CDEX Laboratory Operation Report: CK06-06 D/V Chikyu shakedown cruise offshore Shimokita*: Yokohama (CDEX-JAMSTEC).
- Asikainen, C.A., Francus, P., and Brigham-Grette, J., 2007. Sedimentology, clay mineralogy and grain-size as indicators of 65 ka of climate change from El'gygytyn Crater Lake, Northeastern Siberia. *Journal of Paleolimnology*, 37(1):105–122. <https://doi.org/10.1007/s10933-006-9026-5>
- Bender, M., and Conrad, R., 1994. Methane oxidation activity in various soils and freshwater sediments: occurrence, characteristics, vertical profiles, and distribution on grain size fractions. *Journal of Geophysical Research: Atmospheres*, 99(D8):16531–16540. <https://doi.org/10.1029/94JD00266>
- Dale, N.G., 1974. Bacteria in intertidal sediments: factors related to their distribution. *Limnology and Oceanography*, 19(3):509–518. <https://doi.org/10.4319/lo.1974.19.3.0509>
- Ding, Z.L., Derbyshire, E., Yang, S.L., Yu, Z.W., Xiong, S.F., and Liu, T.S., 2002. Stacked 2.6-Ma grain size record from the Chinese loess based on five sections and correlation with the deep-sea $\delta^{18}\text{O}$ record. *Paleoceanography*, 17(3):1033. <https://doi.org/10.1029/2001PA000725>
- Expedition 337 Scientists, 2013a. Expedition 337 summary. In Inagaki, F., Hinrichs, K.-U., Kubo, Y., and the Expedition 337 Scientists, *Proceedings of the Integrated Ocean Drilling Program*, 337: Tokyo (Integrated Ocean

- Drilling Program Management International, Inc.). <http://dx.doi.org/10.2204/iodp.proc.337.101.2013>
- Expedition 337 Scientists, 2013b. Expedition 337 Site C0020. In Inagaki, F., Hinrichs, K.-U., Kubo, Y., and the Expedition 337 Scientists, *Proceedings of the Integrated Ocean Drilling Program*, 337: Tokyo (Integrated Ocean Drilling Program Management International, Inc.). <http://doi.org/10.2204/iodp.proc.337.103.2013>
- Gealy, E.L., 1971. Grain size of sediments from the western equatorial Pacific: Leg 7, *Glomar Challenger*. In Winterer, E.L. et al., *Initial Reports of the Deep Sea Drilling Project*, 7: Washington DC (U.S. Govt. Printing Office), 1027–1036. <https://doi.org/10.2973/dsdp.proc.7.122.1971>
- Goman, M., and Wells, L., 2000. Trends in river flow affecting the northeastern reach of the San Francisco Bay Estuary over the past 7000 years. *Quaternary Research*, 54(2):206–217. <https://doi.org/10.1006/qres.2000.2165>
- Gross, D., Bechtel, A., and Harrington, G.J., 2015. Variability in coal facies as reflected by organic petrological and geochemical data in Cenozoic coal beds offshore Shimokita (Japan)—IODP Exp. 337. *International Journal of Coal Geology*, 152(B):3–79. <http://doi.org/10.1016/j.coal.2015.10.007>
- Ijiri, A., Ikegawa, Y., and Inagaki, F., 2017. Data report: permeability of ~1.9 km-deep coal-bearing formation samples off the Shimokita Peninsula, Japan. In Inagaki, F., Hinrichs, K.-U., Kubo, Y., and the Expedition 337 Scientists, *Proceedings of the Integrated Ocean Drilling Program*, 337: Tokyo (Integrated Ocean Drilling Program Management International, Inc.). <https://doi.org/10.2204/iodp.proc.337.202.2017>
- Inagaki, F., Hinrichs, K.-U., Kubo, Y., Bowles, M.W., Heuer, V.B., Long, W.-L., Hoshino, T., Ijiri, A., Imachi, H., Ito, M., Kaneko, M., Lever, M.A., Lin, Y.-S., Methé, B.A., Morita, S., Morono, Y., Tanikawa, W., Bihan, M., Bowden, S.A., Elvert, M., Glombitza, C., Gross, D., Harrington, G.J., Hori, T., Li, K., Limmer, D., Liu, C.-H., Murayama, M., Ohkouchi, N., Ono, S., Park, Y.-S., Phillips, S.C., Prieto-Mollar, X., Purkey, M., Riedinger, N., Sanada, Y., Sauvage, J., Snyder, G., Susilawati, R., Takano, Y., Tasumi, E., Terada, T., Tomaru, H., Trembath-Reichert, E., Wang, D.T., and Yamada, Y., 2015. Exploring deep microbial life in coal-bearing sediment down to ~2.5 km below the ocean floor. *Science*, 349(6246):420–424. <http://doi.org/10.1126/science.aaa6882>
- Inagaki, F., Hinrichs, K.-U., Kubo, Y., and the Expedition 337 Scientists, 2016. IODP Expedition 337: Deep Coal-bed Biosphere off Shimokita—microbial processes and hydrocarbon system associated with deeply buried coal-bed in the ocean. *Scientific Drilling*, 21:17–28. <http://dx.doi.org/10.5194/sd-21-17-2016>
- Noffke, N., Knoll, A.H., and Grotzinger, J.P., 2002. Sedimentary controls on the formation and preservation of microbial mats in siliciclastic deposits: a case study from the upper Neoproterozoic Nama Group, Namibia. *PALAIOS*, 17(6):533–544. [https://doi.org/10.1669/0883-1351\(2002\)017<0533:SCOTFA>2.0.CO;2](https://doi.org/10.1669/0883-1351(2002)017<0533:SCOTFA>2.0.CO;2)
- Orton, G.J., and Reading, H.G., 1993. Variability of deltaic processes in terms of sediment supply, with particular emphasis on grain size. *Sedimentology*, 40(3):475–512. <https://doi.org/10.1111/j.1365-3091.1993.tb01347.x>
- Phillips, M.P., Harwood, D.M., and Harrington, G.J., 2016. Neogene and early Pleistocene diatom biostratigraphy and age synthesis of Site C9001/C0020, northwest Pacific. *Marine Micropaleontology*, 128:39–49. <https://doi.org/10.1016/j.marmicro.2016.08.002>
- Phillips, S.C., Johnson, J.E., Clyde, W.C., Setera, J.B., Maxbauer, D.P., Severmann, S., and Riedinger, N., 2017. Rock magnetic and geochemical evidence for authigenic magnetite formation via iron reduction in coal-bearing sediments offshore Shimokita Peninsula, Japan (IODP Site C0020). *Geochemistry, Geophysics, Geosystems*, 18(6):2076–2098. <https://doi.org/10.1002/2017GC006943>
- Sessitsch, A., Weilharter, A., Gerzabek, M.H., Kirchmann, H., and Kandeler, E., 2001. Microbial population structures in soil particle size fractions of a long-term fertilizer field experiment. *Applied and Environmental Microbiology*, 67(9):4215–4224. <https://doi.org/10.1128/AEM.67.9.4215-4224.2001>
- Shepard, F.P., 1954. Nomenclature based on sand-silt-clay ratios. *Journal of Sedimentary Research*, 24(3):151–158. <http://dx.doi.org/10.1306/D4269774-2B26-11D7-8648000102C1865D>
- Sperazza, M., Moore, J.N., and Hendrix, M.S., 2004. High-resolution particle size analysis of naturally occurring very fine-grained sediment through laser diffractometry. *Journal of Sedimentary Research*, 74(5):736–743. <http://dx.doi.org/10.1306/031104740736>
- Thayer, P.A., Hostettler, J., and Smith, S., 1974. Grain-size distribution of sediments from the eastern Indian Ocean: Deep Sea Drilling Project, Leg 27. In Veevers, J.J., Heirtzler, J.R., et al., *Initial Reports of the Deep Sea Drilling Project*, 27: Washington, DC (U.S. Govt. Printing Office), 507–522. <http://dx.doi.org/10.2973/dsdp.proc.27.123.1974>
- Warner, N.R., and Domack, E.W., 2002. Millennial- to decadal-scale paleoenvironmental change during the Holocene in the Palmer Deep, Antarctica as recorded by particle size analysis. *Paleoceanography*, 17(3):8004. <http://dx.doi.org/10.1029/2000PA000602>
- Wentworth, C.K., 1922. A scale of grade and class terms for clastic sediments. *Journal of Geology*, 30(5):377–392. <http://dx.doi.org/10.1086/622910>

Initial receipt: 9 December 2016

Acceptance: 3 October 2017

Publication: 11 December 2017

MS 337-203

Figure F1. Lithostratigraphic summary with coal bed depths, wood/lignite content, diatom content, depositional environment, lithostratigraphic unit, and biostratigraphic age. Modified from Expedition 337 Scientists (2013b) with updated biostratigraphic ages from Phillips et al. (2016).

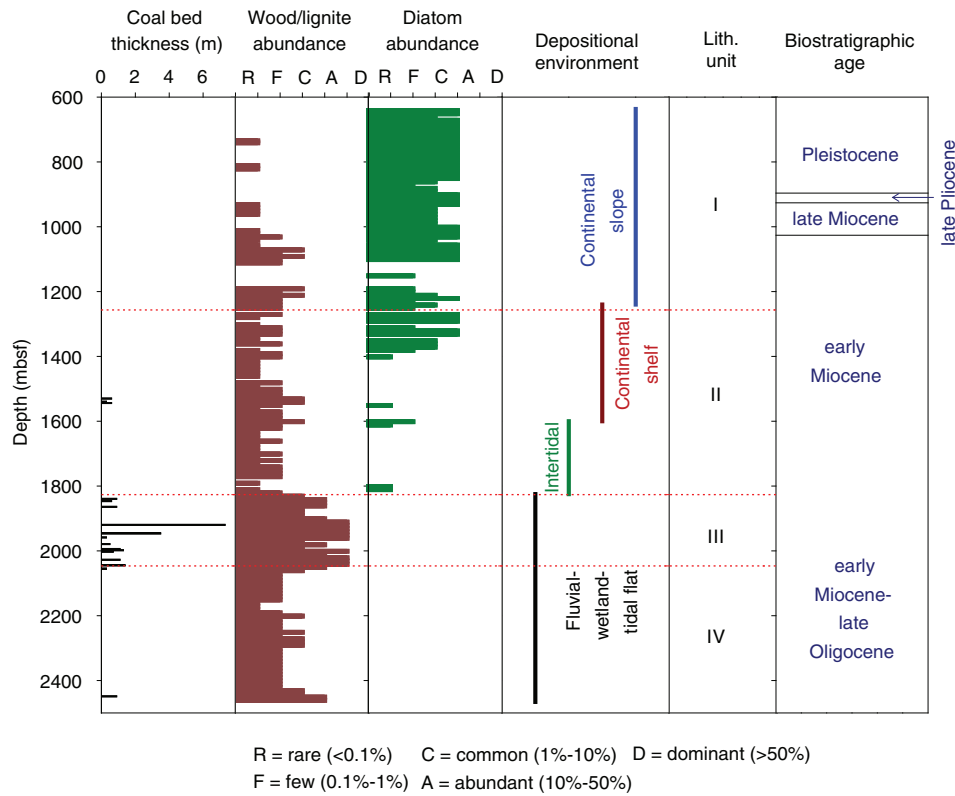


Figure F2. Core images showing the location of selected samples collected for this study. Samples were selected to avoid contamination by drilling mud. Images from Expedition 337 Scientists (2013b).

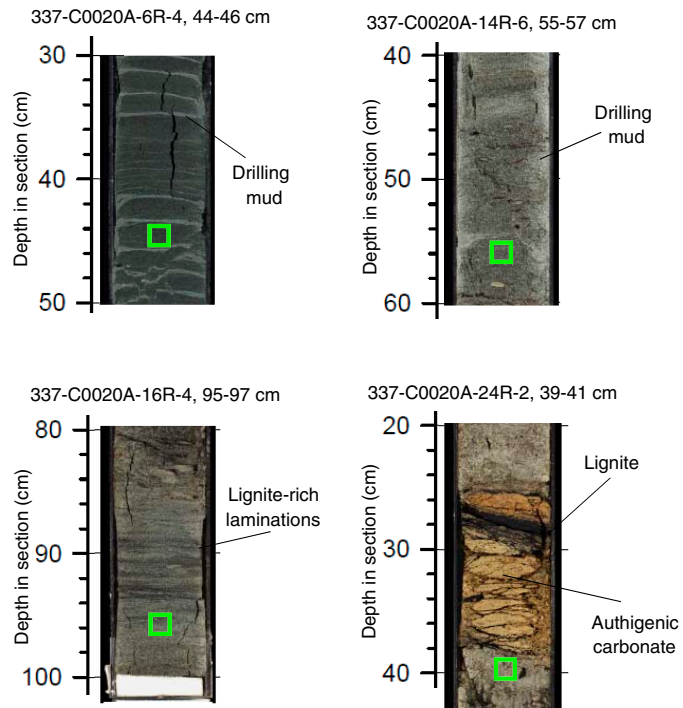


Figure F3. Particle size distribution for unconsolidated sands in Unit II.

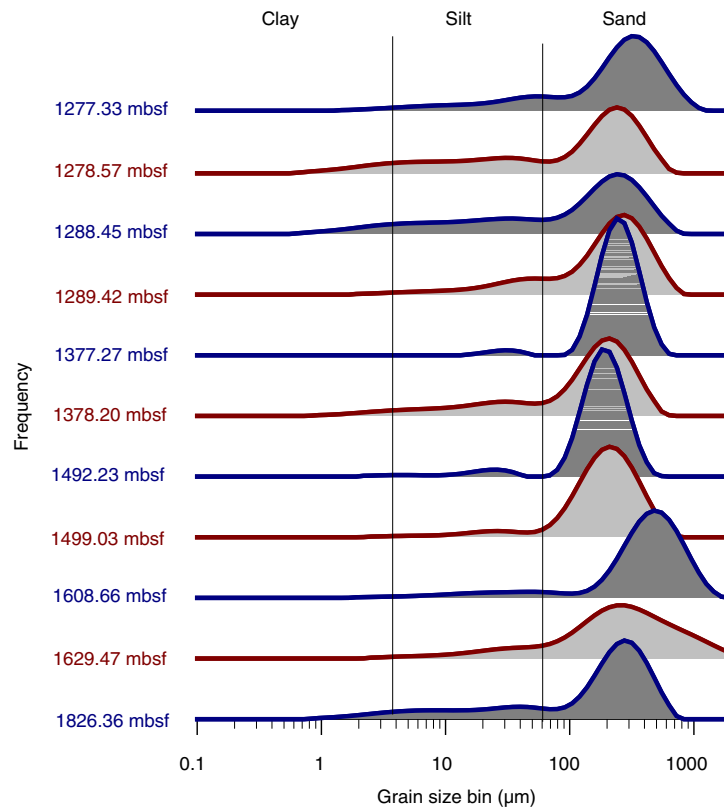


Figure F4. Particle size distribution for unconsolidated sands in Unit III.

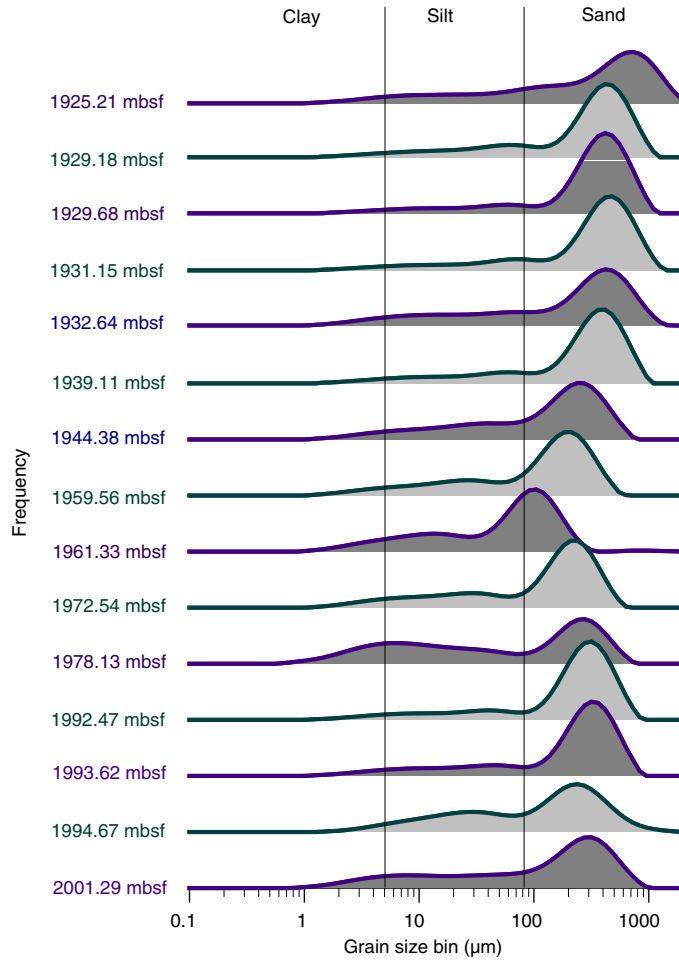


Figure F5. Sand-silt-clay percentages plotted on a ternary diagram based on the classification system of Shepard (1954).

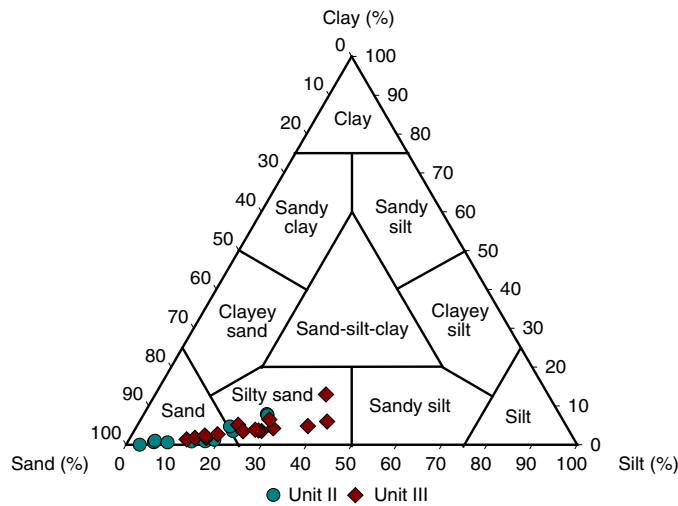


Figure F6. Downhole variation in summary statistics for unconsolidated sand samples in Hole C0020A.

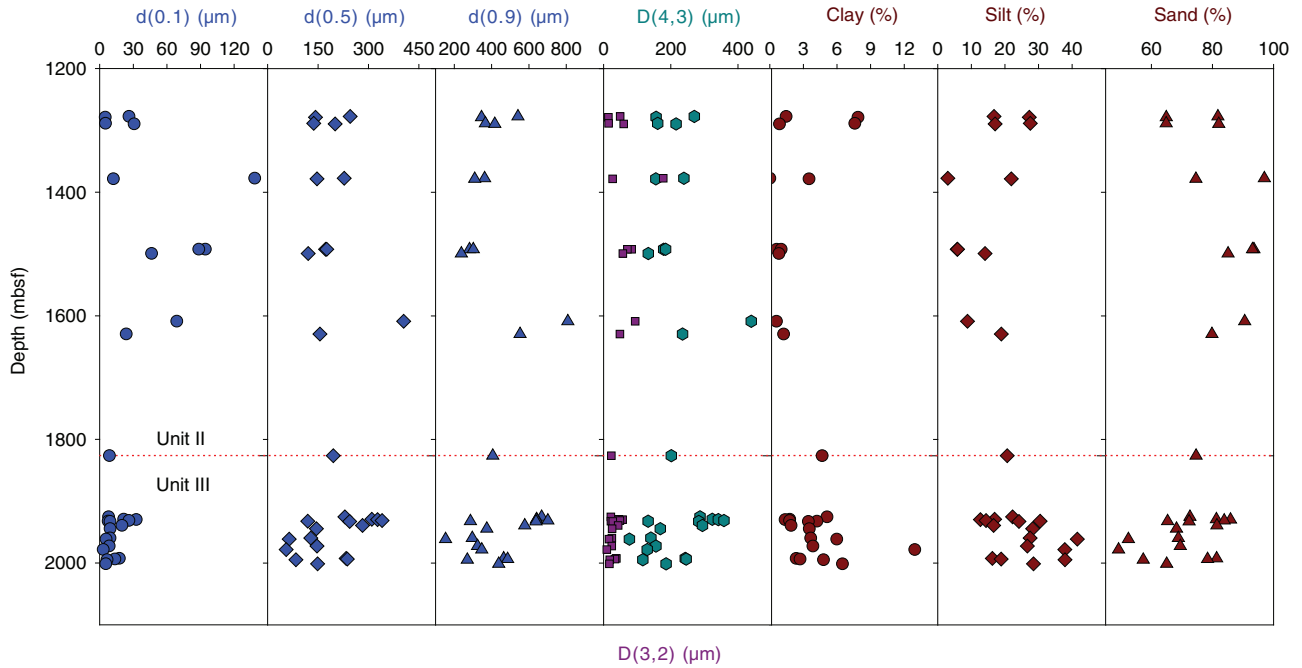


Figure F7. Downhole variation in summary statistics for unconsolidated sand samples in Unit III, Hole C0020A.

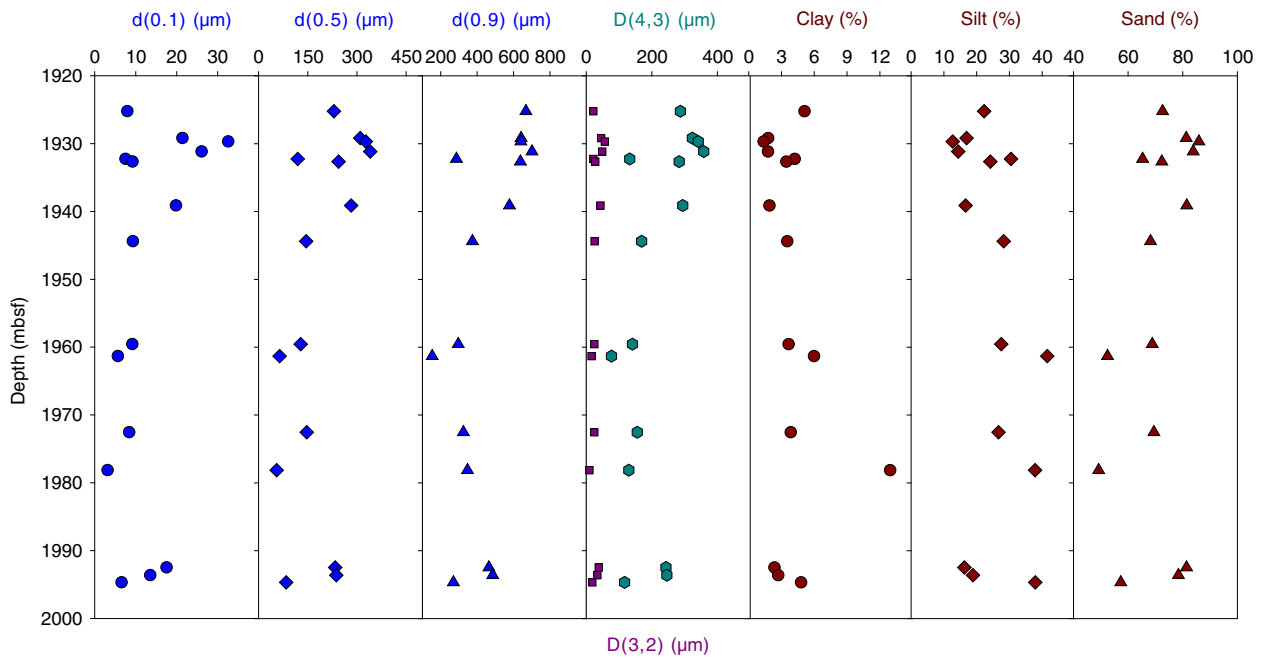


Table T1. Grain size distribution of Hole C0020A samples.

Core, section, interval (cm)	Depth (mbsf)	Lith. unit	D(4,3) (µm)	D(3,2) (µm)	d(0.1) (µm)	d(0.5) (µm)	d(0.9) (µm)	Clay (%)	Silt (%)	Sand (%)
337-C0020A-										
1R-1, 82-84	1277.33	II	270	49	26	246	541	1	17	82
1R-2, 79-51	1278.57	II	156	14	5	142	346	8	27	65
2R-2, 104-106	1288.45	II	161	15	5	137	367	8	28	65
2R-3, 60-65	1289.42	II	216	60	31	201	417	1	17	82
4R-2, 84-87	1377.27	II	239	178	138	228	363	0	3	97
4R-3, 74.5-77	1378.20	II	155	27	12	147	309	4	22	75
5R-3, 36.5-39	1492.23	II	179	83	94	172	282	1	6	94
5R-3, 36.5-39	1492.23	II	185	71	88	176	302	1	6	93
6R-4, 44-46	1499.03	II	133	57	46	121	238	1	14	85
8L-6, 49-51	1608.66	II	440	94	69	405	808	1	9	91
9R-4, 108-110	1629.47	II	235	48	24	156	552	1	19	80
14R-6, 55-57	1826.36	II	202	23	9	195	406	5	21	75
15R-6, 115-116	1925.21	III	287	21	8	230	668	5	22	73
16R-1, 67-68	1929.18	III	324	45	21	310	641	2	17	81
16R-2, 15-18	1929.68	III	342	56	33	327	640	1	13	86
16R-3, 86-88	1931.15	III	358	48	26	341	702	2	14	84
16R-4, 95-97	1932.25	III	132	22	8	119	286	4	30	65
16R-5, 32-34	1932.64	III	283	27	9	244	638	3	24	72
17R-4, 40-41	1939.11	III	294	43	20	282	578	2	17	81
17R-9, 22.4-25	1944.38	III	169	25	9	145	374	4	28	68
20R-1, 5-7	1959.56	III	141	24	9	128	296	4	28	69
20R-2, 41-43	1961.33	III	77	16	6	64	153	6	42	52
21R-CC, 2-4	1972.54	III	156	24	8	147	325	4	27	69
22R-5, 48-50	1978.13	III	130	9	3	55	347	13	38	49
24R-1, 45-47	1992.47	III	243	38	18	234	465	2	16	81
24R-2, 39-41	1993.62	III	246	34	14	237	486	3	19	78
24R-CC, 5-7	1994.67	III	117	18	7	84	270	5	38	57
25R-5, 129-131	2001.29	III	186	17	5	149	438	6	28	65